

1 Botany and Geogenomics: constraining geological hypotheses in the Neotropics with large-scale genetic  
2 data derived from plants

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## 16 Abstract

17 Decades of empirical research have revealed how the geological history of our planet shaped plant  
18 evolution by establishing well-known patterns (e.g., how mountain uplift resulted in high rates of  
19 diversification and replicate radiations in montane plant taxa). Under this approach, information is passed  
20 from geology to botany by interpreting data in light of geological processes. Instead, in this synthesis, I  
21 describe how by integrating natural history, phylogenetics, and population genetics, botanical research  
22 can inform our understanding of past geological and climatic processes. This conceptual shift aligns with  
23 the goals of the emerging field of geogenomics. In the Neotropics, plant geogenomics is a powerful tool  
24 for the reciprocal exploration of two long standing questions in biology and geology: how the dynamic  
25 landscape of the region came to be and how it shaped the evolution of the richest flora. Current challenges  
26 that are specific to analytical approaches for plant geogenomics are discussed. I describe the scale at  
27 which various geological questions can be addressed from biological data, and what makes some groups  
28 of plants excellent model systems for geogenomics research. Although plant geogenomics is discussed  
29 with reference to the Neotropics, the recommendations given here for approaches to plant geogenomics  
30 can and should be expanded to exploring long-standing questions on how the earth evolved with the use  
31 of plant DNA.

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33 Key words: admixture, biogeography, demographic modeling, geogenomics, geology, Neotropics,  
34 orogeny, phylogenetic networks, plant phylogenomics

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36 Geological history in the Neotropics has played a key role in shaping the evolution of the world's richest  
37 flora (Gentry, 1982; Hughes et al., 2013). Geomorphic processes in the region —driven greatly by rapid  
38 mountain building (Hoorn et al., 2010; Boschman, 2021)—, have led to identifiable evolutionary patterns.  
39 These patterns include accelerated diversification (e.g., explosive radiations in Andean bellflowers,  
40 *Lupinus*, and several other taxa living at high-elevations; (Hughes and Eastwood, 2006; Madriñán et al.,  
41 2013; Lagomarsino et al., 2016; Tribble et al., 2023), limited gene flow across landscape units (e.g., in  
42 *Amphirox longifolia* and two species of *Marathrum* across the Rio Negro and Andean mountains  
43 respectively; Nazareno et al., 2017; Bedoya et al., 2021), and replicated radiations (e.g., repeated  
44 evolution of leaf ecomorphs in a Neotropical clade of *Viburnum*; Donoghue et al., 2022). Plant evolution  
45 research has shed light on such patterns by interpreting biological data in reference to known geological  
46 events. Although past geological studies have provided insight into the timing of processes that led to the  
47 configuration of the dynamic and heterogeneous landscape matrix in the Neotropics (e.g., Hoorn et al.,  
48 2010; Montes et al., 2015), precise landscape models for the region carry large uncertainty and are  
49 unavailable across most landscape units through time.

50 The massive amount of genomic data currently being generated, the long-recognized role of landscape  
51 change in shaping the evolution of the Neotropical flora, and the current gap in our knowledge about the  
52 timing and process of formation of many landscape units in the region, presents an opportunity to use  
53 patterns inferred from large-scale genetic data to generate, constrain, or reduce uncertainty in landscape  
54 evolution models in the Neotropics. This approach aligns with the goal of Geogenomics, an emerging  
55 field (Baker et al., 2014) that involves the reciprocal integration of geologic, climatic, ecological,  
56 paleontological, and genomic data to test landscape evolution hypotheses from biological data (Badgley et  
57 al., 2017; Rahbek et al., 2019; Dolby et al., 2022). The reciprocal investigation and testing of biological  
58 and geomorphic processes in geogenomics research, sets it apart from other fields like comparative  
59 phylogeography, which search for landscape and climatic features that are correlated with patterns of

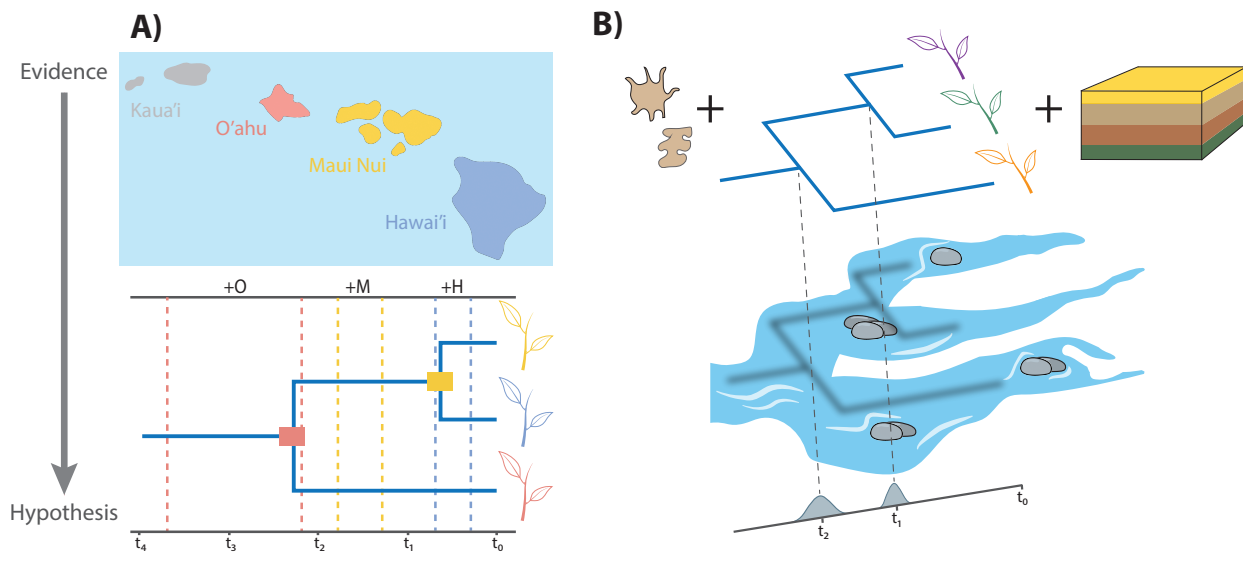
60 genomic variation in co-distributed taxa (Edwards et al., 2022), without explicitly testing models of  
61 landscape or climatic evolution.

62 Here I review recent contributions to plant geogenomics in the Neotropics that show how by integrating  
63 natural history, phylogenetics, and population genetics, botanical research can go beyond investigating  
64 patterns in light of landscape change and inform our understanding of past geologic and climatic  
65 processes in the Neotropics. I also discuss what makes some groups of plants ideal model systems for  
66 learning about past landscapes with the use of plant DNA, and offer practical considerations for plant  
67 geogenomics, including the scale at which geological hypotheses can be tested from genomic data.  
68 Research aligned with the goals of geogenomics and using animals as study systems (e.g., birds,  
69 mammals, and amphibians), has tested landscape hypotheses in the Neotropics (Albert et al., 2006;  
70 Boubli et al., 2015; Thom and Aleixo, 2015; Godinho and Da Silva, 2018; Naka et al., 2022; Rodriguez-  
71 Muñoz et al., 2022). However, responses to a common geological or climatic event across taxa are  
72 idiosyncratic (Donoghue and Smith, 2004; Bacon et al., 2015; Antonelli et al., 2018). Geogenomics  
73 research in plants is fundamental to a more holistic understanding of the geophysical history of the  
74 Neotropics, and instrumental to our understanding of how landscape changes contributed to the high plant  
75 richness in the region.

## 76 WHY AND HOW TO SHIFT THE PARADIGM?

77 *Plant evolutionary patterns can inform paleogeographic dynamics*— Since the proposal of the  
78 ‘riverine barrier’ hypothesis by A.R. Wallace in 1852, historical biogeography has aimed to study  
79 distributional patterns of organisms with reference to landscape units. Research in this area has  
80 tremendously advanced our understanding of plant evolution on our constantly changing planet.  
81 Examples where dated phylogenies were interpreted according to landscape model reconstructions  
82 abound (e.g., biogeographic studies in Asteraceae, Bromeliaceae, Lamiaceae, Malpighiaceae,  
83 Myristicaceae, Orchidaceae, and Solanaceae (Davis et al., 2002; Bell, 2005; Givnish et al., 2011, 2016;  
84 Dupin et al., 2017; Frost et al., 2017, 2022; Mandel et al., 2019; Rose et al., 2022). A more recent

85 approach to modeling species geographic ranges is biogeographic dating, a time-stratified framework  
 86 where molecular, paleogeographic, and biogeographic evidenced are jointly modeled to estimate time-  
 87 calibrated phylogenies and species ancestral ranges. Under this approach, paleogeographic hypotheses  
 88 impose constraints on the rates of biogeographic change and vice versa (Landis, 2017). For example,  
 89 estimated dates of island formation in the Hawaiian archipelago (Lim and Marshall, 2017) were  
 90 incorporated as the maximum ages at which dispersal to each island could have taken place in a  
 91 biogeographic study of Hawaiian silverswords (Asteraceae; Landis et al., 2018; Fig. 1A). This analytical  
 92 framework was expanded to re-examine the ancestral biome reconstruction of *Viburnum* (Landis et al.,  
 93 2021).



95 Figure 1. Approaches to questions at the intersection of geology and plant evolution, from evidence to hypothesis  
 96 generation. A) Biogeographic dating *sensu* Landis et al., 2017; 2018. Modeling of ancestral ranges is conducted by  
 97 imposing proposed times for the formation of islands in the Hawaiian archipelago (top) as constraints on the rates of  
 98 biogeographic change (bottom). B) A geogenomics approach, where divergence times for three populations of river  
 99 plants are used to constrain the approximate timing and pattern of connections of paleodrainages through time.  
 100 Fossil (e.g., phytoliths) and geologic (e.g., stratigraphic) data would further provide support for the formulated  
 101 hypothesis.

102 A major assumption of these approaches is that there is a model of the context in which plants evolved in  
 103 a given area. The example above is unusual in the extent to which we know the exact timing of island

104 formation in the Hawaiian archipelago. However, unifying, accurate, and precise landscape or climatic  
105 models are unavailable for a majority of regions and time scales. This gap can be explained by uncertainty  
106 in analytical tools for dating strata (Schoene et al., 2013), the complexity of past landscape dynamics, and  
107 the high degree of heterogeneity of topographic units. While geological advances continue to improve  
108 paleogeographic reconstructions, biological information can be used for inference of geologic events  
109 (Wegener, 1966). Today's exponential growth of genomic data presents an opportunity to improve what  
110 we know about landscape configurations over time using a geogenomics approach. In the Neotropics,  
111 plant geogenomics research is a tool for the reciprocal exploration of how the dynamic landscape of the  
112 region came to be and how it shaped the evolution of the richest flora.

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114 *Selection of study systems, reciprocity in geogenomics and the problem of scale* — Figure 1B  
115 shows how inferred divergence times of river plants could be used to propose a hypothesis for fluvial  
116 connections in the past. The efficacy of this approach is dependent on the temporal and spatial  
117 concordance of geological and biological processes. In other words, geogenomics research requires  
118 careful selection of model systems whose evolutionary histories reflect that of landscape units, as  
119 organismal responses to geomorphic barriers are highly heterogeneous (Araya-Donoso et al., 2022;  
120 Rodriguez-Muñoz et al., 2022). A famous example of this heterogeneity is the variable timing of dispersal  
121 in terrestrial organisms (i.e., plants, birds, mammals, reptiles, amphibians, and arthropods) across the  
122 isthmus of Panama (Bacon et al., 2015). The hypothesis-testing approach shown in Figure 1B would  
123 necessitate that the plants studied live strictly in riverine habitats, are naturally distributed across rivers,  
124 and have limited gene flow across drainage basins (e.g., low dispersal ability, unsuccessful establishment  
125 of migrant populations, or other pre or pro-zygotic barriers to gene flow).

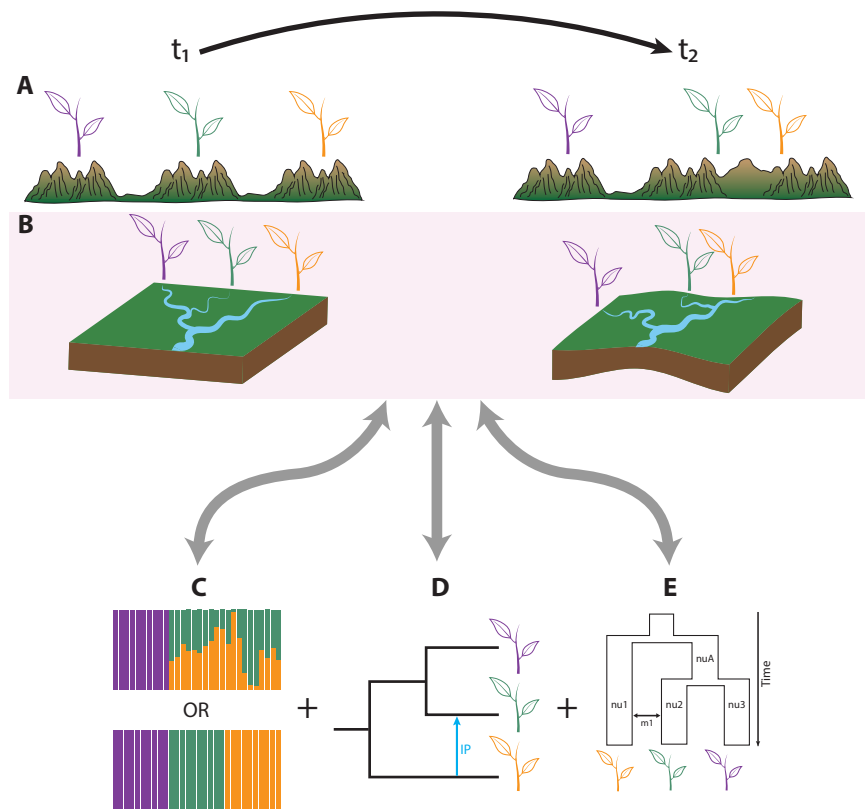
126 By concurrently interpreting geological and biological data, geogenomics reduces biases  
127 introduced from using geological data alone—which inherently carry uncertainty—to constrain  
128 biological processes (e.g., divergence times). It also better informs the interpretation of biological data

129 (e.g., dated phylogenies or algorithms for detection of gene flow events) to constrain geological  
130 hypotheses. For example, although both geological and biological models carry uncertainty, a temporal  
131 and spatial correspondence between a geological hypothesis (i.e., timing of mountain building) and  
132 biological processes for a given study system (i.e., inferred divergence times), would support such  
133 geological hypothesis. Patterns inferred from genomic data (Fig. 2C-E) for allopatric operational  
134 taxonomic units would further inform the study system's responses to mountain building. Then,  
135 biological patterns inferred for this same study system can be used to formulate hypotheses of mountain  
136 building of a different topographic unit, or of other landscape events linked to mountain building. The  
137 paleontological record could further inform plant responses to some geological events (e.g., (Strömberg,  
138 2005; Stiles et al., 2020; Carvalho et al., 2021).

139 Estimating the timing of geomorphic events from genomic data is scale-dependent. The example shown  
140 in Figure 1B is an oversimplification of a deep time (i.e., millions of years) event spanning large  
141 geographic distances. Here, the connections of river drainages— not of single rivers—are inferred. On the  
142 other end of the spectrum, the genomic signature of a landscape event (Fig. 2) could be erased by  
143 demographic processes and the strong influence of genetic drift on plants with small population sizes.  
144 Thus, geologic events may be reconstructed most reliably only if occurring at shallow time scales. For  
145 example, using genomic data to estimate the timing of connections across high-elevation ecosystems in  
146 topographically complex landscapes is likely to be limited to recent dispersal and vicariance events. For  
147 example, páramo plants experienced demographic changes due to the contraction and expansion of high-  
148 elevation ecosystems during the Pleistocene glaciations (Flantua et al., 2019), only the most recent of  
149 which may be identifiable from genomic data (Nevado et al., 2018).

150 *Genomic approaches to plant geogenomics*— Genomic signatures of past geologic events may  
151 be identified with four main approaches: 1) estimation of divergence times of species/populations  
152 distributed across landscape units (Fig. 1B); 2) identification of the distribution of genetic variation across  
153 the landscape (i.e., genetic structure; Fig. 2C); 3) investigation of introgression events (e.g., phylogenetic

154 network inference, summary methods for the detection of gene flow from site patterns, and likelihood-  
 155 based methods for inference of introgression; Fig. 2D; Green et al., 2010; Durand et al., 2011; Blischak et  
 156 al., 2018; Nauheimer et al., 2021; Ji et al., 2023); and 4) demographic modeling (i.e., using site frequency  
 157 spectra and whole genomes; Fig. 2E; Adams and Hudson, 2004; Gutenkunst et al., 2009; Li and Durbin,  
 158 2011; Schiffels and Durbin, 2014).



159

160 Figure 2. Hypothetical example of three lines of genomic evidence that can be used to identify landscape change  
 161 events from biological data. Two hypothetical examples presented correspond to uplift and connection of mountain  
 162 segments A), and a river capture event B) taking place at time  $t_2$ . These events resulted in secondary contact and  
 163 admixture of two previously isolated plant populations (orange and green). Admixture linked to geologic processes,  
 164 leaves a signature in the genetic constitution of populations that may be identifiable from C) inference of admixture  
 165 proportions, D) identification of introgression events through phylogenetic network inference, and E) modeling of  
 166 past demographic dynamics. Bidirectional arrows indicate how geologic processes can inform the interpretation of  
 167 genomic patterns, which can then be used to test geologic hypotheses (see '*Selection of study systems, reciprocity in*  
 168 *geogenomics and the problem of scale*' for how geogenomics avoids circularity). IP: Inheritance probability, or the  
 169 proportion of genes contributed by each parental population to a reticulation node, if gene trees are used for network  
 170 inference.



171 Figure 2 shows a hypothetical example of uplift and connection of mountain segments (A) and a river  
172 capture event (B) taking place at time  $t_2$  and generating patterns of genomic variation that can be  
173 identified with the approaches mentioned above. The examples assume lack of plant dispersal across  
174 landscape units. The hypothetical events resulted in secondary contact and admixture of two previously  
175 isolated populations of plants (green and orange), exemplifying spatial and temporal correspondence of  
176 geological and biological processes. Gene flow following secondary contact may be revealed by an  
177 admixture plot (Fig. 2C, top). However, signatures of admixture could be eroded by subsequent events  
178 such as genetic drift (Fig. 2C, bottom), particularly in small populations (e.g., plant populations in  
179 savanna pockets within the Amazon, sky islands like the páramos, or river-rapids). Exploration of non-  
180 bifurcating relationships through phylogenetic network inference could also detect the branches of the  
181 tree involved in the reticulation event (green and orange populations; Fig. 2D), and the inheritance  
182 probability (IP; the proportion of genes contributed by each parental population to a reticulation node, if  
183 gene trees are used for network inference; Solís-Lemus et al., 2017). Ideally, the demographic model in  
184 Fig. 2E would be identified as the best fit for genomic data generated from the three populations.  
185 Variations of this model would include permitting changes in population size and asymmetric gene flow.  
186 Concordant patterns across approaches would facilitate interpretation of the resulting patterns as  
187 indicators of a specific past geomorphic process (recent connections of previously disconnected mountain  
188 segments (Fig. 2A) and river capture events (Fig. 2B). However, the gene flow patterns detected in Fig. 2  
189 could also result from long-distance dispersal across landscape units. Phylogeographic and population  
190 structure analyses of the study systems would be necessary to determine the extent and direction of gene  
191 flow across populations/species. This would inform the extent to which these patterns are indicative of an  
192 underlying landscape change event. Sedimentological and stratigraphic data in the area where the species  
193 are distributed would further support the landscape change hypothesis formulated. The above highlights  
194 the reciprocal integration of geologic, climatic, and genomic data to test landscape evolution hypotheses  
195 from biological data *sensu* Dolby et al., 2022.

196 Rigorously constraining landscape and climatic hypotheses with the use of time-trees, admixture plots,  
197 phylogenetic networks, and inferred demographic models requires the joint application of at least two of  
198 the four approaches for various reasons. Putative hybrid ancestry in admixture proportions might instead  
199 result from the retention of ancestral polymorphisms from shared ancestry (Lawson et al., 2018),  
200 scenarios that can often be distinguished through demographic modeling. Furthermore, the signatures of  
201 polymorphism in truly admixed populations are expected to be removed with time (Clark, 1997). There  
202 are also practical limits to existing algorithms: phylogenetic network inference is computationally feasible  
203 only for a small number of terminal taxa, does not identify the direction at which gene flow took place,  
204 and may suffer from non-identifiability as complexity of the model increases (Yu et al., 2014; Pardi and  
205 Scornavacca, 2015; Solís-Lemus and Ané, 2016; Solís-Lemus et al., 2017). Similarly, demographic  
206 models inferred from genetic data can only be applied to a limited number of populations (Farleigh et al.,  
207 2021), suffer from non-identifiability (Terhorst and Song, 2015), and require *a priori* specifications of  
208 parameters and demographic events (unknown in most empirical systems; Li and Durbin, 2011; Loog,  
209 2021).

210 Caveats important but not exclusive to plant geogenomics include the scarcity of fossil evidence useful  
211 for calibration for many taxa. (Crane et al., 2004; Schenk, 2016). Dated phylogenies for a large number of  
212 plant groups are inferred from secondary calibrations and are subject to potential bias (Schenk, 2016).  
213 Gene flow can also affect inference of divergence times (Leaché et al., 2014). Inferred chronograms are a  
214 means to exploring plausible evolutionary and landscape change scenarios, but they should be taken  
215 cautiously and re-examined as more sources of data and analytical tools become available. Issues more  
216 specific to plants include escalating costs of sequencing of genomic-scale data for taxa with big genomes,  
217 widespread hybridization, polyploidization, and various mating systems (Stebbins, 1950), as well as the  
218 difficulty of developing analytical tools that accommodate these complexities (Blischak et al., 2023; but  
219 see Garrison and Marth, 2012; Serang et al., 2012; Blischak et al., 2018b; Gerard et al., 2018; Clark et al.,  
220 2019 for genotyping tools that enable ploidy specification and algorithms for genotype likelihood

221 estimation). Other challenges include the identification of paralogous and orthologous loci prior to  
 222 phylogenetic inference (Yang and Smith, 2014; Johnson et al., 2016; Gardner et al., 2021; Morales-  
 223 Briones et al., 2022; Freyman et al., 2023; Mendez-Reneau et al., 2023), a scarcity of resources for  
 224 analysis of shared ancestry in polyploids (but see Kolář, 2021; Shastry et al., 2021), and demographic  
 225 modeling under various ploidy and mating systems scenarios (but see Roux and Pannell, 2015; Blischak  
 226 et al., 2023; Roux et al., 2023). Simulation studies testing the effect of strategies and violations of model  
 227 assumptions in plant geogenomics research (Schenk, 2016; Stift et al., 2019) are critical to better  
 228 constrain and propose plausible geological scenarios from biological data in light of uncertainty.

## 229 PLANT EVOLUTION AND GEOGENOMICS IN THE NEOTROPICS, OR, HOW ELSE ARE 230 PLANTS SPECIAL

231 The original proposal of the geogenomics neologism (Baker et al., 2014) identified major geologic  
 232 questions in the Neotropics to be addressed with biological data. The groundwork for many of these  
 233 questions consists of a model of the geologic history of northern South America since the Paleogene by  
 234 Hoorn et al., 2010, who provided a review of paleontological, sedimentological, ecological, and  
 235 geological data for the region. Below I show how plant evolution research conducted over the past  
 236 decade, exemplify how the hypotheses laid out by Hoorn et al., 2010, by Baker et al., 2014, and by the  
 237 several prior studies that they were based on, can be constrained from biological data.

### 238 *Fine-scale uplift history of different segments and Cordilleras of the Northern Andes—*

239 Landscape models indicate that the uplift of the Andes (which in northern South America are split into the  
 240 Western, Central, and Eastern Cordilleras) was an asynchronous process that started in the Late  
 241 Cretaceous (Horton, 2018; Gianni et al., 2020). A current challenge is to generate a finer-scale geological  
 242 model that reflects the asynchrony and spatial heterogeneity of the Andes, particularly for the Western  
 243 and Central cordilleras (Sanín et al., 2022b). A recent study by Sanín et al., 2022b tackled this challenge  
 244 and revised the geologic history of uplift across the three mountain ranges that constitute the northern

245 Andes, using genomic data generated for the *Geonoma undata*–*G. orbignyana* species complex  
246 (Arecaceae).

247 The study integrated inference of genetic structure, phylogenetic relationships, divergence times, and  
248 demographic modeling with information from past geological studies in the region (Velandia et al., 2005;  
249 Mora et al., 2008; Veloza et al., 2012; Anderson et al., 2016; Montes et al., 2019, 2021; Mora-Páez et al.,  
250 2019; Noriega-Londoño et al., 2020). They provide evidence for a scenario where mountain segments  
251 were already uplifted but disconnected until the Pliocene, when continuous cordilleras were formed.

252 Phylogeographic breaks coincide with the location of strike-slip faults (i.e., vertical fractures where  
253 blocks have moved horizontally). The authors suggest that extensional and compressional features of the  
254 faults may represent initial negative topography that is subsequently inverted (Dewey et al., 1998; Veloza  
255 et al., 2012; Montes et al., 2019; Zapata et al., 2023). This resulted in the connection of previously  
256 isolated mountain segments during the Pliocene. Geological samples of ignimbritic rocks (i.e., a type  
257 volcanic rocks) further indicate that vulcanism contributed to mountain connectivity and the closure of  
258 low mountain passes in the Plio-Pleistocene ( $< \sim 2$  Ma; Sanín et al., 2022a). This resulted in increased  
259 dispersal of high-elevation mountain palms (i.e., *Ceroxylon*) throughout the northern Andean cordilleras.  
260 Phylogeographic inference in the *Tococa-Azteca* plant-ant mutualism (Torres Jimenez et al., 2021) and  
261 biogeographic modeling in the Rubiaceae (Antonelli et al., 2009) further support the prior existence of  
262 lowland passes across the Andean mountains, and the dispersal through lowland portals in frogs,  
263 mammals, and non-avian reptiles (Rodriguez-Muñoz et al., 2022).

264 ***Drainage basin reconfiguration through time***— In addition to being a major driver of  
265 Neotropical diversification (Antonelli et al., 2009; Antonelli and Sanmartín, 2011; Sklenář et al., 2011;  
266 Smith et al., 2014; Hoorn et al., 2018), Andean uplift reconfigured watersheds across South America  
267 (Albert et al., 2006; Hoorn et al., 2010; Ruokolainen et al., 2019). However, the configuration of  
268 paleodrainages through time remains largely unknown. To propose a hypotheses for past river  
269 connections in northern South America, I have used two species of strictly riverine plants (*Marathrum*,

270 Podostemaceae) with broad distributions across the Andes as a model system. Collecting genomic data  
271 for populations from different drainage basins, I inferred population structure, phylogenetic networks,  
272 divergence dates, and population summary statistics (Bedoya et al., 2021). After establishing that gene  
273 flow is limited in populations across drainage basins via population structure analyses, I used the timing  
274 and pattern of population divergence across river drainages as evidence to argue that drainage basins  
275 became separated at ~12 and ~4 Ma as a result of major pulses of Andean uplift (Gregory-Woodzicki,  
276 2000; Garziona et al., 2008; Hoorn et al., 2010; Mora et al., 2010; Anderson et al., 2016; Boschman,  
277 2021). This paleodrainage model was subsequently validated by a paleogeographic model of northern  
278 South America through the Miocene, built from geochronological, petrographic, and thermochronological  
279 data (Zapata et al., 2023). Strictly riverine plants are a great study system in geogenomics in the  
280 Neotropics. Limited gene flow has been reported for various species in the Podostemaceae (Baggio et al.,  
281 2013; Katayama et al., 2016). However, long-distance dispersal may explain the pantropical distribution  
282 of some groups in the family (Kita and Kato, 2004; Koi et al., 2015; Ruhfel et al., 2016). Research in this  
283 group should be expanded to investigate the link between river structure, distance between river rapids  
284 within and across rivers, and plant dispersal. This to better understand the prevalence of range evolution  
285 through river capture events and long distance dispersal in the group, and assess to what extent geologic  
286 hypotheses can be formulated from genomic data across space in the group.

287 Additional botanical studies have introduced promising study systems to constrain hypotheses for river  
288 reconfiguration in the past, including *Mauritia flexuosa* (Arecaceae), whose populations are inferred to  
289 have strong genetic structure across drainage basins in the Amazon (Sander et al., 2018). Other examples  
290 include *Amphirrox longifolia* (Violaceae) and *Buchenavia oxycarpa* (Combretaceae), tree species with  
291 limited gene flow across broad portions of the Rio Negro in the Amazon basin (Nazareno et al., 2017,  
292 2019).

293 *Neogene and Pleistocene marine incursions in the Neotropics*— Landscape change models  
294 developed from palynological, sedimentological, geochemical, and macrofossil data indicate that Pebas

295 Lake, a vast wetland-like system of > 1 million km<sup>2</sup>, was in place in northern and central South America  
296 from ca. 23–10 Ma (Hoorn, 1993; Wesselingh et al., 2002; Hoorn et al., 2010; Linhares et al., 2017;  
297 Jaramillo, 2023). Isotopic and palynological evidence shows a marine influence on this water body, but  
298 conflicting interpretations of the Miocene depositional environment and extent of marine incursions at the  
299 time remain unresolved (Frailey et al., 1988; Hoorn, 1993, 2006; Räsänen et al., 1995; Webb, 1995;  
300 Vonhof et al., 2003; Latrubesse et al., 2007, 2010; Jaramillo et al., 2017).

301 Based on distributional information, dispersal biology, and calibrated phylogenies of plant taxa, Bernal et  
302 al., (2019) proposed the hypothesis that the current occurrence of plants from estuarine or coastal zones in  
303 western Amazonia is the legacy of Miocene marine incursions in the region. Divergence times of  
304 populations of *Pachira aquatica* (Malvaceae) and *Manicaria saccifera* (Arecaceae) suggest that  
305 Amazonian and coastal populations differentiated as a result of the complete retreat of the marine-  
306 influenced embayment. Appropriate testing of this hypothesis to reconstruct the extent of past estuarine  
307 environments requires further geogenomics studies, for which the authors provide a list of 28 candidate  
308 species from deltaic or estuarine environments in the region. Evidence for past estuarine connections from  
309 botanical data also include a study of a relict inland mangrove ecosystem (*Rhizophora mangle*) in the  
310 Yucatán peninsula (Aburto-Oropeza et al., 2021). Population genetic analyses of genotyping-by-  
311 sequencing data, together with floristics, sedimentological, paleontological evidence, and sea-level  
312 modeling, point to the mangrove relict reaching its current inland location in the Yucatán peninsula  
313 during last interglacial (ca. 120 K years).

#### 314 FUTURE CONSIDERATIONS FOR PLANT GEOGENOMICS

315 Given their putatively limited gene flow across landscape units, distributional patterns, and affinity to  
316 specific habitats, many plant groups are excellent systems for geogenomics research. Recent botanical  
317 investigations demonstrate that the analysis of data derived from plant DNA can improve our knowledge  
318 of topography, riverine, and habitat connections through time.

319 Geogenomics research advocates for building interdisciplinary networks (e.g., Sanín et al., 2022a and the  
320 Baja GeoGenomics consortium; <https://bajageogenomics.org>). Future questions in geogenomics would  
321 necessitate such integrative collaborations. For example, investigating the extent of expansion,  
322 contraction, and flickering connectivity of high elevation Neotropical grasslands (i.e., páramos) requires  
323 reciprocally testing the concordance of inferred timing and patterns of gene flow (e.g., in *Lupinus*;  
324 Nevado et al., 2018), and landscape models derived from palynological evidence during the Pleistocene  
325 (e.g., Flantua et al., 2019). Exploring concordance across other páramo plant taxa, may further support or  
326 reject landscape hypotheses for how high-elevation grasslands connected and disconnected in the past.

327 The timing of formation, long-term persistence, geographical isolation, and expansion or contraction of  
328 seasonally dry tropical forests (SDTFs), are other future questions to be addressed with plant geogenomics  
329 in the Neotropics. To explore these questions, biological (e.g., divergence time estimation from sequence  
330 data and biogeographic modeling; Pennington et al., 2004; Särkinen et al., 2012; Fernandes et al., 2022)  
331 and paleontological data (Toby Pennington et al., 2000; Burnham and Carranco, 2004; Mayle et al., 2004;  
332 Werneck et al., 2011; Martínez et al., 2020, 2021), have been provided independently. Integrating these  
333 findings into a geogenomics framework would involve selecting study systems in the SDTFs where  
334 geomorphic, ecological (Pennington et al., 2006), and evolutionary processes are concordant, and  
335 applying the approaches here described.

336 This synthesis focuses on recent and future approaches to understanding how the richest flora and one of  
337 the most species-rich landscapes evolved through time. However, the recommendations given here for  
338 plant geogenomics would allow for the exploration of questions at the intersection of geology and biology  
339 across the globe (e.g., the origin of disjunct arid and semi-arid regions across the Americas, and the  
340 evolution of arid-adapted species with amphitropical distributions). At a time when rapid technological  
341 innovation allows for the generation of massive amounts of genomic data, classical botanical knowledge  
342 is critical for identifying good study systems where plant DNA can help us address long-standing  
343 questions on earth's geologic history.

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