

Global exposure risk of frogs to increasing environmental dryness

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39 **Data accessibility:** Data and code to reproduce the study are available on the GitHub repository:
40 <https://github.com/nicholaswunz/global-frog-drought> and will be available on Zenodo upon
41 acceptance.

42 **ABSTRACT**

43 Species exposed to prolonged drying are at risk of population declines or extinctions. A key missing
44 element for assessments of climate change risk is the sensitivity of species to water loss and their
45 microhabitat preference, or ecotype, as both dictate the risk of environmental drying. Here, we
46 identified globally where water-sensitive ectotherms, i.e. anurans, are at risk to increasing aridity
47 and drought, examined which ecotypes are more sensitive to water loss from 238 species, and
48 estimated how behavioural activity is impacted by future drought and warming scenarios through
49 biophysical models. Under an intermediate and high emission scenario, 6.6 and 33.5% of areas
50 occupied by anurans will increase to arid-like conditions, and 15.4 and 36.1% are at risk of
51 exposure to a combination of increasing drought intensity, frequency, and duration by 2080–2100,
52 respectively. Critically, increasing arid-like conditions will increase water loss rates and anurans in
53 dry regions will almost double the water loss rates under a high emission scenario. Biophysical
54 models showed that during the warmest quarter of the year, the combination of drought and
55 warming reduced an anuran’s potential activity by 17.9% relative to the current conditions
56 compared to warming alone which reduced potential activity by 8%. Our results exemplify the
57 widespread exposure risk of environmental drying for anurans, posing a serious challenge for the
58 lives of water-sensitive species beyond the effects of temperature alone.

59 **Keywords:** amphibian decline, climate change, dehydration, desiccation, hydroregulation,
60 macrophysiology, thermoregulation

61 **INTRODUCTION**

62 Global warming and land modification are expected to increase the frequency and intensity of
63 droughts, which are anomalous periods of low precipitation (Slette et al., 2019; Pokhrel et al., 2021;
64 Zhao and Dai, 2022). This synergistic interaction is accelerating extinction risk for species sensitive
65 to rapid water loss (Zylstra et al., 2019; Lowe et al., 2021). Amphibians are prime examples of
66 terrestrial species that are heavily dependent on water, and often have higher rates of water turnover
67 than other terrestrial animals (Lillywhite, 2006; Hillman et al., 2009). The group is characterised by
68 their sensitivity to both environmental temperature and moisture. Thus, negative or challenging
69 water balances can reduce opportunities for foraging, finding mates, and dispersal (Sodhi et al.,
70 2008; Ficetola and Maiorano, 2016; Lertzman-Lepofsky et al., 2020). Despite synthesis studies on
71 the effects of climate change on amphibians (Li et al., 2013; Campbell Grant et al., 2020), global
72 assessments of climate vulnerability have mainly been through the lens of temperature (Snyder and
73 Weathers, 1975; Gunderson and Stillman, 2015; Murali et al., 2023; Pottier et al., 2024b). Yet, the
74 impact of extreme moisture deficit events such as drought, needs to be assessed, given the

75 sensitivity of amphibians to water loss and their threatened status relative to other taxonomic groups
76 (Luedtke et al., 2023).

77 Assessing species vulnerability to climate change relies on the degree of environmental
78 exposure and the species' sensitivity to the exposure (Williams et al., 2008). The extent of **exposure**
79 to drying is mainly influenced by the frequency, intensity, and duration of the stressor. However,
80 the impact of climate change is not singular (Rozen-Rechels et al., 2019). While environmental
81 water stress primarily depends on surrounding water availability (in the air, substrate, and water
82 bodies) and how much is evaporated or transpired by plants (Trenberth et al., 2014), warming alone
83 will increase evaporation rates by increasing the vapor pressure deficit (VPD). VPD is one of the
84 primary forces driving evaporation rates and atmospheric drought (Park Williams et al., 2013). High
85 VPDs are known to be major drivers of mortality events in plants (Eamus et al., 2013; Grossiord et
86 al., 2020), and the rise in evaporation rates reduces activity in many ectotherms (Kearney et al.,
87 2018; Lertzman-Lepofsky et al., 2020) and increases exposure to lethal dehydration events in birds
88 and mammals (Kearney et al., 2016; Albright et al., 2017; Iknayan and Beissinger, 2018).
89 Experimental and field studies have suggested amphibians behaviourally prefer environments that
90 maintain hydration over thermoregulation, emphasising the importance of water availability in
91 informing climate vulnerability (Navas et al., 2007; Anderson and Andrade, 2017; Galindo et al.,
92 2018). Understanding the water loss vulnerability is therefore a necessary component for
93 determining species vulnerability to global environmental change.

94 Behavioural and physiological **sensitivity** to water loss can vary by species. Some
95 ecological types ('ecotypes') have unique adaptations to drying conditions (Toledo and Jared,
96 1993). For example, fossorial species living in arid regions burrow underground during the non-
97 breeding period to aestivate (Carvalho et al., 2010), and some are able to create water-proof
98 cocoons to conserve water (Withers, 1995; Tracy et al., 2007). In contrast, some arboreal anurans
99 produce skin waxy secretions that reduce evaporative water loss (Shoemaker and McClanahan,
100 1980; Stinner and Shoemaker, 1987; Amey and Grigg, 1995). Most amphibians, however, are
101 vulnerable to dehydrating conditions compared to other terrestrial vertebrates and must therefore
102 behaviourally maintain water balance by using shade, burrows, or by adopting a water conserving
103 postures when inactive (Pough et al., 1983). Because of this diversity in natural history, some
104 ecotypes are likely to be more sensitive to environmental drying than others due to the microclimate
105 preferences and limited physiological adaptations to conserve water. The increased frequency and
106 intensity of drought conditions—even for arid specialists—can impact anurans by restricting the
107 activity periods for foraging, seeking mates, dispersing, and delaying breeding cycles (Wassens et

108 al., 2013), all of which ultimately affects population dynamics, disease susceptibility, and shifts in
109 geographic distribution (Kohli et al., 2019; Zylstra et al., 2019; Kupferberg et al., 2022).

110 In this study, we aimed to (i) identify regions where anurans are at risk to increased aridity
111 and drought by quantifying the spatial overlap between anuran species richness and areas of
112 increasing aridification and drought (including intensity, frequency and duration) under future
113 warming scenarios, (ii) examine ecotype sensitivity of adult anurans to evaporative water loss and
114 water uptake through comparative meta-analyses, and (iii) simulate physiological limits for activity
115 of a hypothetical adult frog in relation to future drought and warming scenarios in order to assess
116 the relative effects of warming and/or drought on changes in potential activity hours for each water-
117 conserving strategy under different climate scenarios. We focused on water loss risk for adult
118 terrestrial anurans because most larval stages are aquatic and rely on water bodies.

119 **METHODS**

120 **Aridity and drought risk**

121 *Global climatic water balance data*

122 The global water cycle is complex, and many indices have been used to understand and predict
123 environmental water stress across space and time (Spinoni et al., 2020; Ukkola et al., 2020; Li et al.,
124 2021). Here, we used two common metrics of environmental dryness, the aridity index (AI), and
125 the Palmer Drought Severity Index (PDSI). The AI represented the degree of dryness of the climate
126 and is calculated as the ratio of the total amount of water supply (precipitation) relative to the
127 amount of water loss (potential evapotranspiration). More arid regions have a smaller index value
128 (Cherlet et al., 2018). In this study, the AI represented the broadscale dryness experienced by
129 amphibians in that area. The PDSI is a widely used meteorological drought index calculated using
130 soil moisture and precipitation data of previous months. By using surface air temperature and a
131 physical water balance model, the PDSI takes into account the basic effect of global warming
132 through potential evapotranspiration and is effective in determining long-term drought (Palmer,
133 1965; Dai et al., 2004). In this study, the PDSI represented extreme changes in meteorological water
134 deficit or surplus experienced. The AI was categorised to five categories (Humid, Dry sub-humid,
135 Semi-arid, Arid, and Hyper-arid) and the PDSI was categorised to seven categories (Extremely
136 moist, Very moist, Moderate moist, Normal, Moderate drought, Severe drought, and Extreme
137 drought) based on descriptions from Budyko (1961) and Palmer (1965) in **Table S1**.

138 High resolution (~4km²) global data on precipitation (mm/month) and potential
139 evapotranspiration (mm/month) were obtained from Abatzoglou et al. (2018) under 1) the current
140 climate (1970–2000), 2) an intermediate greenhouse gas emission scenario of +2°C (Shared

141 Socioeconomic Pathway 2–4.5; SPP2–4.5), and 3) a high greenhouse gas emission or “business-as-
142 usual” scenario of +4°C (SSP5–8.5) by 2080–2099 (**Fig. S1a–i**). We used a self-calibrated PDSI
143 with Penman–Monteith potential evapotranspiration representing the current climate (1970–2000)
144 and an intermediate and high emission scenario by 2080–2099 (SSP2–4.5 and SSP5–8.5) from Zhao
145 and Dai (2022). The SSP2–4.5 and SSP5–8.5 scenarios were based on the average of 25 Coupled
146 Model Intercomparison Project Phase 6 (CMIP6) models of precipitation, evapotranspiration, soil
147 moisture, and runoff (Eyring et al., 2016), where the mean annual surface temperature is expected
148 to increase by 2.7°C (2.1–3.5°C range) and 4.4°C (3.3–5.7°C range), respectively by 2080–2100
149 (IPCC, 2021). Zhao and Dai (2022) noted that there are large uncertainties in the drought
150 projections. Therefore, our interpretation is based on the average projected drought which is
151 typically in line with other indices of drought (Ukkola et al., 2020; Li et al., 2021; Cook et al., 2022;
152 Zhao and Dai, 2022; Qing et al., 2023).

153 *Anuran species richness, ecotype, and IUCN status data*

154 We extracted the geographic range (excluding introduced range), threatened status, and ecological
155 types of all anuran species (6,416 species) listed in the International Union for Conservation
156 (IUCN) of Nature's Red List of Threatened Species platform (extracted on 11/01/2021; IUCN,
157 2022). Data-deficient Red Listed species were assigned IUCN classifications based on predictions
158 from González-del-Pliego et al. (2019). Ecotypes were classified based on descriptions from Moen
159 and Wiens (2017) focusing on adult behaviour and microhabitat preferences outside the breeding
160 season given that many anurans breed in water but are not adapted to live in water all year (**Table**
161 **S2**). Ecotypes were defined to provide generalised microhabitat use and for modelling the activity
162 in different microhabitats. For example, arboreal species regularly access vertical environments like
163 trees, rock crevices, while stream-dwelling species are restricted to permanent streams or flowing
164 water bodies. We acknowledge that ecotype grouping does not capture the diversity of microclimate
165 preferences. Where possible, we included secondary ecotypes in the raw data, but we analysed
166 species based on their primary ecotype. For example, some species frequent around permanent
167 water bodies, but occasionally have been observed climbing on trees. We would classify them
168 primarily as ‘semi-aquatic’ and secondarily as ‘arboreal’ as they do not spend a substantial amount
169 of time climbing. Species richness was defined as the sum of species in each grid cell (0.5°), based
170 on the geographic range, and was calculated using the ‘calcSR’ function from the *rasterSp* package
171 (<https://github.com/RS-eco/rasterSp>). All raster data were examined and visualised at a spatial
172 resolution of 0.5 arc minute.

173 *Species richness with aridity and PDSI*

174 To examine the relationship between species richness and aridity, the number of species per grid
175 cell was overlapped with the aridity raster layer, where each grid cell was assigned an AI category.
176 The change in species richness between the current and projected (either +2 or +4 °C warming) AI
177 category was calculated as the change in AI category grids occupied by anurans relative to the
178 future projection; a lower number indicates a reduced number of species with the assigned AI
179 category and vice versa.

180 With a monthly prediction of PDSI from 1950 to 2100 globally available from Zhao and Dai
181 (2022), we classified future drought risk in three ways: 1) increase in mean drought intensity
182 (change in PDSI or $\Delta\text{PDSI}_{[\text{intensity}]}$), increase in mean drought frequency (monthly PDSI count below
183 -2 per year or $\Delta\text{PDSI}_{[\text{frequency}]}$), increase in mean drought duration (number of consecutive months
184 with PDSI values below -2 or $\Delta\text{PDSI}_{[\text{duration}]}$). Change in mean $\Delta\text{PDSI}_{[\text{intensity}]}$, $\Delta\text{PDSI}_{[\text{frequency}]}$, and
185 $\Delta\text{PDSI}_{[\text{duration}]}$ under a +2 or +4 °C warming scenario (2080–2100) was calculated relative to the
186 1970–2000 monthly climatology per grid cell ($\Delta\text{PDSI} = \text{PDSI}_{[\text{future}]} - \text{PDSI}_{[\text{current}]}$). Results were
187 presented in the main text as $\Delta\text{PDSI}_{[\text{intensity}]}$, $\Delta\text{PDSI}_{[\text{frequency}]}$ and $\Delta\text{PDSI}_{[\text{duration}]}$ relative to current
188 scenario (1970–2000). The absolute $\text{PDSI}_{[\text{intensity}]}$, $\text{PDSI}_{[\text{frequency}]}$ and $\text{PDSI}_{[\text{duration}]}$ under a +2 or +4
189 °C warming scenario (2080–2100) were also presented in **Fig. S2a–i**.

190 The simultaneous risk of increasing drought intensity, frequency, and duration within a grid
191 cell that are occupied by anurans (species assemblages) was calculated by converting each risk
192 category as binary. Grid cells with a $\Delta\text{PDSI}_{[\text{intensity}]}$ below -2 (indicating increased drought intensity
193 relative to current scenario) were assigned a ‘1’ binary. Both $\Delta\text{PDSI}_{[\text{frequency}]}$ and $\Delta\text{PDSI}_{[\text{duration}]}$ were
194 assigned a binary of ‘1’ if the grid cell has a value of 1 month or higher (indicating increase in
195 frequency or duration relative to current scenario). The number of overlapping binaries were
196 summed up per grid cell. Therefore, a risk factor of 2 indicate species assemblages in the grid cell
197 are at increasing risk of two drought events. We estimated which species assemblages were at risk
198 of experiencing drought events using an arbitrary risk factor scale (species richness \times drought risk),
199 where grid cells with high drought risk and high species richness have higher “assemblage-level
200 risk” than grid cells with high drought risk and low species richness (low assemblage-level risk).

201 **Ecotype sensitivity**

202 *Water loss and uptake data*

203 We systematically reviewed the literature on experimental water loss and uptake studies for adult
204 anurans following the PRISMA-EcoEvo protocol (O'Dea et al., 2021). Full detail of the systematic
205 search, exclusion criteria, and data extraction are presented in the Supplementary Information and
206 visualised in **Fig. S3**. Studies were included if the following hydrological metrics were reported:

207 water loss as the rate of evaporative water loss (EWL), resistance to water loss (total r_t or relative
208 r_i), and change in body mass (absolute or relative), and water uptake as the rate of cutaneous water
209 uptake (WU). Water loss measurements either by forced convection (via wind speed manipulation)
210 or free convection (desiccant) were included in the dataset. The rate of EWL typically incorporates
211 cutaneous and respiratory water loss as most studies do not include methods to distinguish the
212 difference between the two routes of water loss. However, respiratory EWL is assumed to
213 contribute little (5.5% at 25 °C) to the total EWL for amphibians (Senzano and Andrade, 2018).
214 Additionally, the daily turnover of water for a typical amphibian from other processes such as food
215 and water gain, metabolic water production, and urine and faecal loss is considered negligible and
216 equivalent between species for comparative purposes (Hillman et al., 2009), thus we did not
217 consider these aspects when comparing ecotype risks. The rate of EWL, skin resistance (r_i), and
218 cutaneous WU were standardised to a common unit of mg h^{-1} , and s cm^{-1} for relative skin resistance
219 using standard conversions and approximations via biophysical principles (see Supplementary
220 Information section “Calculations and conversions”).

221 We extracted the mean values, variance (either standard deviation or standard error), and
222 sample size from either the study’s main text, supplementary text, table, figures via metaDigitize
223 (Pick et al., 2019), or calculated from the raw data. We also extracted the mean body mass (g), the
224 exposed ambient temperature (°C) during the experiment, skin temperature measured (°C),
225 coordinates of collection site (latitude and longitude in decimal degrees), flow rate (as cm s^{-1} or ml
226 s^{-1}), duration of the experiment (min), the relative humidity (RH; %) and noted whether the animal
227 was in a physiological water-conserving state (e.g., aestivation, cocoon-forming, or water-proof via
228 cutaneous surface fluid or lipids). Dry air from desiccants without RH presented were assumed to
229 be 1% RH. Missing information were obtained by contacting the study authors where possible. We
230 noted if urine was voided prior to the EWL experiments, body mass was defined as “standard body
231 mass” as the presence of urine will affect water loss. Additional unpublished data from 39 species
232 of anurans were included and detailed in the ‘Unpublished data’ section of the Supplementary
233 Information. Note, while we aimed for a comprehensive and representative dataset, there are
234 geographical biases in anuran water loss and uptake studies, especially around the Amazon, central
235 Africa, and Eurasia with high amphibian richness, but little to no water loss and uptake studies (**Fig.**
236 **S4**).

237 Unpublished data from Brazil were collected according to the experimental procedures
238 approved by the Ethical Committee in Animal Use (CEUA) of the Biosciences Institute (#0820),
239 affiliated with UNESP, Rio Claro, São Paulo, Brazil. License for animal capture and transport was
240 provided by Instituto Brasileiro do Meio Ambiente e dos Recursos Naturais Renováveis (IBAMA,

241 #29703-1 and 22028-1), and Secretaria do Meio Ambiente – Instituto Florestal (#26018 -
242 013.054/2011). Unpublished data from South Africa were collected in accordance to Western Cape
243 Nature Permit AAA007-00073-0056, and all experimental procedures were approved by the
244 University of Stellenbosch REC:ACU, Research Ethics Committee: Animal Care and Use (SU-
245 ACUM13-00005).

246 *Phylogenetic data*

247 To account for phylogenetic relatedness in the model, we used the amphibian phylogeny of Jetz and
248 Pyron (2018) comprising of 7,238 amphibian species. There were 5,636 species from our dataset
249 matched the tree, and the remaining species were not available as they were described after 2016.
250 Names were matched with the IUCN. Polytoamy was removed using the function ‘multi2di’, and
251 branch lengths were estimated using the ‘compute.brlen’ function from the *ape* package (Paradis
252 and Schliep, 2018). The generated tree was converted to a phylogenetic relatedness correlation
253 matrix for subsequent analysis.

254 *Water loss and uptake analyses*

255 All models used a No-U-Turn Sampler extension of Hamiltonian Monte Carlo (Hoffman and
256 Gelman, 2014) Bayesian procedure implemented in the *R* package *brms* (Bürkner, 2017) to derive
257 posterior distributions and associated credible intervals (CIs) for the fitted parameters. For all
258 models, we constructed four chains with 5,000 steps per chain, including 2,500-step warm-up
259 periods, hence a total of 10,000 steps were retained to estimate posterior distributions [i.e., (5,000 –
260 2,500) × 4 = 10,000]. Adapt delta was set at 0.99 to decrease the number of divergent transitions,
261 and the maximum tree depth was set to 15 when the depth of tree evaluated in each iteration was
262 exceeded. Fixed effects were assigned weakly informative priors following a Gaussian distribution
263 ($\ln\beta_0$ mean = 0, SD = 3) to speed up model convergence, and Student’s *t* prior with three degrees of
264 freedom was used for group-level, hierarchical effects. The degree of convergence of the model was
265 deemed as achieved when the Gelman–Rubin statistics, \hat{R} (Gelman and Rubin, 1992), was 1.

266 Differences in resistance to water loss (r_i) and WU between ecotypes were examined with
267 the ‘brms’ function from the *brms* package. We analysed r_i because it represents the most
268 physiologically relevant metric of water loss (Feder and Burggren, 1992; Riddell et al., 2017). For
269 the r_i water loss model, we included the natural logarithm transformed body mass (lnMass), VPD
270 (lnVPD), and air flow rate (lnFlow) as additive fixed effects. Study ID, lnMass nested in species,
271 and the phylogenetic correlation matrix were included as group-level effects to account for variance
272 between studies, repeated measures within species, and evolutionary history (Pottier et al., 2024a).
273 For the WU model, the fixed effects were lnMass, treatment temperature, and initial hydration level.
274 Relative humidity was never reported during the WU experiments, thus VPD was replaced with

275 exposure temperature. Group-level effects were the same as for the water loss model. Only ecotypes
276 with five or more recordings were analysed (Nakagawa et al., 2017). The R^2_{marginal} and $R^2_{\text{conditional}}$
277 were calculated from the water loss and WU models using the 'r2_bayes' function from the
278 *performance* package (Lüdecke et al., 2021), and posterior predictive checks were presented in
279 Supplementary **Fig. S5**. All statistical outcomes are presented as mean posterior estimates \pm 95%
280 credible intervals (95% CI). The model results were also visualised in **Fig. S6–7**.

281 *Water loss vulnerability*

282 We estimated the impact of SSP2–4.5 (+2 °C) and SSP5–8.5 (+4 °C) scenarios on EWL globally
283 for a typical 8.7 g frog (geometric mean body mass of the study) using the 'ectotherm' function
284 from the *NicheMapR* package (Kearney and Porter, 2020) and compared the estimated EWL to the
285 spatial distribution of species richness. We extracted high resolution ($\sim 4\text{km}^2$) global dataset on the
286 mean VPD (kPA/month) and mean wind speed ($\text{m s}^{-1}/\text{month}$) from Abatzoglou et al. (2018) under
287 the current climate (1970–2000) and under a +2 and +4°C scenario to estimate EWL under the
288 current and future scenario. Wind speed at ground level (1 cm) was corrected based on the reference
289 height of 10 m and assuming a level surface roughness of 0.15 m in an open landscape
290 (<https://mrke.github.io/NicheMapR/inst/doc/microclimate-model-theory-equations>).

291 **Effect of environmental warming and drying on behavioural activity**

292 To demonstrate the role of water-conserving strategies on the potential hours for activity, we
293 simulated a hypothetical frog in Karawatha, Queensland, Australia. Karawatha has a record of daily
294 rainfall from 1994–current and this area has experienced drought recently (2017–2019). The aim is
295 to explore the interactions between different climate processes on frog activity rather than reflect
296 realistic changes in the ecosystem. We also note that these simulations can be applied to any
297 terrestrial location in the world with species-specific thermal and hydrological variables.

298 We used *NicheMapR* to construct a biophysical model of the water, and energy balance of a
299 typical adult frog (8.7 g) from Kearney and Porter (2020) and Kearney and Enriquez-Urzelai (2022)
300 to test the role of behavioural and physiological regulation on potential activity under different
301 thermal and hydric conditions. The default frog model was based on the leopard frog *Lithobates*
302 *pipiens* and was applied for all ecotypes because r_i and WU rate did not differ between ecotypes
303 (**Table S3–4**). The model was based on the 'thermodynamic niche' modelling scheme, considering
304 thermodynamic effects on the biophysical landscape (**Fig. S8a**). Three hypothetical frog models
305 with differing water saving strategies were simulated to estimate the influence of behaviour and
306 physiological modification on activity (summarised in **Table S5**): 1) a shade model where the frog
307 can thermo-hydroregulate with shade available (0–90% shade), but is not able to burrow or climb
308 up trees to regulate temperature and hydration (representative of a typical ground-dwelling frog), 2)

309 an arboreal model with water-proof skin (high skin resistance) and the ability to climb trees up to
310 150 cm high to regulate their body temperature and hydration (representative of arboreal frog), 3)
311 burrowing model where the frog is able to burrow underground to a maximum of 200 cm to
312 regulate body temperature and hydration (representative of fossorial species). Semi-aquatic and
313 stream-dwelling ecotypes were not simulated as they are typically found around permanent water
314 bodies, which allows them to quickly rehydrate; our models focused on activity change under
315 terrestrial conditions.

316 We simulated the potential hours (in a 24 hour day) active in a year (t_{act}) which represents
317 the suitable thermal and hydric conditions for the animal to move beyond their retreat to either catch
318 prey or find mates (Kearney and Porter, 2020). While frogs can survive losing up to 50% of their
319 body weight before death (Hillman et al., 2009), locomotor performance declines rapidly around
320 10–30% loss of water (Beuchat et al., 1984; Titon Jr et al., 2010; Anderson and Andrade, 2017;
321 Greenberg and Palen, 2021), and toads have been observed to seek shade after losing ~14% of their
322 water mass (Bartelt, 2000). Therefore, we used a threshold of 80% of hydrated body mass to stop
323 activity, as locomotion is highly impaired below these conditions.

324 For each hypothetical frog model, we simulated t_{act} under four climate conditions, 1) current
325 normal scenario (typical air temperature and rainfall), 2) a current drought scenario (typical air
326 temperature and low rainfall), 3) a warming normal scenario (+4°C only), and 4) a warming drought
327 scenario (+4°C and low rainfall). Vapour pressure was computed from the environmental
328 temperature and humidity values which directly affects EWL in the model (driven by the vapour
329 density gradient between the skin surface and the air), while the simulated soil water potential
330 influences WU rate (Kearney and Enriquez-Urzelai, 2022). The full detail of model specifications
331 and climate model verification are provided in the Supplementary Information “*NicheMapR*”
332 section and **Fig. S9**.

333 **RESULTS**

334 **Species richness, aridity, and drought risk**

335 Amphibian species richness is inversely correlated with the aridity, where the higher the AI, the
336 fewer amphibian species per grid cell (**Fig. 1a-c**). On average, most species classified as stream-
337 dwellers inhabit wetter regions (mean AI of 1.46 ± 0.62 s.d), while most fossorial species live in
338 drier regions (mean AI of 0.80 ± 0.61 s.d; **Fig. 1d**). Under an intermediate warming scenario of
339 +2°C, around 6.6% of areas occupied by anurans will increase to arid-like conditions. Humid
340 regions will reduce by 1.4%, dry sub-humid regions will increase by 2.5%, semi-arid regions will
341 increase by 1.4%, arid regions will increase by 5% and hyper-arid regions will decrease by 2.3%
342 (**Fig. 1e and 1f**). Under a high warming scenario of +4°C, around 33.5% of areas occupied by all

343 anurans will increase to arid-like conditions. Humid regions will reduce by 5.8%, dry sub-humid
344 regions will increase by 9%, semi-arid regions will increase by 8.4%, arid regions will increase by
345 13.2% and hyper-arid regions will increase by 3% (**Fig. 1e** and **1g**).

346 By 2080–2100, 15.4 % of regions occupied by anurans are expected to face a combination
347 of increased drought in relation to the three metrics we considered: intensity, frequency, and
348 duration. This is especially so in large areas of South America, northern America, and most of
349 eastern Europe under an intermediate emission scenario (SPP2–4.5; **Fig. 2a**). Under a high emission
350 scenario (SPP5–8.5), 36.1 % of regions occupied by anurans, mainly the Americas, southern Africa,
351 Europe, and southern Australia, will be at risk of increased exposure to drought according to all
352 three metrics (**Fig. 2b**). Anuran species-assemblages in the Amazon region had the highest risk
353 from the combination of high species richness and is predicted to be exposed to increases in all
354 three metrics under both intermediate and high emission scenario (**Fig. 2c–d**).

355 We estimate around 21% of regions that anurans occupy will be at risk of increasing drought
356 intensity under an intermediate scenario (**Extended Data Fig. 1a–b**). Under a high emission
357 scenario, 38% of the area anurans occupy will be at risk of increasing drought intensity (**Extended**
358 **Data Fig. 1c–d**). Anurans in central America, southern America, western and central Europe,
359 southern Africa, and southern Australia are expected to be subjected to increased average drought
360 frequency by 1–4 months per year under an intermediate emission scenario (41.1% of areas
361 occupied by anurans; **Extended Data Fig. 2a–b**), while the mean frequency of over 4 months per
362 year are considered rare (0.9%). 15.3% of regions occupied by anurans are expected to reduce in the
363 frequency of drought. Under a high emission scenario, anurans in central America, the Amazon
364 region, Chile, northern United States, and the Mediterranean regions are predicted to experience
365 increases in drought frequency by over 4 months per year (16.3% of areas occupied by amphibians
366 more than 4 months per year; **Extended Data Fig. 2c–d**). 27.4% of regions occupied by anurans are
367 not expected to change and 11% are expected to reduce in drought frequency.

368 The duration of drought will increase in most of the America's, Europe, southern and central
369 Africa, and southern Australia by 1–4 consecutive months under an intermediate emission scenario
370 (28.6% of areas occupied by anurans; **Extended Data Fig. 3a–b**). The mean durations of over 4
371 months occurred over a limited area (3.1%). Some areas are expected experience increases in
372 drought duration by 10 months in the northern United States, Honduras, the Amazon region,
373 Guyana, Chile, Spain, and Türkiye (1.6% of areas occupied by anurans; **Extended Data Fig. 3c–d**).

374 **Water loss and uptake**

375 Resistance to water loss (r_i) did not differ between ecotypes but varied by water-conserving
376 strategies (**Table S3; Fig. S6**). Frogs with water-proof skin (morphological), with a cocoon layer
377 during aestivation (physiological), and inside artificial hollow structures (behavioural) had higher r_i
378 than frogs with no specialised water conserving strategies, with the highest r_i from frogs that are in
379 aestivation (**Fig. S6b**). Body size, and flow rate influenced r_i , which explained 64% [60–68] of the
380 variation in r_i (**Table S3**). r_i increased with body mass with a scaling exponent of 0.10 [0.04–0.17],
381 and higher wind speed decreased r_i (-0.22 [-0.34–-0.10]). There was a moderate phylogenetic signal
382 for water loss ($\lambda = 0.38$ [0.04–0.65]). EWL was high for a 8.7 g frog in regions with high annual
383 VPD (indicator of dryness) such as hot, arid regions (**Fig. 3a**). Anuran species richness is negatively
384 related to the potential EWL where areas with climates that allow low EWL have a higher number
385 of anuran species (**Fig. 3b**). Under an intermediate and high-emission scenario, EWL increased the
386 most in arid regions (**Fig. 3c-d**), where EWL nearly doubles in the Sahara, Arabian, Taklamakan,
387 and Australian deserts under a high-emission scenario (**Fig. 3d**).

388 Water uptake did not differ between ecotypes or treatment temperature (**Table S4**).
389 However, water uptake was influenced by body size and initial hydration, both moderators
390 explained 69% [62–75] of the variation in water uptake (**Extended Data Fig. S5**). WU increased
391 with body mass with a scaling exponent of 0.80 [0.70–0.91], and more dehydrated frogs had higher
392 water uptake rates (-0.5 [-0.7– -0.04]). There was a strong phylogenetic signal for water uptake ($\lambda =$
393 0.78 [0.38–0.94]).

394 **Effect of environmental warming and drying on behavioural activity**

395 Biophysical simulations of the current climate scenario (no warming, no drought) in Karawatha,
396 Queensland, showed a hypothetical ground-dwelling frog could be active 3,971 hours out of 8,760
397 total hours of the year (45%) (**Fig. 4a**). The corresponding values for a water-proof arboreal frog
398 was 4,767 h (54%), and 4,286 h (49%) for a burrowing frog (**Fig. 4c and e**).

399 Under a warming climate scenario alone (+4 °C, no drought), the ground-dwelling frog
400 increased in potential activity by 346 h (8.8%), the waterproof arboreal frog by 197 h (3.6%), and
401 the burrowing frog by 428 h (10%) relative to the current climate scenario. The increase in activity
402 was driven by warmer winters (June to August) allowing for more activity during the coldest
403 quarter (**Fig. 4b, d, g, Table S6**). When restricted to the warmest quarter of the year (December to
404 February), where climate warming is predicted to have the greatest impact, warming decreased
405 activity by 8.5% for a ground-dwelling frog, 8.8% for a waterproof frog, and 4.6% for a burrowing
406 frog (**Table S6**).

407 Under a drought climate scenario (no warming, drought) matching the 2018–2019 historical
408 drought, the ground-dwelling frog decreased potential activity by 174 h (4.5%), the waterproof
409 arboreal frog by 192 h (4.5%), and the burrowing frog by 97 h (2.2%) across the year. The effects
410 of drought were highest in the warmest quarter with an average reduction in potential activity for all
411 ecotypes by 8.3% relative to the current scenario (**Table S6**). When simulating both warming and
412 drought climate scenarios together, the ground-dwelling frog decreases its potential activity by 188
413 h (17.7% reduction), the waterproof arboreal frog by 224 h (19%), and the burrowing frog by 125 h
414 (10%) relative to the current scenario in the warmest quarter. For all ecotypes, a reduction in rainfall
415 has a greater effect on yearly activity than an increase in temperature alone (**Table S6**).

416 **DISCUSSION**

417 Amphibians are the most threatened class of vertebrates, and the number of species directly
418 impacted by climate change has increased by 39% over the last decade (Luedtke et al., 2023). We
419 find the exposure risk of anurans to increasing environmental dryness is global with most regions
420 where anurans occupy increasing in the average aridity and in the intensity, frequency, and duration
421 of extreme drought events. The Amazon rainforest in South America is of particular concern
422 because high species diversity overlaps with high risk to increasing drought. Our results
423 demonstrate how combining the effects of temperature and environmental dryness is more
424 impactful for amphibians' biology, thus providing a more comprehensive understanding of the
425 vulnerability of ectotherms to climate change.

426 Stream-dwelling or semi-aquatic ecotypes are expected to have a greater risk of water loss
427 than arboreal or fossorial ecotypes due to differences in adaptation to differing water exposure in
428 their environment. However, similarities in skin resistance and WU across all ecotypes in our study
429 suggest habitat preference does not drive variation in water loss and WU as suggested in previous
430 studies (Thorson, 1955; Katz and Graham, 1980; Withers et al., 1984; Wygoda, 1984). Instead, a
431 species' risk of water loss may be more tightly related to physiological and behavioural traits. For
432 example, some arboreal species have lower rates of EWL due to their high skin resistance than
433 other ecotypes. The high skin resistance may be attributed to either secretions (mucus, lipids,
434 proteins) or iridophores (Lillywhite, 2006) that allow some “waterproof” species to have r_i values as
435 high as squamates, turtles, and crocodylians (Hillman et al., 2009). Concerning the vulnerability to
436 climate change, most anuran species (apart from waterproof and cocoon-forming frogs) lack
437 physiological ways to reduce EWL. Such species must therefore rely on behavioural
438 hydroregulation such as microhabitat shelter to reduce dehydration risk (Schwarzkopf and Alford,
439 1996; de Andrade and Abe, 1997; Seebacher and Alford, 2002), otherwise it is expected that EWL
440 may double in arid regions in open habitats under a high-emission scenario (**Fig. 3d**).

441 The spatial extent of increasing drought events tends to overlap with extreme thermal events
442 described by Murali et al. (2023). Heatwaves are typically most frequent and intense in arid regions
443 (Murali et al., 2023), but these areas tend to have few if any anuran species. Of serious concern are
444 anurans in tropical regions such as the Amazon borders and southeast Asia that may experience
445 extreme drought and extreme warming (Malhi et al., 2008; Phillips et al., 2009). The combined
446 effects drought and warming could act synergistically on physiological functions in ways that
447 further reduce potential activity and survival. For example, the thermal tolerance and preferred body
448 temperature tend to decrease as frogs are dehydrated or subjected to dry conditions (Mitchell and
449 Bergmann, 2016; Anderson and Andrade, 2017; Guevara-Molina et al., 2020). Likewise,
450 amphibians under heat stress will attempt to keep cool via evaporative cooling. When exposed to
451 40–50°C air temperatures, anurans can maintain a skin surface temperature around 35°C (**Fig. S11**).
452 However, this comes at a cost of increased EWL. The negative effects of high temperature on frog
453 locomotor performance is exacerbated by dehydration in anurans (Prest and Pough, 1989;
454 Walvoord, 2003). Because extreme atmospheric evaporative demand under rising temperatures and
455 less predictable rainfall will impact the biology and phenology of terrestrial organism that depend
456 on water, the interactions of water and temperature must be considered for accurate forecasting of
457 physiological vulnerability to climate change (Rozen-Rechels et al., 2019; Lertzman-Lepofsky et
458 al., 2020).

459 Dehydration impacts on behavioural activity are ecologically important because
460 physiological-based restrictions on activity can limit the ability of organisms to disperse (Wu and
461 Seebacher, 2022), and shorten the time window for breeding (Wassens et al., 2013). The reduction
462 in activity when dehydrated is likely due to decline in ability to synthesise ATP by aerobic means
463 and increase in glycolysis during locomotion under water stress (Gatten Jr, 1987). However,
464 drought may not only reduce foraging and mating seeking opportunities, but also the environmental
465 resource available to fuel locomotion. Environmental aridity is highly correlated with primary
466 productivity and, therefore, food availability (Janzen and Schoener, 1968; Qiu et al., 2023). Thus,
467 environmental drying restricts both the quantity of available food as well as the capacity to find it
468 (Lertzman-Lepofsky et al., 2020). Due to the paucity of amphibian studies directly linking
469 environmental drying and activity in the field (Tracy et al., 2013), our understanding of amphibian
470 sensitivity would benefit from field research focused on species-specific changes in activity in
471 regions to be most at risk to drying.

472 Physiological plasticity can increase an organism's resistance to climate change (Seebacher
473 et al., 2015). Whether amphibians can adapt to increasing environmental dryness will depend on
474 their acclimation capacity. Previous research have indicated amphibians acclimation to drier

475 conditions by increased skin resistance (Wygoda, 1988), and regenerated capillary beds in the skin
476 (Riddell et al., 2019) which help increase WU capacity. Whether these changes permit longer
477 activity under drier conditions requires further investigation, especially for small, narrowly
478 distributed species, as observed in insects (Hoffmann et al., 2003; Chown et al., 2011). Field
479 observations have shown frog community assemblages are larger in average size in warmer, drier
480 regions (Guo et al., 2019; Castro et al., 2021; Sheridan et al., 2022). Larger frogs have lower EWL
481 proportionally to smaller frogs, and have higher water storing capacity in the bladder (Hillman et
482 al., 2009). Therefore, we may observe phenotypic change in amphibian community composition,
483 where larger species may be able to cope better with increased environmental drying relative to
484 smaller species. Consideration of plasticity is important for evaluating species risk to climate
485 change because plasticity can often, but not always, buffer the effects of environmental stressors
486 (Wu and Seebacher, 2021), and not accounting for plasticity and evolutionary potential can
487 overestimate the impacts of environmental change on predicting species distributions (Gerick et al.,
488 2014; Kellermann et al., 2020).

489 To understand and manage the effects of climate change on biodiversity we must integrate
490 knowledge on the biologically relevant processes for different types of organism in different
491 habitats. Thermal constraints on activity are often a focus for studies of ectotherms but integrating
492 thermal and hydric constraints gives a deeper understanding of anurans and other wet-skinned
493 terrestrial organisms risk to climate change than temperature alone.

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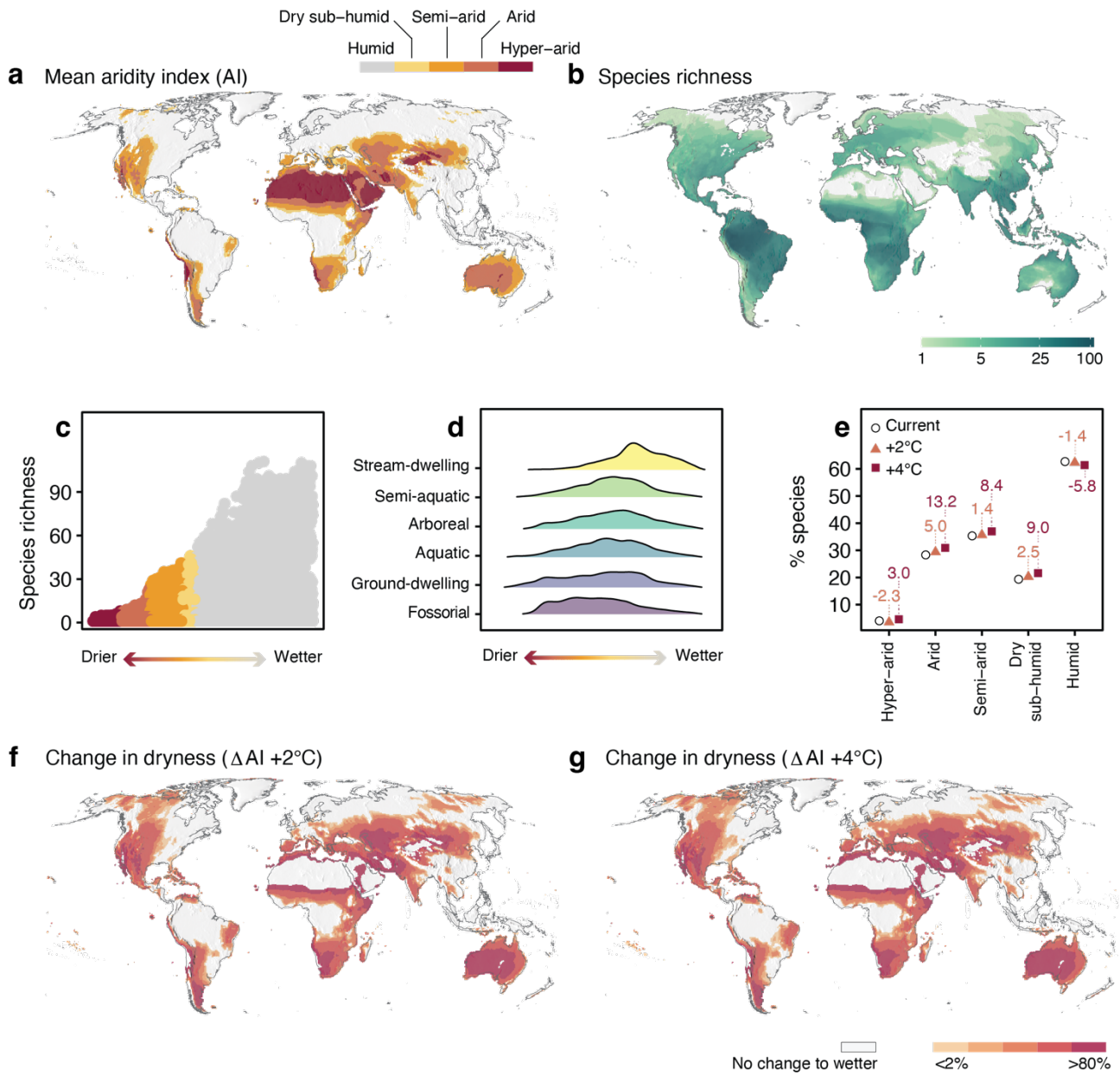
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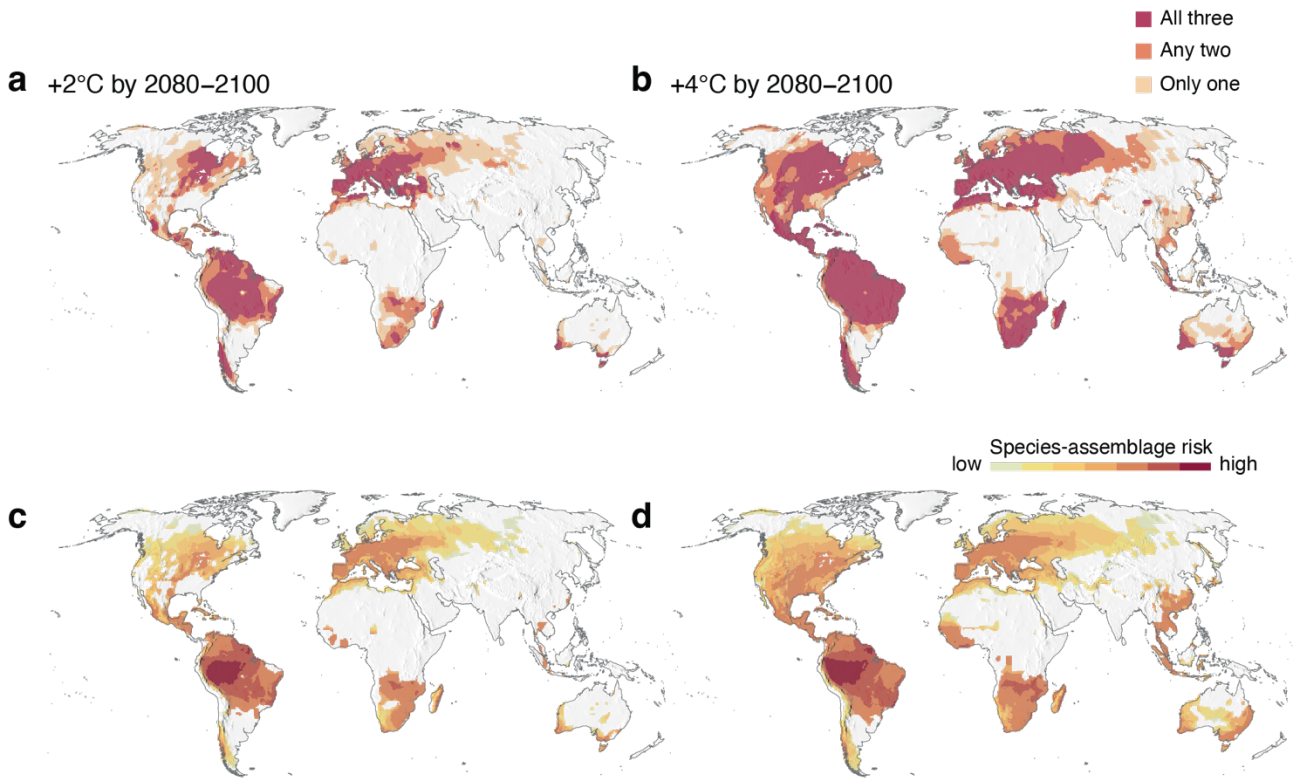
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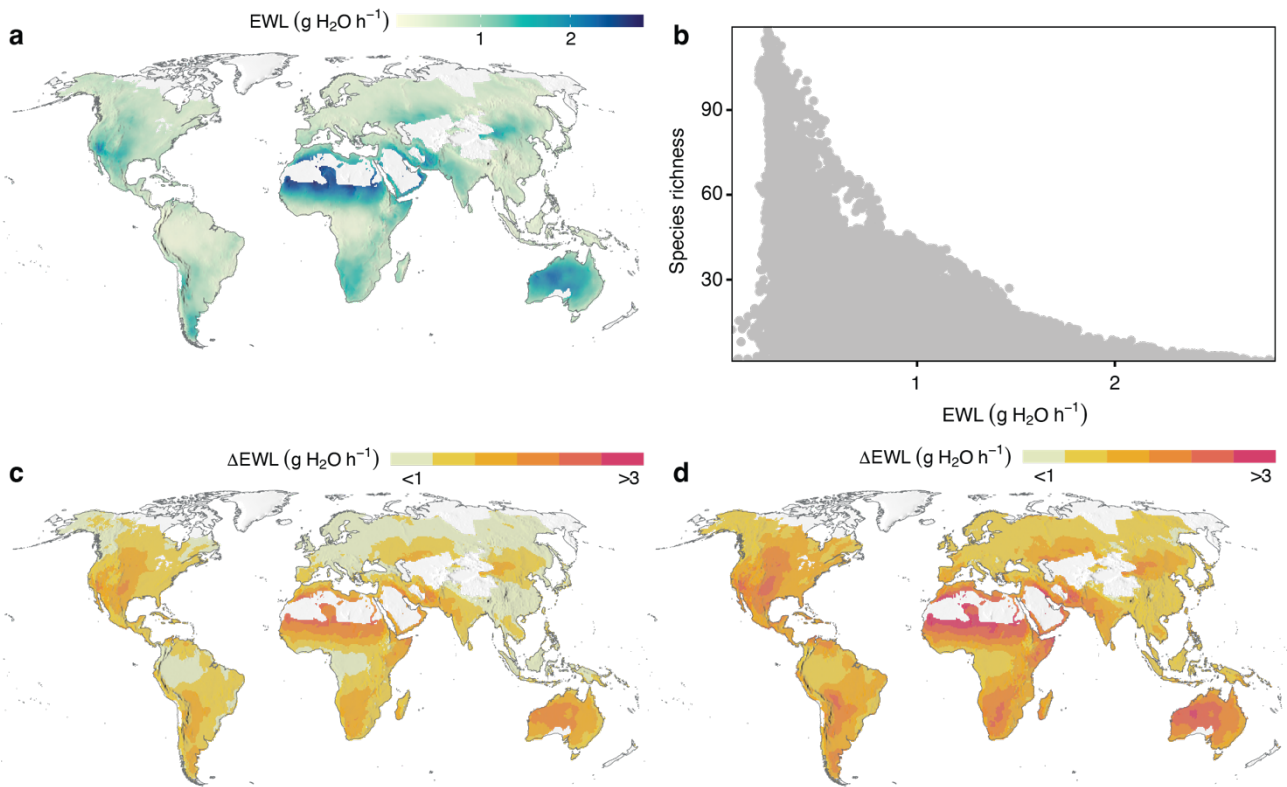
719 **Figure 1. Relationship between the spatial distribution of anuran species and the degree of**
 720 **dryness of the climate. (a)** The spatial distribution of the mean aridity index (AI) based on the
 721 precipitation and potential evapotranspiration between 1981–2010, and **(b)** is the anuran species
 722 richness based on the IUCN spatial assessment. **(c)** The relationship between mean AI and species
 723 richness, where the wetter the climate (lower aridity), the more species richness within the grid cell
 724 (0.5°) than drier climates. **(d)** Ecotype-specific distribution of anurans across AI, where stream-
 725 dwelling species tend to be conjugated in humid areas, while fossorial species tend to be found in
 726 drier areas. **(e)** Change in the percentage of species in each grid cell (0.5°) grouped by AI category
 727 between the current climate (yearly average from 1981–2010) and future warming scenario of $+2^{\circ}C$

728 and +4°C by 2080–2099. Under a +2°C warming scenario, there will be a -1.4 % decrease in
 729 percentage of species occupying humid regions and a 6.6% increase in percentage of species
 730 occupying dry sub-humid to arid regions. Under a +4°C warming scenario, there will be a -5.8 %
 731 decrease in percentage of species occupying humid regions and a 33.5% increase in percentage of
 732 species occupying dry sub-humid to arid regions. **(f–g)** Spatial change in mean dryness (decrease in
 733 AI) under a +2°C and +4°C warming scenario.



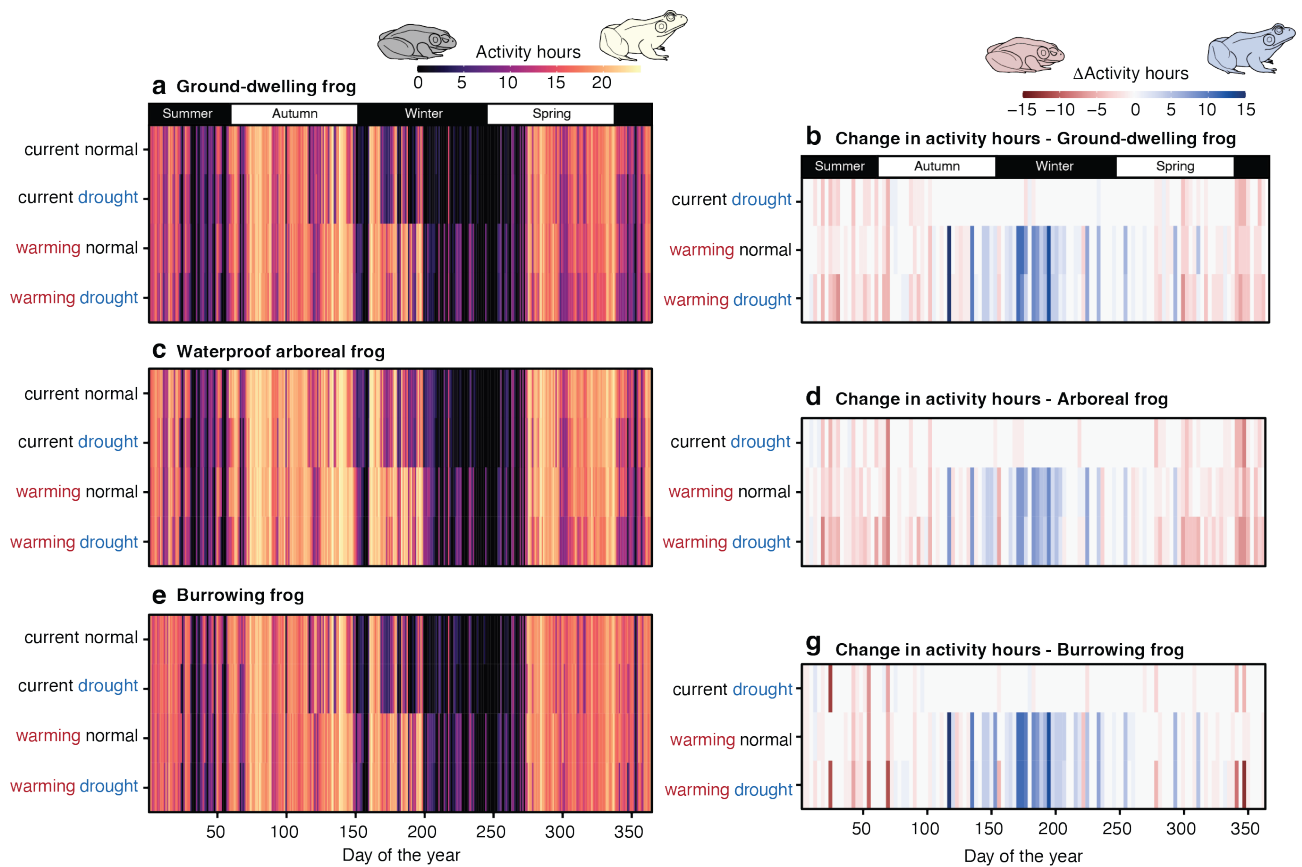
734

735 **Figure 2. Assemblage-level risk due to multiple aspects of increasing drought events.** Grid cells
 736 with anuran occurrences exposed to either a single metric (frequency, duration, or intensity: only
 737 one), any two metrics (frequency and intensity, frequency and duration, duration and intensity: any
 738 two), or all three metrics (frequency, duration, and intensity: all three) of extreme drought events
 739 beyond their historic levels by 2080–2100 under **(a)** a +2°C warming scenario (SSP2–4.5) and **(b)** a
 740 +4°C warming scenario (SSP5–8.5). The results for drought frequency, intensity, and duration
 741 alone were presented in the **Extended Data Figs. 1–3**. **(c–d)** Assemblage-level risk based on the
 742 number of drought risk by number of species per grid cell under **(c)** a +2°C warming scenario
 743 (SSP2–4.5) and **(d)** a +4°C warming scenario (SSP5–8.5). The drought risk was weighed by species
 744 assemblage, where a grid cell with high drought risk and high species richness has a higher risk
 745 score than a grid cell with high drought risk and low species richness.



746

747 **Figure 3. Variation in evaporative water loss (EWL) for a typical 8.7 g frog.** (a) Spatial
 748 variation in EWL ($\text{g H}_2\text{O h}^{-1}$) under the current (1981–2010) scenario. (b) Relationship between
 749 EWL and species richness where the number of species tends to be higher in areas that provide low
 750 EWL potential (Ecotype specific relationship is presented in **Supplementary Fig. S11**). Bottom
 751 maps represent change in EWL (ΔEWL ; $\text{g H}_2\text{O h}^{-1}$) under (c) a $+2^\circ\text{C}$ warming scenario (SSP2–4.5)
 752 and (d) a $+4^\circ\text{C}$ warming scenario (SSP5–8.5).



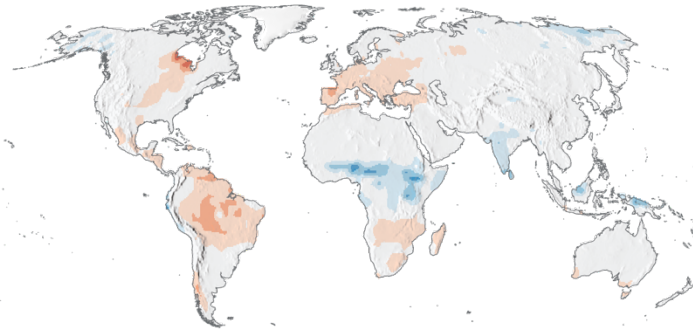
753

754 **Figure 4. Total hours for potential activity and change in activity under different climate**
 755 **conditions for a hypothetical 8.7 g frog in Queensland, Australia.** Daily hours of potential
 756 activity within suitable thermal and hydric conditions (left panels) and change in activity relative to
 757 the current normal scenario (right panels) for (a-b) a ground-dwelling frog, (c-d) a water-proof
 758 arboreal frog, and (e-g) a burrowing frog. Total potential activity represented as numbers presented
 759 in Supplementary Table S6.

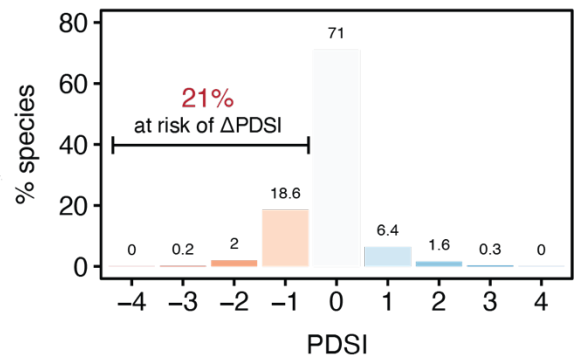
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761 **EXTENDED FIGURES**

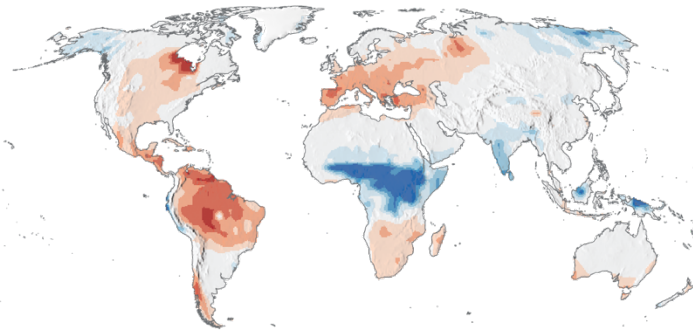
a Δ PDSI intensity (+2°C)



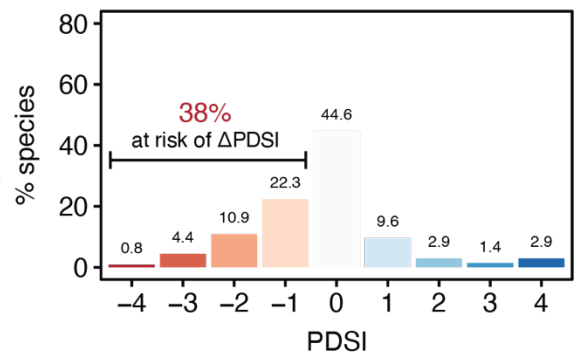
b Grid cell occupied +2°C



c Δ PDSI intensity (+4°C)



d Grid cell occupied +4°C



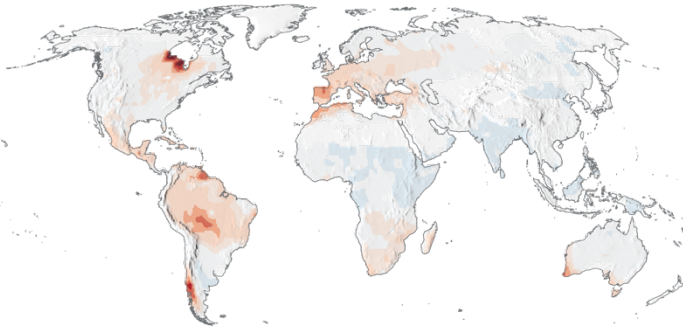
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763 **Extended Data Fig. 1 | Risk to increasing drought intensity for anurans by 2080–2100. (a)**

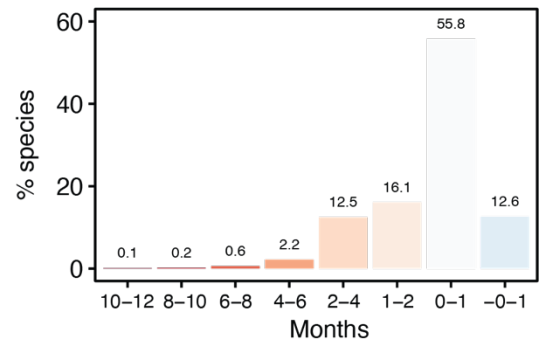
764 Change in the Palmer Drought Severity Index (Δ PDSI) under a +2°C warming scenario (Shared
 765 Socioeconomic Pathways 2–4.5; SPP2–4.5) by 2080–2100 relative to the current scenario (1970–
 766 1999). A decrease Δ PDSI indicates higher drought occurrences, while an increase Δ PDSI indicates
 767 more extreme wetness. **(b)** Percentage of anuran species occupancy in each PDSI category grid cell
 768 (0.5°) under a +2°C warming scenario, where 21% of species are in areas that are at risk of
 769 increasing drought. **(c)** Change in Δ PDSI under a +4°C warming scenario (SPP5–8.5). **(d)**
 770 Percentage of anuran species occupancy in each PDSI category grid cell (0.5°) under a +4°C
 771 warming scenario, where 38% of species are in areas that are at risk of increasing drought.

772

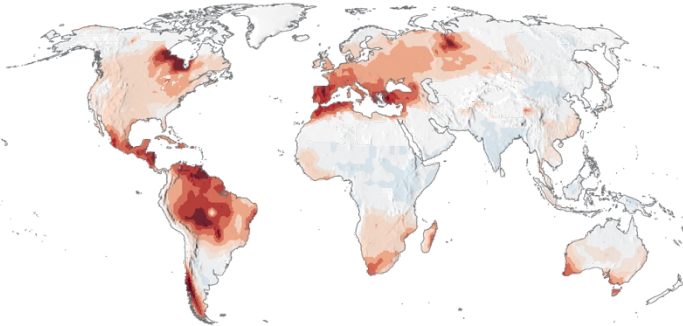
a Change in drought frequency (Δ PDSI +2°C)



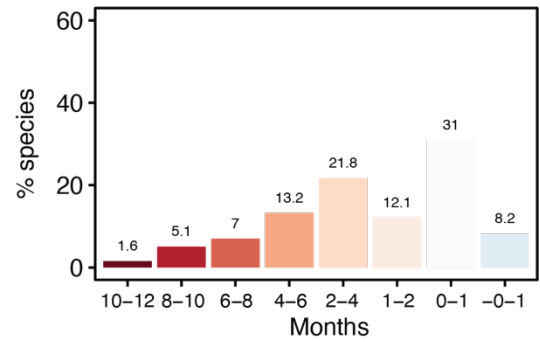
b Grid cell occupied +2°C



c Change in drought frequency (Δ PDSI +4°C)



d Grid cell occupied +4°C

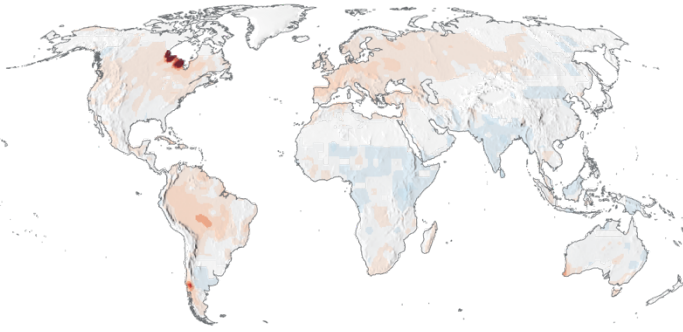


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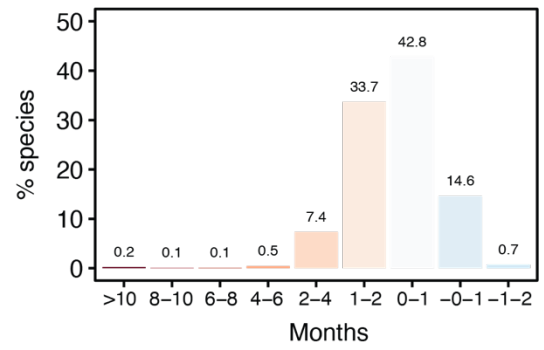
774 **Extended Data Fig. 2 | Risk to increasing drought frequency for anurans by 2080–2100. (a)**
775 **Change in the Palmer Drought Severity Index frequency (Δ PDSI_[frequency]) under a +2°C warming**
776 **scenario (Shared Socioeconomic Pathways 2–4.5; SPP2–4.5) by 2080–2100 relative to the current**
777 **scenario (1970–1999). Δ PDSI_[frequency] was defined as change in monthly drought below -2**
778 **(moderate to extreme drought) within a 20 year period. (b) Percentage of anuran species occupancy**
779 **in each frequency category grid cell (0.5°) under a +2°C warming scenario. (c) Change in the**
780 **Δ PDSI_[frequency] under a +4°C warming scenario (SPP5–8.5). (d) Percentage of anuran species**
781 **occupancy in each frequency category grid cell (0.5°) under a +4°C warming scenario.**

782

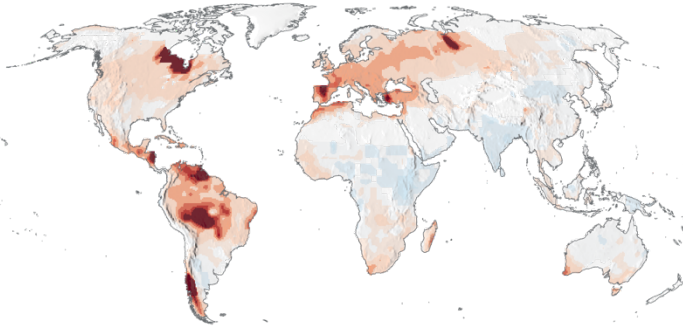
a Change in drought duration (Δ PDSI +2°C)



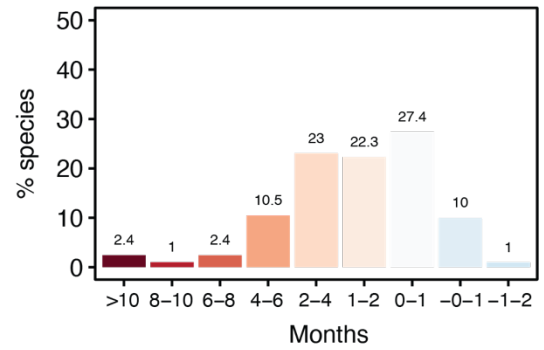
b Grid cell occupied +2°C



c Change in drought duration (Δ PDSI +4°C)



d Grid cell occupied +4°C

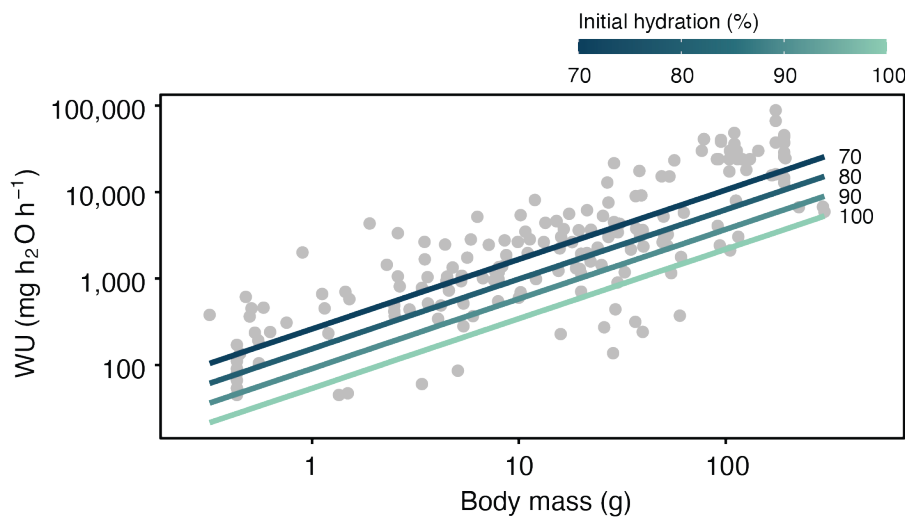


783

784 **Extended Data Fig. 3 | Risk to increasing drought duration for anurans by 2080–2100. (a)**

785 Change in the Palmer Drought Severity Index frequency (Δ PDSI_[duration]) under a +2°C warming
 786 scenario (Shared Socioeconomic Pathways 2–4.5; SPP2–4.5) by 2080–2100 relative to the current
 787 scenario (1970–1999). Δ PDSI_[duration] was defined as consecutive months under moderate to extreme
 788 drought (PDSI < -2) within a 20 year period. **(b)** Percentage of anuran species occupancy in each
 789 duration category grid cell (0.5°) under a +2°C warming scenario. **(c)** Change in the Δ PDSI_[duration]
 790 under a +4°C warming scenario (SPP5–8.5). **(d)** Percentage of anuran species occupancy in each
 791 duration category grid cell (0.5°) under a +4°C warming scenario.

792



793

794 **Extended Data Fig. 5 | Relationship between cutaneous water uptake with body mass of adult**
795 **anurans.** Gradient coloured lines represent the influence of initial hydration of the frog on the rate
796 of water uptake. Individual data points represent mean water uptake measurements ($n = 143$).
797