

1 **Title**

2 The ecosystem-climate-human nexus in the Arctic

3

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50

51

## 52 **Abstract**

53 The Arctic has warmed at nearly four times the global average since 1979, which has intensified  
54 the disruption of its biotic and local human communities under ongoing environmental change.

55 Here, we explore the ecosystem-climate-human nexus in the Arctic region. We summarize  
56 current knowledge of regional climate change and its impact on ecosystems and their functions,  
57 highlight gaps and uncertainties, and explore future outlooks to provide an overview of key  
58 areas for ongoing and future research. By detailing how the combination of biodiversity,  
59 environmental, and functional changes affect humans, we highlight the necessity of expanding  
60 climate change research to better incorporate environmental and social change, and predict  
61 ecological response, thereby increasing the resilience of Arctic communities.

62

## 63 **Keywords**

64 Biodiversity, Arctic, climate change, human impacts, terrestrial, tundra

## 65 Introduction

66

67 The Arctic region, which covers 5% of Earth's terrestrial surface (Meltotte et al. 2013), is in  
68 crisis. The greatest contemporary changes in both average and year-to-year variability of air  
69 temperature are occurring in the Arctic, which has already warmed at nearly four times the  
70 global average since 1979 (Rantanen et al. 2022). Climate change is already affecting marine  
71 and terrestrial biodiversity in the Arctic by driving changes in species' ranges and modifications  
72 of their phenotypes and life cycles across the Tree of Life (e.g. Pecl et al. 2017, Bjorkman et al.  
73 2018, van Beest et al. 2021). Further progression of climate change will cause increased  
74 permafrost thaw, retreat of glaciers and ice sheets, and decreased snow cover and sea ice  
75 (Intergovernmental Panel on Climate Change [IPCC] 2022), exacerbating these changes  
76 (Intergovernmental Science-Policy Platform on Biodiversity and Ecosystem Services [IPBES]  
77 2019). Arctic permafrost thaw and increased wildfire frequency has disrupted global carbon  
78 storage, accelerating the pace of environmental change (McCarty et al. 2020, Witze 2020).  
79 Permafrost thaw, extractive industries, and thinning and retreating glaciers and sea ice  
80 contribute to reduced land accessibility and loss of territorial rights of local human communities  
81 (Kumpula et al. 2011, Hanaček et al. 2022).

82

83 Livelihoods of Indigenous Peoples in the Arctic are based on nomadic herding and informed by  
84 millenia of Indigenous Knowledge, and weather instability under climate change is creating  
85 novel hurdles that cannot be navigated with previous experience (Shadrin 2021). Environmental  
86 change is increasing human access to natural resources in the Arctic, and heavy investments in  
87 economic activities including agriculture, maritime trade, natural resource exploration and  
88 extraction, immigration, and tourism are expected over the next decades (Barnhart et al. 2016,  
89 Constable et al. 2022), creating a high risk of social conflicts (Hanaček et al. 2022). Increased  
90 human population sizes and economic activities will introduce new risks associated with  
91 extreme weather events and warmer temperatures, such as damage to infrastructure from  
92 permafrost thaw. Damage to roads and seasonal pathways can also impact transport options  
93 and food accessibility for Indigenous Peoples and cause loss of connectivity between Arctic  
94 communities (Constable et al. 2022). Conversion of natural areas into cropland will also impact  
95 biodiversity and biogeochemical cycling, and further reduce the areas used for hunting, fishing,  
96 gathering, and reindeer (*Rangifer tarandus*) and muskox (*Ovibos moschatus*) pasture, i.e. the  
97 basis of society of many Indigenous Peoples (Mustonen and Shadrin 2021, Unc et al. 2021).

98

99 The multi-faceted changes in the Arctic will feed back on Earth system processes. Changes in  
100 the extent of snow, sea ice, permafrost, and glaciers all contribute to climate regulation at a  
101 global scale. Potentially irreversible and rapid non-linear changes of the Earth climate system  
102 (such as loss of permafrost and increased melting of the Greenland ice sheet) may occur under  
103 intermediate and high global warming scenarios (Peterson et al. 2020). As thinner ice  
104 dominates and spring snow cover and summer sea-ice cover decrease, more energy is  
105 absorbed at Earth's surface level, influencing the latitudinal temperature gradient and thereby  
106 global weather and climate (IPCC 2022). Permafrost thaw will cause drastic carbon sink-source  
107 shifts, and impact global climate (Callaghan and Jonasson 1995, Post et al., 2019). Yet,  
108 projections of responses to greenhouse gas emission scenarios vary widely, with potential for

109 slow to fast carbon cycle response in Arctic ecosystems (Schuur et al. 2022). Despite spatial  
110 variability in Arctic near-surface temperature projections, uncertainty in future socioeconomic  
111 pathways and emissions, and inconsistent estimation of future sea ice area and thickness,  
112 warming in the Arctic is projected to continue under all scenarios (Cai et al. 2021). While  
113 ecosystem response may buffer impacts of disturbance with increased plant community  
114 turnover, productivity, and greening, extensive abrupt permafrost thaw and limited plant  
115 response could lead to net positive greenhouse gas emissions (Schuur et al. 2022).

116  
117 Here, we define true tundra areas as High Arctic, Low Arctic, and high elevation northern tundra  
118 as Oroarctic, while Subarctic refers to low elevation regions above 55°N and below treeline,  
119 following the definitions of the Arctic Climate Impact Assessment (ACIA 2004), Virtanen et al.  
120 (2015) and Berner et al. (2024). Canada, the Kingdom of Denmark, Finland, Iceland, Norway,  
121 The Russian Federation, Sweden, and the United States are the eight Arctic States (**Figure 1**).  
122 Understanding and predicting changes in Arctic biodiversity caused by climate change poses a  
123 special problem, due to diverse political and legal priorities, a lack of representative ecological  
124 sampling and relatively few research stations across the Arctic (Schmidt et al. 2017a, Metcalfe  
125 et al. 2018). An uneven representation of the Arctic's two ecoregions (tundra and boreal forest,  
126 respectively) in monitoring programs limits opportunities to inform policy development and  
127 protection of Arctic biodiversity. For instance, due to difficulty obtaining data from Russia,  
128 exacerbated by the exclusion of Russia from the Arctic Council, data from stations in the Arctic's  
129 largest state are inaccessible (López-Blanco et al. 2024, Kasten et al. 2024). Nearly half of the  
130 International Network for Terrestrial Research and Monitoring in the Arctic stations occur in the  
131 Siberian Subarctic (López-Blanco et al. 2024), and the inaccessibility of this data results in a  
132 less comprehensive view of the impacts of climate change on ecosystem function and  
133 biodiversity. Further, circumarctic sites of long-term vegetation monitoring are notably limited  
134 (**Figure 1**), as only 45 monitoring sites across 32 unique locations (above 63°N) fit this criterion  
135 (Bjorkman et al. 2020). A restricted spatial coverage of monitoring stations and access to  
136 monitoring station data in the Arctic poses challenges in the understanding of all threats to  
137 Arctic biodiversity and ecosystems, as well as actions for conserving biodiversity in the region.

138  
139 Previous reviews have explored consequences of environmental change in the Arctic in different  
140 dimensions, including shifting ecological and trophic relationships (Post et al. 2009), impacts of  
141 changing phenology (Ernakovich et al. 2014, Prévay et al. 2017, Post et al. 2019, Prévay et al.  
142 2019), methane fluxes and shifts in carbon cycles and storage (Callaghan and Jonasson 1995,  
143 Post et al. 2019), as well as changes in biodiversity and species abundance (Callaghan and  
144 Jonasson 1995, Callaghan et al. 2004, Pecl et al. 2017, Taylor et al. 2020). Multiple studies  
145 provide synopses of warming treatment experiments, variability in species response to warming  
146 and climate sensitivity, as well as recent findings from long-term biodiversity monitoring  
147 programs (Elmendorf et al. 2012, Lehikoinen et al. 2014, Johnson et al. 2015, Taylor et al.,  
148 2020, Maes et al. 2024). Although these syntheses provide important summaries of our  
149 understanding of biodiversity and environmental change in the Arctic, it is less appreciated how  
150 these changes have influenced ecosystem function and how the combination of biodiversity,  
151 environmental, and functional changes affect humans.

152

153 The preceding sections highlight the entwinement of climate change, ecosystem functions, and  
154 human livelihoods in the Arctic, advocating for a nexus approach to understand their  
155 interconnectedness. This approach emphasizes the importance of identifying opportunities to  
156 mitigate threats, minimize trade-offs, and foster synergies (Liu et al. 2018). Here, we identify  
157 major knowledge gaps in our understanding of the nexus between climate, humans, and  
158 ecosystems in the Arctic region. We first identify what has been observed and is known about  
159 climate change and environmental change in the Arctic. We then outline knowledge gaps and  
160 uncertainties that scientists acknowledge. Many uncertainties are acknowledged that cannot be  
161 properly quantified; such deep uncertainties represent voids in our understanding as to how  
162 climate, humans, and ecosystems in the Arctic will respond. Finally, we detail future outlooks for  
163 the Arctic and consequential directions for future research. Due to the critical situation in the  
164 Arctic, our goal is not only to review these topics, but provide an overview of key areas of  
165 current and future research, creating a roadmap of the Arctic nexus.

166

## 167 **1. Climate change**

168

### 169 **1.1 Tipping points, feedbacks, and carbon dynamics**

170 Arctic summers are now the warmest they have been in 2,000 years, and air temperatures are  
171 expected to increase by 3°C (SSP1-2.6) to 12°C (SSP5-8.5) by 2081-2100 (Lee et al. 2021),  
172 with warming projected to be stronger over the Arctic Ocean than over land (e.g. ~5.9°C over  
173 the Arctic Ocean vs. ~4.6°C over Arctic land under the intermediate SSP2-4.5 scenario; Cai et  
174 al. 2021). Warming in the Arctic could lead to modification of multiple cryosphere and biosphere  
175 elements resulting in tipping points, or critical thresholds that once surpassed can qualitatively  
176 alter a system, often irreversibly (Lenton et al. 2008). Many uncertainties in climate projections  
177 associated with feedbacks from sea ice, ocean, and atmosphere interactions remain, as models  
178 are unable to consistently incorporate local feedbacks or simulate changes in spatial patterning  
179 of sea ice loss (Cai et al. 2021). One critical issue is how the Earth climate system will be  
180 impacted when elements like the Atlantic Meridional Overturning Circulation (AMOC) reach  
181 tipping points (van Westen et al. 2024). The AMOC is an integral component of ocean  
182 thermohaline circulation and North Atlantic climate regulation. Freshening of the North Atlantic  
183 Ocean by Arctic freshwater sources, notably from melting of the Greenland ice sheet, Canadian  
184 Arctic Archipelago glaciers, and Arctic sea ice, is influencing convection in the Labrador Sea  
185 and likely weakening the AMOC (Yang et al. 2016). The complexity of the AMOC system has  
186 led to widely opposing estimates of timing and degree of AMOC collapse, with some studies  
187 predicting a weakening and recovery after temperature stabilization of 1.5°C to 3°C of global  
188 warming (Jackson and Wood 2018, Sigmond et al. 2020) and others forecasting a potential full  
189 collapse as early as mid-century (Ditlevsen and Ditlevsen 2023, van Westen et al. 2024).

190

191 Global warming is expected to reach 1.5°C by the 2030s, even with drastic emission reductions  
192 (Lee et al. 2021). Exceeding the 1.5°C threshold potentially means a 3°C warmer Arctic, a shift  
193 of the Subarctic northwards, increased melt of the Greenland ice sheets, widespread permafrost  
194 thaw, and frequent loss of Arctic summer sea ice (Scheffer et al. 2012, Schuur et al. 2015,  
195 Hoegh-Guldberg et al. 2018, Armstrong McKay et al. 2022). Regardless of the emission  
196 scenario, recent projections have shown that the Arctic will be sea ice-free in September as

197 early as the 2030s-2050s (Kim et al. 2023). Furthermore, if global air temperatures warm to 2°C,  
198 Arctic temperatures could increase by 4-8°C or more in some regions. This could lead to  
199 substantial areas of permafrost degradation, which would seriously impact hydrology,  
200 ecosystems, and building and road infrastructure (Kokelj and Jorgenson 2013, Chadburn et al.  
201 2017, Biskaborn et al. 2019). Abrupt permafrost thaw, potentially leading to permafrost collapse,  
202 could impact global temperatures, especially if deeper carbon-rich permafrost (Yedoma  
203 deposits) are exposed (Schurr et al. 2015, Armstrong McKay et al. 2022). Permafrost thaw and  
204 subsequent surface water redistribution, leading to either drier or wetter conditions, can impact  
205 ecosystem carbon dynamics irreparably (Jorgenson 2013, Schuur et al. 2022).

206  
207 Phenological shifts such as earlier green-up and delayed senescence (Treharne et al. 2020)  
208 may influence biogeochemical processes (e.g. carbon sequestration and photosynthesis), and  
209 physical attributes (e.g. water balance and the surface energy balance) (Peñuelas et al. 2009)  
210 with direct and indirect climate feedback. At the biogeochemical level, increased carbon dioxide  
211 uptake from a longer period of photosynthetic activity would reduce warming from greenhouse  
212 gases by decreasing atmospheric carbon dioxide (Schuur et al. 2022), while increased  
213 temperatures and plant productivity increases ecosystem respiration rates (Maes et al. 2024).  
214 Increased plant activity and species composition change could also lead to changes in the  
215 strength and type of the emission of biogenic volatile organic compounds (such as terpenoids),  
216 which could enhance or counteract global warming due to aerosol formation (Peñuelas et al.  
217 2009, Tang et al. 2023). For instance, a cooling effect could be expected where evergreen  
218 conifers expand their ranges into the High Arctic and monoterpene emissions increase, while a  
219 warming feedback is expected where broad-leaved deciduous trees replace evergreens and  
220 monoterpene emissions decrease (Tang et al. 2023). The degree to which soil organic matter  
221 decomposition in permafrost may enhance global warming is uncertain and influenced by local  
222 conditions of soil saturation, freeze and thaw cycles (especially duration and soil layer depth of  
223 thaw), and movement of labile organic matter through the soil profile (Walz et al. 2017).

224  
225 Record high permafrost temperatures have already been recorded at multiple long-term  
226 monitoring sites (boreholes down to ~10-20m depth) around the Arctic (Biskaborn et al. 2019).  
227 As permafrost thaws, hydrological modifications become more common and the soil organic  
228 carbon stored for millenia becomes available for microbes to degrade. The amount of carbon in  
229 permafrost soils is estimated to be nearly double that of the atmosphere (Hugelius et al. 2014).  
230 Even partial loss of this carbon through microbial degradation as carbon dioxide and methane,  
231 can lead to drastic increases of greenhouse gases in the atmosphere (IPCC 2019, 2022). High  
232 emission scenarios (i.e. RCP 8.5) estimate a potential release of 5 to 15% of the carbon pool in  
233 Arctic permafrost by 2100, a feedback that would influence climate change on a similar scale as  
234 a major land-use change such as deforestation (Schuur et al. 2022). Carbon budgets limiting  
235 warming to 1.5°C need to be reduced to account for permafrost thaw, with further reduction to  
236 maintain temperature stabilization (Rogelj et al. 2018). Up to 5,300 teragrams of carbon from  
237 methane release and 240,000 teragrams of carbon from carbon dioxide from permafrost thaw  
238 may occur during the 21st century, but the timing and amount of greenhouse gases released  
239 from permafrost soils are still highly uncertain (Canadell et al. 2021). If climate mitigation goals  
240 (e.g. the Paris Agreement on climate change, the Kyoto Protocol) are achieved then northern

241 peatlands may remain net carbon dioxide sinks, but under RCP 8.5 increased carbon dioxide  
242 and methane emissions could augment warming by 0.21°C by 2300 (Qiu et al. 2022). Soil  
243 methane uptake can also reduce carbon release, as Arctic wetland net methane emissions  
244 predicted for 2100 increased only by 18% (29 to 35 teragrams of methane a year) after  
245 incorporating methane-oxidizing bacteria and methanogens dynamics (Oh et al. 2020). Yet, as  
246 atmospheric carbon dioxide concentration increases, a smaller fraction of atmospheric carbon  
247 will be stored in Arctic terrestrial and oceanic sinks as these sinks become saturated (Canadell  
248 et al. 2021).

249

## 250 **1.2 Changes in Arctic weather, wildfire frequency, and browning events**

251 Increases in winter precipitation and mean summer temperature are predicted under climate  
252 change in the Arctic (Leffler et al. 2016). Despite projected increases in precipitation over high  
253 latitudes, soil moisture shows high regional variability across warming levels, with both drier and  
254 wetter soils projected due to the influence of evapotranspiration and future regional permafrost  
255 conditions (Bring et al. 2016, Wrona et al. 2016). Fluctuations in terrestrial precipitation,  
256 evaporation, and soil moisture in the Arctic have contributed to increased tundra wildfire  
257 frequency, magnitude, and severity due to unprecedented drought, extreme weather events,  
258 and earlier snowmelt (Bokhorst et al. 2011, Wrona et al. 2016). The Arctic region is also  
259 experiencing increasingly extreme fire seasons, with recent large-scale fires across Alaska and  
260 Siberia (McCarty et al. 2021). Fires in the northern latitudes are driven by climate and fuel  
261 conditions, lightning, and human activity (McCarty et al. 2021). Wildfire seasons, particularly in  
262 tundra regions of Alaska, northwest Canada, and Siberia, have been significantly correlated to  
263 summer heatwave activity (Thoman et al. 2023, Hegedús et al. 2024). While the majority of  
264 global wildfires are intentionally or accidentally set by humans, lightning is the predominant  
265 cause of burning in Arctic regions (Veraverbeke et al. 2017). Lightning is also projected to  
266 become twice as frequent in the Arctic tundra than in the Subarctic, increasing by about 150%  
267 by 2100 (Chen et al. 2021a).

268

269 Warmer winters with both snow and rain, and more extreme weather events are expected  
270 (Walsh et al. 2020, Henry et al. 2022). However, the scarcity of controlled winter field  
271 experiments in the Arctic that manipulate single factors limits understanding of the impacts of  
272 warmer weather conditions and severe events (Bokhorst et al. 2023). Extreme winter warming  
273 events, ‘frost drought’, and rain-on-snow events which cause an ice layer at the surface of,  
274 within, or below the snowpack, may lead to destruction of living plant biomass, and result in  
275 browning events (reviewed by Phoenix et al. 2024). Extreme winter warming events can also  
276 have severe consequences for large mammals and small animals that live under the snow. For  
277 instance, a severe rain-on-snow event in the Canadian Arctic Islands region killed ca. 20,000  
278 muskoxen as they could not access food sources (Putkonen et al. 2009). Similar findings have  
279 been reported for reindeer both on Svalbard (Kohler and Aanes 2004) and Western Siberia  
280 (Forbes et al. 2016). Earlier snowmelt has been associated with spring phenology  
281 advancements of many organismal groups, and while thinner snowpack often increases  
282 mortality in plants via frost drought; impacts on animals are not as consistent or clear (Slatyer et  
283 al. 2022). Winter warming can result in a thinner snowpack, earlier snowmelt and earlier  
284 flowering, which can lead to increased exposure to freezing spring conditions and decreased

285 flower abundance, as found for *Cassiope tetragona* and *Salix arctica* in High Arctic Greenland  
286 (Wheeler et al. 2015). If flowering continues to occur earlier, plants will likely be exposed to  
287 freezing temperatures more often, which may further reduce flower production and negatively  
288 impact plant survival (Wheeler et al. 2015). Yet, the response of different plant species to earlier  
289 snowmelt varies widely, and it often takes multiple years of consistent earlier snowmelt for a  
290 response to become evident (Frei and HSenry 2021). Landscape heterogeneity and  
291 microclimate variables are also important to consider in terms of their potential influence on  
292 snow accumulation and snowmelt timing on phenological shifts.

293

294 While Arctic vegetation is expected to increase carbon storage through treeline expansion  
295 northward, increased photosynthesis, water use efficiency, unknowns such as drought  
296 occurrence, changes in land use, nitrogen availability, limits to migration rates, and wildfire  
297 prevalence could diminish carbon sink capacity (Canadell et al. 2021, Chen et al. 2021b,  
298 Gustafson et al. 2021). Further, decreased albedo from forest encroachment into tundra is  
299 expected to increase the rate of global warming due to positive biogeochemical feedbacks to  
300 the earth-atmosphere energy balance (Zhang et al. 2013). Shrub expansion constitutes one of  
301 the most noticeable climate-driven changes in tundra vegetation (Sturm et al. 2001, Elmendorf  
302 et al. 2012, Bjorkman et al. 2018, Vowles and Björk 2019), and landscape attributes could  
303 influence the direction of vegetation state transition (Chen et al. 2021b). For instance, in dry  
304 tundra, climate change and increased wildfire frequency could promote Arctic shrub expansion,  
305 whereas, in wet tundra, wildfires could counteract the climate-driven shrub expansion (Chen et  
306 al. 2021b). However, there remains high uncertainty regarding how multiple stressors (e.g.  
307 combination of climatic drivers and increased fires frequencies) will impact Arctic ecosystems.  
308 Will these stressors be additive, counteractive, or even drive Arctic ecosystems to alternative  
309 states?

310

## 311 **2. Human activity**

312

### 313 **2.1 Vulnerability of humans in the Arctic**

314 The Arctic region is home to almost four million people, of which 10% are Indigenous Peoples  
315 (Johnson et al. 2015, IPCC 2019, Constable et al. 2022). Population density and economic  
316 activity vary significantly across the eight Arctic States (**Figure 1**), ranging from large areas with  
317 no human settlements or industry, either on- and off-shore, to areas with larger cities and  
318 significant economic activities (European Environment Agency 2017).

319

320 Climate change increases the vulnerability of Arctic Indigenous communities by modifying  
321 resource availability and mobility within and between communities. Humans in the Arctic are  
322 highly vulnerable due to their dependency on limited natural resources (e.g. fish, reindeer, game  
323 birds, wild plants), which are sensitive to climate change (Meredith et al. 2019, Mustonen and  
324 Shadrin 2021, Fedewa et al. 2020, Jørgensen et al. 2019). For instance, the productivity and  
325 phenology of Pacific and Atlantic salmon (*Oncorhynchus* spp., *Salmo salar*) in freshwater and  
326 coastal areas in Alaska, Norway, and Finland have been altered due to warming and algal  
327 blooms (Brattland and Mustonen 2018, Cline et al. 2019, Mustonen et al. 2021). Declines in  
328 reindeer abundance and habitat (Vors and Boyce 2009), have led to the categorization of



329 reindeer as vulnerable on the IUCN Red List of Threatened Species (Gunn 2016). Road, harbor,  
330 and water treatment infrastructures are negatively impacted by climate-induced flooding,  
331 thawing permafrost, erosion, and sea level rise (Ford et al. 2021, Mustonen and Shadrin 2021).  
332 Less ice coverage has also allowed for increased shipping, growth in marine trade, tourism, and  
333 mineral and oil resource extraction, all which may lead to increased nitrogen deposition and  
334 pollution (Stephen 2018, Ford et al. 2021, Parmesan et al. 2022). Beyond physical risks for  
335 these local human communities, there is a risk for heightened political tensions playing out at  
336 the expense of vulnerable and marginalized communities that have historically had little to no  
337 voice in their governance (Dawson et al. 2018, Drewniak et al. 2018). Climate change can  
338 effectively erase important locations and traditions which keep Indigenous communities alive  
339 and united, and provide them a source of income through hunting and fishing tourism, for  
340 example (Fenger-Nielsen et al. 2020, Jensen 2020).

341  
342 Arctic Indigenous communities are facing multiple cascading and compounded risks, which  
343 creates an existential threat for these communities (Constable et al. 2022). Food insecurity in  
344 the Arctic is expected to increase, partly due to rising sea water temperature, influencing  
345 subsistence fish species to migrate (e.g. the brown trout (*S. trutta*) in Russia) and invasive  
346 marine species to expand their ranges (e.g. mackerel (*Scomber scombrus*) in Iceland)  
347 (Astthorsson et al. 2012, Pecl et al. 2017, Huntington et al. 2020, Mustonen et al. 2021). Other  
348 factors, such as global shifts in phytoplankton distribution and abundance in response to  
349 increased ocean temperature, acidity, and stratification also contribute to subsistence fish  
350 distribution patterns (Hoegh-Guldberg and Bruno 2010). Direct exposure to climate-induced  
351 hazards, coupled with food insecurity and an increasing rate of waterborne and vector-borne  
352 diseases (such as gastroenteritis and tularemia), and environmental contamination are  
353 projected to worsen the health of humans in the Arctic (Waits et al. 2018, Grigorieva 2024). For  
354 example, Arctic populations, Inuit in particular, have been exposed to elevated levels of mercury  
355 (Basu et al. 2022), which pose serious health effects. Exposure pathways are mainly through  
356 diet (Basu et al. 2022, Donaldson et al. 2010), and while concentrations of chemicals (like lead  
357 and mercury) have been declining in some areas of the Arctic, contaminants like selenium have  
358 been reported for the first time (Gibson et al. 2016, Abass et al. 2018). Estimated mercury  
359 release from permafrost thaw has been projected to increase mercury concentrations in the  
360 Yukon River, Canada by 14% to 50% by 2100, depending on the emission scenario (Schaefer  
361 et al. 2020). Most emerging disease outbreaks in the Arctic over the last 30 years have been  
362 vector-borne or zoonotic, both of which are indirectly affected by climate (Ruscio et al. 2015). As  
363 temperatures increase and permafrost thaws, factors such as renewed pathogen activity,  
364 increased pathogen survival, and vector range expansion increase transmission risk (Waits et  
365 al. 2018, Mohite et al. 2023). Due to decreases in water quantity and quality, associated with  
366 water treatment failures, water rationing, and the absence of indoor plumbing, an increase in  
367 waterborne disease outbreaks have been reported, including new, emerging waterborne  
368 pathogens (Thomas et al. 2016, Harper et al. 2020, Mustonen and Shadrin 2021). Increasing  
369 temperatures and precipitation are predicted to have the strongest impacts on promoting the  
370 spread of pathogens (Raheem 2018).

371

372 The loss of archeological sites, cultural sites, and cultural practices, such as ice-fishing and  
373 reindeer herding due to ice thaw, represents a major risk, and potentially an existential one, for  
374 the heritage and identity of Indigenous communities (Raheem 2018, Nicu and Fatorić 2023).  
375 Mental health and well-being of these communities are also at stake due to indirect exposure to  
376 risks heightened by climate change, such as loss of cultural heritage, place-based knowledge,  
377 and livelihoods (Cunsolo Willox et al. 2015, Harper et al. 2020). These health impacts are and  
378 will be unequally distributed among populations depending on age and gender (Kowalczewski  
379 and Klein 2018, Feodoroff 2021). Primary knowledge gaps of humans in the Arctic regard  
380 adaptation to climate change and how impacts from global markets will influence human-natural  
381 systems interactions (e.g. subsistence harvesting, increased transport and tourism, and natural  
382 resource development projects) (Kapsar et al. 2022). These interactions are complex and reliant  
383 on cooperation from the local to global level to be sustainable as climate change progresses.  
384 For instance, in many rural Indigenous communities of Alaska, subsistence harvest accounts for  
385 over 50% of the local diet (Fall 2016). The traditional practice of subsistence harvest is essential  
386 for cultural wellbeing and depends on healthy ecosystems and populations to continue  
387 sustainably. If subsistence harvest, and therefore food security, is compromised, an increase in  
388 migration from rural to urban areas is likely, which could lead to environmental degradation and  
389 impacts on local biodiversity as urban centers grow (Kapsar et al. 2022).

390

## 391 **2.2 Adaptation in local Arctic human communities**

392 Knowledge is still limited regarding how Indigenous communities have changed their practices  
393 and habits to match changing climate and environments. Adaptation in Arctic communities has  
394 mostly been behavioral and reactive, with few examples of transformational adaptation that, for  
395 instance, consider the concerns of Indigenous communities and integrate their solutions in a  
396 long-term and potentially more sustainable way (Canosa et al. 2020). The Alaska Native Tribal  
397 Health Consortium, a local environmental observation platform, is an example of  
398 transformational adaptation that is used for sharing information on environmental impacts and  
399 local human community health effects (Berner et al. 2016). In fact, climate change appears not  
400 to be explicitly included in planning and design in many sectors (Canosa et al. 2020), due to the  
401 limited participation of Indigenous communities in decision-making bodies, a manifestation of  
402 colonialism and historical inequities and injustices (Ford et al. 2021). The integration of local  
403 Arctic residents' knowledge and involvement in the framework, design, and production of  
404 human-natural system research is essential to create effective policy in a changing Arctic  
405 (Kapsar et al. 2022). Additionally, while certain adaptation practices may work in the present,  
406 they may result in maladaptive behavior in the future. For instance, due to reduced caribou  
407 availability, Inuit in the Northwest Territories of Canada have relied more on muskox for food, as  
408 well as income through guided hunts. This adaptation has led to a decrease in the population  
409 size and health of muskox (Fawcett et al. 2018). Evidence of successful adaptation,  
410 maladaptation, and limits to adaptation are still scarce, and represent major knowledge gaps  
411 which need to be investigated to understand where, how, and if participatory governance has  
412 included climate change impacts in planning to reduce exposure and risk to local populations.

413

## 414 **3. Ecosystem evolution**

415

416 **3.1 Evolution of the Arctic**

417 The biogeography of the Arctic regions has influenced the ecological and evolutionary  
418 trajectories of its biotic communities. The macrofossil record indicates that the Arctic tundra  
419 became a recognizable biotic community at the end of the Neogene (the Pliocene) into the early  
420 Pleistocene (ca. 3-2 Ma; Bennike and Böcher 1990, Matthews and Ovenden 1990, Murray  
421 1995), as cooling cycles and continental glaciation shifted the tundra from continuous forest to  
422 treeless landscapes (Repenning and Brouwers 1992, Murray 1995). Molecular data suggests  
423 that the Arctic flora assembled from dispersal events as old as 10 mya, with dispersal increasing  
424 over time as a function of climate and peaking at 1.0-0.7 mya (Zhang et al. 2023). A relatively  
425 homogenous tundra flora persisted until the Last Glacial Maximum (~20 kya), when increases in  
426 temperature and precipitation led to higher plant diversity as herbaceous and woody  
427 communities emerged during the early to mid-Holocene (MacDonald et al. 2000, Wang et al.  
428 2021). Currently, the Arctic flora includes 2,218 vascular species (Elven et al. 2011), and Low  
429 Arctic areas are generally situated within larger landmasses like the American and Eurasian  
430 continents. In contrast, most of the landmass in High Arctic regions above 70°N feature  
431 numerous large archipelagos. This geographical distinction holds implications for the  
432 susceptibility of ecosystems to novel species introductions and the pace of species turnover.  
433 However, there is a lack of studies exploring how the biogeographical patterns impact current  
434 biotic community change.

435  
436 Changes in phenology, such as the timing of vegetation green-up (or leaf emergence), and  
437 abundance are commonly used indicators of species and ecosystem response to climate  
438 change, due to their responsiveness to warming and wide-ranging impacts on ecological  
439 processes (Peñuelas et al. 2009). Earlier green-up trends across the Arctic have been well-  
440 supported with long-term ecological research plot observations, remote sensing, and warming  
441 treatments (Bjorkman et al. 2020, Jenkins et al. 2020, Collins et al. 2021). Yet, vegetation  
442 green-up and abundance response is often not detectable, is species- and/or site-specific, or  
443 dependent on specific temperature windows (Myers-Smith et al. 2015, Bjorkman et al. 2020,  
444 Scharn et al. 2021), thus often making species response over time difficult to predict. Arctic  
445 animals that can respond phenologically, such as modifying the timing of breeding events or  
446 migration, may have an advantage under climate change (Gilg et al. 2012). For instance, a  
447 colony of the Arctic seabird, Mandt's black guillemot (*Cepphus grylle mandtii*), advanced its egg  
448 lay date by nearly eight days between 1976 and 2017 on the northern coast of Alaska (Sauve et  
449 al. 2019). Interestingly, long-distance migrant Arctic shorebird birds, such as the Western  
450 sandpiper (*Calidris mauri*) and Red Phalarope (*Phalaropus fulicarius*), have shown greater  
451 phenological response to spring green-up than short- and medium-distance migrants (Tavera et  
452 al. 2024).

453  
454 Due to warming and environmental modification at northern range limits, many Arctic species  
455 are expanding their ranges northward and/or to higher elevations. Continued northward  
456 expansion of generalist Arctic bird species and increases in herbivorous Arctic bird taxa will  
457 likely further degrade habitat quality, cause vegetation loss, and potentially induce trophic  
458 cascades (Davey et al. 2013, Smith et al. 2020). Arctic species and genotypes at the northern  
459 edge of their ranges are especially at risk of extirpation when further poleward and/or higher

460 elevation range expansion is not possible. In the Swedish Scandes mountain birch (*Betula*  
461 *pubescens* ssp. *tortuosa*) saplings were found ca. 400m higher than the species' uppermost  
462 limit in 1955 (Kullman 2002), though stunted growth of birch saplings above treeline suggest  
463 that tundra plant communities are not yet shifting to birch forest (Scharn et al. 2022). Non-Arctic  
464 species are also expanding their ranges northward. Range expansion of the red fox (*Vulpes*  
465 *vulpes*) into the southern range limit of the arctic fox (*V. lagopus*) is linked to increased food  
466 availability with a resulting retreat of the arctic fox populations from Low Arctic habitat due to  
467 climate warming (Killengreen et al. 2007). Additionally, more humans and infrastructure  
468 correlate with higher density of red fox populations and increased human activity in the Arctic  
469 will likely promote further encroachment of the red fox into arctic fox habitat and further retreat  
470 its population (Elmhagen et al. 2017, Kapsar et al. 2022).

471

### 472 **3.2 Genetic attributes and plastic response of Arctic species**

473 Most Arctic plant and animal species are characterized by low genetic diversity (Mcgraw 1995,  
474 Gilg et al. 2012). Range contractions risk further decreasing genetic diversity, within and across  
475 populations, due to evolutionary processes like genetic drift and reduced gene flow (Ellstrand  
476 and Elam 1993, Rubidge et al. 2012). Reduced genetic diversity of High Arctic plants and  
477 animals has been attributed to genetic bottlenecks and repeated loss of habitat due to range  
478 contractions during warmer interglacial periods (Alsos et al. 2002, Gilg et al. 2012, Birkeland et  
479 al. 2017, Stojak and Jędrzejewska 2022). Global-scale population genetic data shows that while  
480 plants in taiga and tundra ecosystems support relatively low genetic diversity, it tends to be  
481 higher in refugial populations and lower in more recently deglaciated areas, which has been  
482 attributed to persistence in glacial refugia, demographic processes, and glacial and post-glacial  
483 range contractions and expansions (Stenström et al. 2001, Taberlet et al. 2012, Eidesen et al.  
484 2013). Ancient DNA, microfossil, and pollen records from sediment samples have been used to  
485 examine shifts in plant communities over the past 24,000 years in the Polar Ural Mountains  
486 (Clarke et al. 2020) as well as support the persistence of Arctic species in refugia (Alsos et al.  
487 2016, 2020). Refugia from the early (1.0 to 1.6 mya) to late Pleistocene in Alaska, Russia,  
488 Europe, and Greenland were recently identified for a widespread Arctic plant, the Arctic Bell-  
489 Heather (*Cassiope tetragona*; Elphinstone et al. 2024). Areas within the Beringian region, such  
490 as eastern Russia and western Alaska, as well as locations within British Columbia, Europe,  
491 and Greenland, remained mostly unglaciated during Pleistocene glacial cycles and served as  
492 refugia for many Arctic plant and animal species (Abbott and Brochmann 2003, Alsos et al.  
493 2005, Elphinstone et al. 2024).

494

495 Arctic species with small effective population sizes, long life cycles, and narrow distribution  
496 ranges are at a particular disadvantage in the warming Arctic (Hamrick and Godt 1996, Gilg et  
497 al. 2012). Wind-pollinated Arctic plants with wide geographical ranges are generally expected to  
498 retain higher genetic diversity as compared to animal-pollinated species, selfing species, and  
499 species with narrow geographical ranges (Hamrick and Godt 1996, Alsos et al. 2012). The  
500 primary pollination and dispersal mechanism, wind, may not be sufficient to overcome the  
501 negative consequences of range contractions since it does not guarantee that populations will  
502 be able to mate or establish over long distances. Low levels of seed germination and seedling  
503 recruitment are common in Arctic plant species (Bliss 1958), and low seedling survival in

504 addition to dense forest canopy reducing population connectivity can inhibit establishment and  
505 range expansion of plants in the Arctic. For instance, treeline has been shown to be a barrier to  
506 gene flow between ecotypes of the tussock cottongrass (*Eriophorum vaginatum*) found south  
507 and north of the Brooks Range in Alaska (Stunz et al. 2022), and white spruce (*Picea glauca*)  
508 north of the Brooks Range demonstrates treeline advance in the region (Dial et al. 2022).  
509 Advancement of trees, and specifically the advancement of the forest-tundra ecotone, may  
510 fragment open tundra and further exacerbate landscape resistance, limiting gene flow for some  
511 Arctic plant species (Stunz et al. 2022). Long-distance dispersal with higher seed germination  
512 rates and establishment under climate change is expected for some species, especially wind-  
513 pollinated and wind-dispersed trees (Hamrick and Godt 1996). Additionally, more open water in  
514 the Arctic Ocean leads to more southerly winter winds which can facilitate seed and pollen  
515 dispersal further northward, above treeline (Dial et al. 2022). Increased open water can lead to a  
516 deeper snowpack, which can protect juvenile trees from severe winter winds, and snowmelt can  
517 alleviate soil moisture limitations during the growing season (Dial et al. 2022). Multiple animal  
518 species utilize the forest-tundra ecotone for shelter, breeding, and foraging (e.g. passerine and  
519 wader bird species that breed in peatlands; Järvinen et al. 1987), and structural changes in  
520 forest habitat could strongly alter the distribution of animal species as well.

521  
522 Major knowledge gaps remain regarding the potential of certain species and/or genotypes to  
523 expand their ranges and experience an increase or decrease in establishment success in a  
524 warming Arctic (Alsos et al. 2002). Increased encroachment of red fox into Low Arctic areas of  
525 Siberia may lead to further population fragmentation and inbreeding of Scandinavian arctic fox  
526 genotypes by limiting connectivity of Siberian and Scandinavian arctic fox subpopulations  
527 (Elmhagen et al. 2017, Cockerill et al. 2022). The Arctic is rich in polyploid plant species, and  
528 some polyploids may be able to expand their ranges due to their increased genetic variability,  
529 especially if they are long-distance dispersers or derived from multiple colonizations  
530 (Brochmann et al. 2004, Meimberg et al. 2009, Mata et al. 2023). Species like grayleaf willow  
531 (*Salix glauca*) and dwarf birch (*Betula nana*) have been shown to have narrower distributions,  
532 low rates of seed germination, and increased clonal reproduction near northern range limits and  
533 during periods of climate cooling (e.g. 4 kya-2.5 kya; Birks 1991, Alsos et al. 2002). As more  
534 extreme weather events increase and climate fluctuations occur in a warming Arctic, the  
535 frequency of sexual reproduction, especially in thermophilous species like the grayleaf willow  
536 and dwarf birch, is expected to increase as well (Alsos et al. 2002). Increased sexual  
537 reproduction leading to higher adaptation potential of aggressive shrub species could have  
538 major implications for ecosystem functioning and biodiversity in the Arctic as their abundance  
539 increases in Subarctic and Arctic areas (Tape et al. 2006, Berner et al. 2018). If range  
540 expansion of polyploids and higher rates of sexual reproduction under climate change leads to  
541 increased genetic variability of Arctic plant species (Alsos et al. 2002, Brochmann et al. 2004),  
542 the survival, abundance, and adaptation of these plants and the composition of their associated  
543 arthropod communities may also be altered (Colella et al. 2020).

544  
545 Phenotypic plasticity is common in Arctic plants and allows species to produce different  
546 phenotypes under modified environmental conditions, allowing for direct response and potential  
547 resilience to environmental change (Bret-Harte et al. 2001, Deslippe and Simard 2011, Dobbert

548 et al. 2021). While the long lifespan of many Arctic species suggests limited evolutionary  
549 potential to evolve fast enough under rapid environmental change, plasticity, increased mutation  
550 rate, and increased sexual reproduction (with subsequent seed germination and seedling  
551 establishment) may promote adaptation and persistence of some species (Colella et al. 2020).  
552 For instance, the developmentally plastic dwarf birch can shade out plants like the tussock  
553 cottongrass, a species that has a disproportionate impact on ecosystem function due to  
554 enhanced nutrient cycling and deep rooting that increases active soil layer depth (Chapin and  
555 Shaver 1985, Curasi et al. 2022), which could lead to drastic modifications in ecosystem and  
556 biotic community structure, both aboveground and belowground (Keuschnig et al. 2022).

557

### 558 **3.3 Changes in biological community composition and species turnover**

559 Arthropods influence Arctic plant and animal diversity and are a fundamental component of food  
560 webs (Hodkinson and Coulson 2004). Large variability in arthropod abundance and biodiversity  
561 trends are common in the Arctic, such as cyclical abundances of the autumnal moth observed  
562 from 1968-2020 in the Swedish Lapland (Nielsen et al. 2013, Andersson et al. 2022).  
563 Contrasting temporal trends in total abundance of multiple arthropod groups have been  
564 observed in Greenland, where decomposers increased in some habitats and key pollinator fly  
565 species decreased in others (Gillespie et al. 2020). As permafrost thaws and the soil active  
566 layer increases, we can expect impacts on belowground biodiversity, such as increased  
567 abundance of soil microbes and proliferation of extensive mycorrhizal networks (Callaghan et al.  
568 2004). Depending on soil drainage following permafrost thaw, arbuscular mycorrhizal-  
569 dependent plant species such as marsh cinquefoil (*Potentilla palustris*) may promote  
570 mycorrhizal network formation and dominate in upland, drier sites, while the non-mycorrhizal  
571 tussock cottongrass will likely have an advantage in inundated soils (Schütte et al. 2019). Drier  
572 top soils following permafrost loss have led to drastic modifications in methanogenic and  
573 methanotrophic microbial communities (Keuschnig et al. 2022). While soil biodiversity and  
574 vegetation composition changes, the degree of change in soil temperature, soil moisture, and  
575 active layer depth will also be impacted by unknown factors, such as invasive species, extreme  
576 weather-events and the frequency and timing of freeze-thaw events (Nielsen and Wall 2013).  
577 For instance, extreme winter warming has shown to negatively impact microarthropods reliant  
578 on wet soils and water films, and declines are expected to be more severe in warmer and drier  
579 conditions (Bokhorst et al. 2012).

580

581 Large knowledge gaps exist regarding how Arctic biological community composition will evolve  
582 in the face of environmental change (**Table 1**). Arctic bird communities are impacted by a  
583 multitude of factors beyond climate change, such as predation, egg harvesting, industrial  
584 activity, and increased vegetation cover (Doyle et al. 2020). Denser and taller vegetation can  
585 have contrasting effects on Arctic birds, as some species like the ground-nesting passerines,  
586 may benefit, but most species decline when vegetation becomes too tall (Thompson et al.  
587 2016). Many Arctic species are expected to experience changes in population connectivity  
588 (Niskanen et al. 2019), and in recently deglaciated areas (i.e. glacier forelands), glacier melt will  
589 expose more land, serving as potential colonization grounds for many Arctic species. What  
590 colonizers arrive first can also impact soil attributes, especially if vegetation shifts occur faster  
591 than soil development. If cryptogams are the first colonizers to arrive in a newly deglaciated

592 area, for instance, the formation of slightly acidic soil may restrict some plants, soil microbes,  
593 and other species from establishing. These areas can serve as potential refugia for cold-  
594 adapted species as environmental change persists, but may also be difficult for thermophilic  
595 Arctic plants to establish due to colder microhabitats. Some deglaciaded areas may support  
596 'novel ecosystems,' where modified abiotic factors create incipient ecosystems with no previous  
597 or current analog under climate change (Reu et al. 2014, Parmesan et al. 2022).

598  
599 Other unknowns concern how environmental change will influence changes in community  
600 composition, species turnover, and migration, especially across geopolitical boundaries: what  
601 will Arctic biotic community composition look like in the future? For instance, intergovernmental  
602 cooperation influencing species management and conservation policy can influence population  
603 connectivity and abundance, as gray wolf (*Canis lupus*) populations migrate frequently between  
604 Norway and Sweden (Kaspar et al. 2022). Changes in the range and abundance of Arctic  
605 species and ensuing shifts in biotic community composition are expected to impact biodiversity  
606 at both micro- and macro scale, and mismatches between peak prey availability and loss of pre-  
607 migration and staging (areas to refuel and rest) habitat (Smith et al. 2020) can influence Arctic  
608 bird abundance. Based on a dataset from 1980-2017 across circumarctic sites, waders, which  
609 comprise nearly half of Arctic terrestrial bird species, had the largest proportion of species  
610 abundance declines, while almost half of waterfowl species increased in abundance (Smith et  
611 al. 2020). Increases and decreases in soil microbial diversity and arthropod diversity may impact  
612 plant and animal diversity and abundance (Hodkinson and Coulson 2004, Schmidt et al. 2017b).  
613 Lack of arthropod monitoring and limited inclusion in long-term research suggests that arthropod  
614 species may disappear before being identified as environmental change progresses in the Arctic  
615 (Gillespie et al. 2020), highlighting the necessity to better catalog biodiversity of these  
616 communities and improve understanding of their influence on ecosystem functioning and trophic  
617 structure (Taylor et al. 2020). In many Arctic ecosystems, muscid flies have been identified as  
618 the most common (and efficient) pollinators, and declining muscid fly diversity and abundance,  
619 along with increasing flowering phenological mismatch, could threaten reproductive success of  
620 many Arctic plant species (Tiusanen et al. 2016). The frequency of Arctic plant sexual  
621 reproduction via animal pollination is predicted to decrease in some species and increase in  
622 others due to a shorter flowering window under climate change (Schmidt et al. 2016). Yet,  
623 earlier flowering can also lead to increased reproductive fitness as more flowers and fruits are  
624 produced due to longer development time (Collins et al. 2024). Extrapolating trends in animal  
625 pollination in the Arctic is uncertain because of high turnover of pollinator species (Cirtwill et al.  
626 2018), as well as conflicting projections regarding range expansions and contractions of  
627 arthropods (Elberling and Olsen 1999, Gillespie et al. 2020).

628  
629 How vegetation composition will change in the future, especially regarding species turnover and  
630 invasive species, is uncertain and can impact ecosystem function. While dynamic global  
631 vegetation models consistently project woody shrub and boreal forest expansion into tundra  
632 regions, disturbance factors, such as insect damage, wildfires, extreme weather events,  
633 dynamics of permafrost thaw/active layer depth, and nutrient availability are poorly simulated or  
634 unrepresented (Peng et al. 2021, Parmesan et al. 2022, Heffernan et al. 2024). Soil moisture  
635 dynamics can inform on plant community composition as thirty years of vegetation survey data

636 showed that the strong spatial relationships between plant traits (such as height, specific leaf  
637 area, leaf nitrogen content) along temperature and soil moisture gradients were mostly  
638 explained by species turnover (Bjorkman et al. 2018). Plant community height increased across  
639 all sites as taller, thermophilous plant species increased, particularly in areas where soil  
640 moisture was high (Bjorkman et al. 2018). Disturbance events and soil nutrient content can also  
641 inform on Arctic vegetation shifts, as abrupt permafrost thaw (leading to thaw ponds) and low  
642 nutrient availability can promote graminoid dominance, and ultimately how carbon cycling may  
643 shift under vegetation change (Wein and Bliss 1974, van der Kolk et al. 2016, Parmesan et al.  
644 2022). Other underrepresented plant and soil communities, such as cryptogams, rhizosphere  
645 microbes, and biocrusts (moss, lichen and cyanobacteria communities on the soil surface), will  
646 likely be impacted as shrubs and other vegetation types increase in abundance and expand  
647 their ranges northwards (Wullschlegel et al. 2015, Bokhorst et al. 2023, Gu et al. 2023).  
648 Biocrust soils in the Arctic can fix and store significant amounts of carbon in the surface soil  
649 layer (Juottonen et al. 2020, Jung et al. 2018), and changes in biocrust extent and thickness  
650 need to be investigated to better understand impacts on carbon cycling (Williams et al. 2017,  
651 Colella et al. 2020). High grazing intensities have also been shown to drive plant communities  
652 towards graminoid and forb-dominated communities (Olofsson et al. 2001, van der Wal 2006,  
653 Kitti et al. 2009, Olofsson and Post 2018). Furthermore, herbivores can mitigate the climate-  
654 driven expansion of deciduous shrubs (Post and Pedersen 2008, Olofsson et al. 2009), whereas  
655 less palatable evergreen shrubs are not influenced by large herbivores to the same extent  
656 (Vowles et al. 2017a, b). Impacts of reindeer grazing on vegetation have also been shown to be  
657 habitat-specific (influenced by historic and current land use) and indirect, as reindeer primarily  
658 modify soil nutrient cycling (Stark et al. 2023). While biomass increases have been predicted for  
659 the Arctic tundra, there is growing evidence that tree expansion into tundra may not increase  
660 carbon storage as models predict (Canadell et al. 2021). In Abisko, for instance, total ecosystem  
661 carbon storage is greater in tundra heath (owing to greater soil carbon stocks) than in the  
662 mountain-birch forest, and high plant activity in the forest during the growing season stimulates  
663 the decomposition of older soil organic matter (Hartley et al. 2012). Mountain birches have been  
664 shown to facilitate birch seedling performance and survival at high-stress sites (Eränen and  
665 Kozlov 2008) and can also facilitate establishment of other species by improving soil fertility  
666 (Mikola et al. 2018). It is uncertain whether tundra plant community composition in the direct  
667 vicinity of birch individuals will become more similar to the composition found in the nearby birch  
668 forest (Sundqvist et al. 2008). Thus, with multiple factors driving vegetation change in the Arctic,  
669 it is uncertain if adding single factors will create additive effects on biotic community structure or  
670 generate new communities (Wang et al. 2021, Parmesan et al. 2022, Scharn et al. 2022).

671

## 672 **Future outlook of the ecosystem-climate-human nexus**

673

674 The extent that ocean acidification, surface water and surface air temperature rise will influence  
675 the northward range expansions of Arctic species, the reorganization of polar systems, and  
676 deterioration of the cold barrier between the Subarctic forest and Arctic tundra is uncertain  
677 (**Table 1**; Constable et al. 2022). Reductions of seasonal sea ice, loss of multi-year ice, and  
678 altered wind patterns are expected to facilitate interactions between coastal communities,  
679 tourism, shipping, and commercial fishing industries. While Coupled Model Intercomparison



680 Project Phase 6 (CMIP6) climate projections perform better than earlier models at predicting  
681 sea ice loss for a given amount of carbon dioxide emissions, they are unable to project the  
682 future extent of Arctic sea-ice cover with confidence (Notz and SIMIP Community 2020).  
683 Therefore, the degree to which the health, security, and subsistence resource availability of  
684 Indigenous Peoples could be impacted by dwindling sea-ice area is largely unknown (Constable  
685 et al. 2022). While decreased abundance of caribou and reindeer are affected by multiple  
686 climate factors, caribou population declines have been partly attributed to reduced lichen  
687 consumption (observations by Inuit hunters; Knotsch and Lamouche 2010). The degree to  
688 which lichen cover and diversity will be lost as tundra plant communities shift under climate  
689 change (Elmendorf et al. 2012), and how this will further impact reindeer populations are  
690 unknown. Several components of terrestrial biodiversity are severely understudied in the Arctic  
691 (e.g. Taylor et al. 2020). We currently lack a sufficient overview of the Arctic diversity for several  
692 organism groups such as arthropods, bryophytes, and lichens, and the absence of these groups  
693 in climate change models hampers the ability to assess how biodiversity and ecosystem  
694 functioning will be impacted. The extent to which economic development and natural resource  
695 extraction will increase in the Arctic is also uncertain and has the potential to dramatically  
696 impact Arctic local human communities and species due to increased ship traffic, noise (and  
697 other) pollution, increased carbon emissions, damage to Arctic summering and breeding  
698 grounds, and interruption of subsistence practices (Kaspar et al. 2022). Additionally, as  
699 permafrost thaw and glaciers melt, sequestered pollutants, microparasites, and hazardous  
700 waste in Arctic environments will likely remobilize and increase risk exposure to ecosystems  
701 and humans (Wang et al. 2019, Colella et al. 2020), potentially leading to wide-scale negative  
702 impacts on biodiversity, ecosystem function, and health. Integrative frameworks to examine how  
703 human and natural systems are coupled (socio-environmental interactions) and how drivers in  
704 other regions affect the Arctic (and vice-versa) are needed to increase resilience of Arctic  
705 species and systems under climate change (Kapsar et al. 2022).

706  
707 Warming in the Arctic has occurred at nearly four times the global average since 1979  
708 (Rantanen et al. 2022), prompting environmental and ecosystem changes such as shifts in  
709 phenology, subsistence resources, biodiversity, species distributions, snow regimes, sea ice,  
710 permafrost extent, and carbon and nutrient cycling. Arctic tipping elements, such as loss of  
711 Arctic sea ice, AMOC collapse, and boreal forest dieback may be triggered by anthropogenic-  
712 caused climate change by 2100, potentially leading to a cascade of impacts that could  
713 drastically alter Arctic ecosystems and the global climate system (Armstrong McKay et al.  
714 2022). Wide-ranging impacts on humans in the Arctic have only been superficially explored at  
715 present, especially in relation to changes in climate, ecosystem function, and biodiversity  
716 response.

717  
718 Pairing robust scientific models that utilize long-term monitoring, spatial genomics, and baseline  
719 and resampled datasets with integrative assessments are needed to better predict ecological  
720 response as well as capture complex causal chains in coupled human and natural systems. To  
721 capture these complex dynamics, human responses to environmental change must be included  
722 in models, such as integrated climate assessments (Maxwell et al. 2015, Beckage et al. 2022).  
723 Strategies to increase the resilience of species and humans to environmental change are

724 imperative, and urgently needed to protect Arctic ecological and social systems. These  
725 strategies should include pathways to facilitate swift response to unprecedented change, such  
726 as decentralized governance and management policy to promote inclusivity, (especially in  
727 regards to gender, social justice, and equity), incorporation of Indigenous Peoples and  
728 knowledge in co-management of Arctic research and resources, creating local, publicly-  
729 available tools for regional weather, climate, and ecosystem response, and international  
730 collaboration and governance for adaptation and carbon mitigation policy (Andersson 2021,  
731 Constable et al. 2022).

732  
733 As it pertains to science, few current approaches to inter- and trans-disciplinary science policy  
734 interfaces integrate Indigenous knowledge and local knowledge, which can be used to better  
735 understand how environmental change in the Arctic impacts species distributions at the  
736 ecosystem level (Ksenofontov et al. 2019, Knopp et al. 2022). The inclusion of Indigenous- and  
737 local knowledge leverages the use of diverse knowledge types to inform biodiversity  
738 assessments (Pascual et al. 2021). Thus, peoples' perspectives and participation matter when it  
739 comes to designing conservation policies, because any effective policy needs to be  
740 implemented and respected by people in the end (Barry et al. 2015, Neuteleers et al. 2021). In  
741 consideration of the continued influence of colonialism and rare inclusion of Indigenous  
742 knowledge in science policy and governance, achieving climate justice for Indigenous  
743 populations is still a distant reality (Whyte et al. 2019, Chakraborty and Sherpa 2021, IPBES  
744 2022).

745  
746 Due to colonial legacies, Indigenous Arctic populations have not been included in the decision-  
747 making processes pertaining to their land. This is one of several factors contributing to the  
748 heightened impacts of climate change they are now experiencing (Whyte et al. 2019). While not  
749 being historically responsible for anthropogenic climate change, they are the ones bearing the  
750 consequences. Indigenous knowledge, grounded on an ontology that understands people and  
751 the Earth as one, has been consistently overlooked and neglected by settler states, increasing  
752 the vulnerability of these populations and perpetrating historical injustices and inequities (Whyte  
753 et al. 2019, Snook et al. 2020, Ford et al. 2021). Settler states, such as Canada, are moving to  
754 rectify this situation by encouraging inclusive governance and Indigenous leadership, (e.g. the  
755 development of local community-based monitoring systems for hunters and fishers) to  
756 document changes in the environment (Danielsen et al. 2014, AMAP 2017). Yet, Indigenous  
757 knowledge remains largely tokenized (at best) in governmental frameworks (Whyte et al. 2019).  
758 For instance, the incorporation of biodiversity and nature value into national-level policy is very  
759 limited, and  $\leq 5\%$  of valuation studies have reported integration into policy (IPBES 2022).

760  
761 To date, research has focused on the influence of environmental change in the North American  
762 Arctic, rather than climate change impacts in the Arctic at a global level, leading to a biased  
763 view of the impacts on humans and natural systems (Metcalf et al. 2018, Kaspar et al. 2022).  
764 Additionally, when considering climate change impacts on the environment, Arctic residents  
765 often cite factors such as tourism, international trade, and natural resource development, in  
766 contrast to researchers that typically consider climate change to be the sole external influence  
767 (Moerlein and Carothers 2012). These gaps in the academic approach highlight the necessity of

768 incorporating both social and environmental change into Arctic climate change research and  
 769 assessments. The migration of species across geopolitical boundaries as environmental change  
 770 progresses will have a dramatic impact on local Arctic communities and future management of  
 771 fauna and flora of the region. The identification and protection of refugial populations and habitat  
 772 corridors, to uphold and potentially increase population connectivity and migration, can promote  
 773 species and overall ecosystem resilience in the Arctic. Indigenous Peoples and culture should  
 774 be explicitly integrated in research and resource management (Johnson et al. 2015, Brattland  
 775 and Mustonen 2018). Increased intergovernmental collaboration at a circumarctic and global  
 776 scale needs to be prioritized to create comprehensive and effective policies and a nexus  
 777 approach will foster deeper knowledge of the region and inclusive conservation and  
 778 management. By reviewing current understanding and outlining knowledge gaps, we call for  
 779 mobilization towards more adaptable governance systems to promote resilience of humans and  
 780 natural systems in a changing Arctic.

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783 **Tables**

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785 **Table 1.** Some outstanding questions regarding future conditions within the ecosystem-climate-  
 786 human nexus across a range of future greenhouse gas emission pathways.

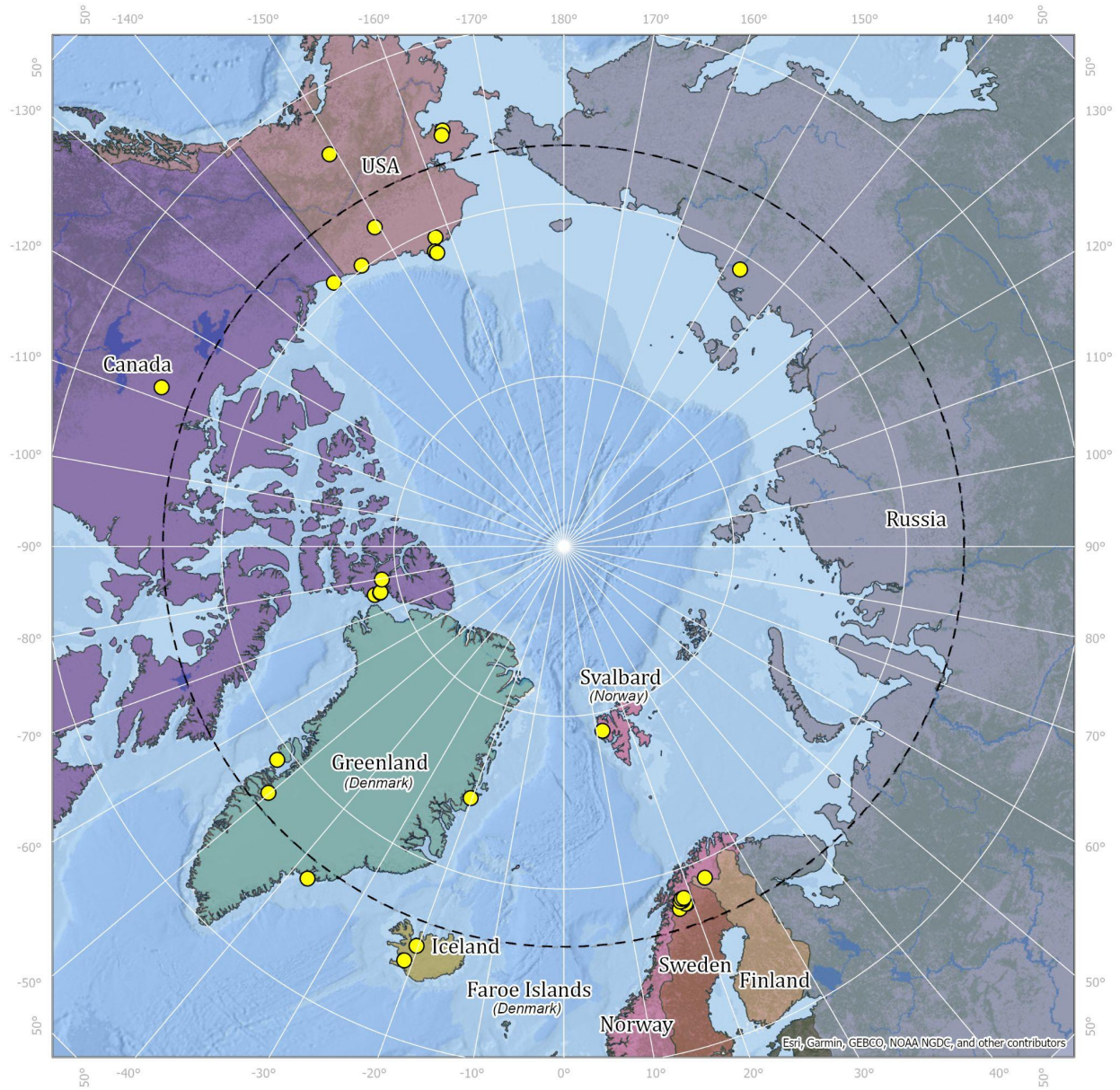
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Nexus Component	Future Outlook Questions
<b>Ecosystems</b>	<ul style="list-style-type: none"> <li>● How will “overshooting” (temporarily or not) 1.5°C warming affect Arctic species and ecosystems?</li> <li>● What ecosystem changes are possible due to (individually and in combination): i. loss of surface soil because of permafrost thaw and collapse, ii. shrubification of the landscape; iii. movement and biotic community change of the forest-tundra ecotone, iv. winter warming events that could potentially cause massive vegetation/foraging source loss, v. change in growing season length, vi. large-scale losses through wildfires and other disturbances?</li> <li>● To what extent could the rate of genetic mutation change (mutational meltdown/extinction vortex) in Arctic species under climate change?</li> </ul>
<b>Humans</b>	<ul style="list-style-type: none"> <li>● How much of the Arctic will be converted to cropland, pasture and/or otherwise modified due to natural resource extraction, tourism, and other development projects as climate change progresses?</li> <li>● Will silviculture expand into the Arctic, and if so, to what extent and how will permafrost thaw, ecosystem processes, and biodiversity be impacted?</li> </ul>

	<ul style="list-style-type: none"> <li>● How will cultural heritage, subsistence harvesting, society, and language of Sámi, Inuit, and other Indigenous Peoples be impacted?</li> <li>● How will forest-ecotone advancement facilitate disease transmission to humans by increasing habitat for animal and insect vectors?</li> <li>● How will the release of sequestered pollutants and disease from permafrost thaw impact biological and human communities?</li> </ul>
<b>Climate</b>	<ul style="list-style-type: none"> <li>● What greenhouse gas emission scenario will most likely follow at the end of the 21st century?</li> <li>● What happens when climate change in the Arctic triggers global tipping points?</li> <li>● How will terrestrial ecosystem and vegetation changes influence the surface energy balance, aerosols and clouds?</li> <li>● How often will extreme events occur and what kinds of regional and global impacts will they have?</li> <li>● How strong will the climate feedback from thawing permafrost and wetland methane emissions be?</li> </ul>

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813 **Figures**  
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816 **Figure 1.** Map showing the territory of the eight Arctic States (Canada, The Kingdom of  
817 Denmark, Finland, Iceland, Norway, The Russian Federation, Sweden, and The United States  
818 of America (USA)) and plot-based vegetation research studies examining abundance and  
819 phenological change at 32 ITEX sites in the Arctic (above 63°N) from Bjorkman et al., 2020.  
820 Yellow circles indicate site location. The Arctic States and their associated territories are as  
821 follows: Canada, The Kingdom of Denmark, Finland, Iceland, Norway, The Russian Federation,  
822 Sweden, and The United States of America. The Arctic Circle is represented with a black  
823 dashed line.

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