

1 Botany and Geogenomics: constraining geological hypotheses with large-scale genetic data
2 derived from plants

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16 Abstract

17 Past studies in plant phylogenetics have shed light on how the geological history of our planet shaped
18 plant evolution by establishing well-known patterns (e.g., how mountain uplift resulted in high rates of
19 diversification and replicate radiations in montane plant taxa). Under this approach, information is
20 transferred from geology to botany, by interpreting data in light of geological processes. In this synthesis,
21 I propose a conceptual shift in this traditional approach to specifically transfer information from botany to
22 geology. This conceptual shift follows the goals of the emerging field of geogenomics and emphasizes
23 that plant phylogenetics can go beyond investigating patterns in light of landscape change, to reduce the
24 inherent uncertainty in models of paleotopography, river system structure, and land connections through
25 time. Current challenges that are specific to analytical approaches for plant geogenomics are discussed. I
26 describe the scale at which various geological questions can be addressed from biological data, and what
27 makes some groups of plants excellent model systems for geogenomics research. This synthesis
28 highlights the critical role of classical botanical knowledge in identifying good study systems to unveil
29 long-standing questions on how the earth evolved with the use of plant DNA.

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31 Key words: admixture, biogeography, demographic modeling, geogenomics, geology, orogeny,
32 phylogenetic networks, plant phylogenomics

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34 Geological history has played a key role in shaping plant evolution (Gentry, 1982; Hughes et al., 2013;
35 Sundaram et al., 2019). Geomorphic processes have resulted, among others, in the formation of islands
36 and elevational gradients (Humboldt and Bonpland, 1807; Clague and Sherrod, 2014; Rahbek et al.,
37 2019), leading to identifiable evolutionary patterns. These patterns include accelerated diversification
38 (e.g., explosive radiations in Andean bellflowers and *Lupinus*; Hughes and Eastwood, 2006; Lagomarsino
39 et al., 2016), convergent evolution (e.g., evolution of reproductive traits in Hawaiian Lobeliads and of life
40 forms in temperate and arctic-alpine *Androsace*; Givnish et al., 2009; Roquet et al., 2013), limited gene
41 flow across landscape units (e.g., in *Amphirrox longifolia* and two species of *Marathrum* across the Rio
42 Negro and Andean mountains respectively; Nazareno et al., 2017; Bedoya et al., 2021), and replicated
43 radiations (e.g., repeated evolution of leaf ecomorphs in a Neotropical clade of *Viburnum*; Donoghue et
44 al., 2022). Plant phylogenetics research has shed light on such patterns by interpreting biological data in
45 reference to geological processes. Here, I propose a shift in this traditional approach, emphasizing that
46 plant phylogenetics can go beyond investigating patterns in light of landscape change, to specifically
47 transfer information from botany to geology (Bedoya et al., 2021).

48 The use of patterns inferred from large-scale genetic data to constrain and reduce uncertainty in
49 geological hypotheses, was defined as the goal of geogenomics by Baker et al. (2014). This emerging
50 field involves the reciprocal integration of geologic, climatic, and genomic data to test landscape
51 evolution hypotheses at various time scales from biological data (Dolby et al., 2022). Here I describe how
52 by integrating natural history, phylogenetics, and population genetics, botanical research can inform
53 geological processes. In this synthesis, I also contribute practical considerations to geogenomics that are
54 specific to plant research. I review studies that, by introducing promising systems to constrain geological
55 hypotheses and informing current models of defining geological and climatic events, exemplify my
56 description on how to shift the paradigm in plant phylogenetics.

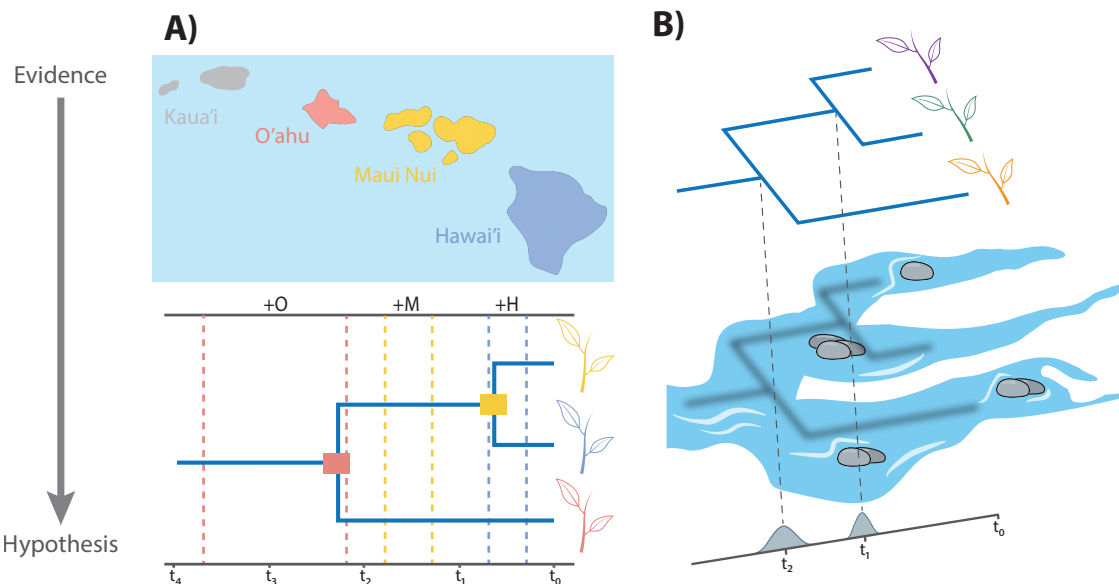
57 WHY AND HOW TO SHIFT THE PARADIGM?

58 *Plant evolutionary patterns can inform paleogeographic dynamics*— Since the proposal of the
59 ‘riverine barrier’ hypothesis by A.R. Wallace in 1852, historical biogeography has aimed to study
60 distribution patterns of organisms with reference to landscape dynamics. Research in this area has
61 tremendously advanced our understanding of plant evolution on our constantly changing planet.
62 Examples where dated phylogenies were interpreted according to landscape model reconstructions
63 abound (e.g., biogeographic studies in Asteraceae, Bromeliaceae, Lamiaceae, Malpighiaceae,
64 Myristicaceae, Orchidaceae, and Solanaceae (Davis et al., 2002; Bell, 2005; Givnish et al., 2011, 2016;
65 Dupin et al., 2017; Frost et al., 2017, 2022; Mandel et al., 2019; Rose et al., 2022). A recent and more
66 sophisticated approach to modeling species geographic ranges is biogeographic dating, a time-stratified
67 framework where paleogeographic hypotheses impose constraints on the rates of biogeographic change
68 and vice versa (Landis, 2017). This framework allows for the incorporation of uncertainty in estimates of
69 formation and change of landscape units in the past. For example, proposed times of island formation in
70 the Hawaiian archipelago (Lim and Marshall, 2017) were incorporated as the maximum ages at which
71 dispersal to each island could have taken place in a biogeographic study of Hawaiian silverswords
72 (Asteraceae; Landis et al., 2018; Fig. 1A). This analytical framework was expanded to re-examine the
73 ancestral biome reconstruction of *Viburnum* (Adoxaceae; Landis et al., 2021).

74 A major assumption of these approaches is that there is a paleogeographic or climatic model of the
75 context in which plants evolved in a given area. The example above is unusual in the extent to which we
76 know exact dates for island formation in the Hawaiian archipelago. However, unifying, accurate, and
77 precise landscape models, are unavailable for a majority of regions and across time scales. This lack of
78 precise models can be explained by uncertainty in analytical tools for dating strata (Schoene et al., 2013),
79 the complexity of past landscape dynamics, and the high degree of heterogeneity of topographic units.
80 While geological advances continue to improve paleogeographic reconstructions, the growing amount of
81 genomic data being generated in current times presents an opportunity to improve what we know about

82 landscape configurations over time. A shift in the paradigm of using landscape models to inform
 83 biological processes to instead using biological data (i.e., large-scale genetic data) to constrain geological
 84 hypotheses (Fig 1B), is a powerful approach to explore long-standing questions in geology and plant
 85 evolution.

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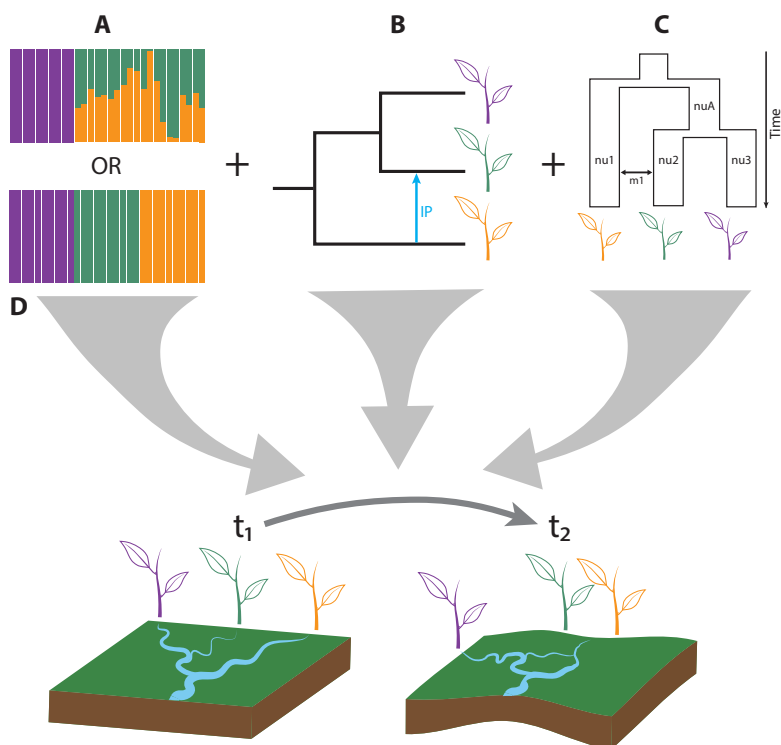
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88 Figure 1. Approaches to questions at the intersection of geology and plant evolution from evidence to generation of
 89 a hypothesis. A) Biogeographic dating *sensu* Landis et al., 2017; 2018. Modeling of ancestral ranges is conducted by
 90 imposing proposed times for the formation of islands in the Hawaiian archipelago (top) as constraints on the rates of
 91 biogeographic change (bottom). B) Geogenomics approach, where divergence times for three populations of river
 92 plants are used to constrain the approximate timing and pattern of connections of paleodrainages through time.

93

94 *Approaches to plant geogenomics*— Genomic signatures of past geologic events may be
 95 identified with four main approaches: 1) estimation of divergence times of species/populations distributed
 96 across landscape units (Fig. 1B); 2) identification of the distribution of genetic variation across the
 97 landscape (i.e., genetic structure; Fig. 2A); 3) investigation of introgression events (e.g., phylogenetic
 98 network inference, summary methods for the detection of gene flow from site patterns, and likelihood-

99 based methods for inference of introgression; Fig. 2B; Green et al., 2010; Durand et al., 2011; Blischak et
 100 al., 2018; Nauheimer et al., 2021; Ji et al., 2023); and 4) demographic modeling (i.e., using site frequency
 101 spectra and whole genomes; Fig. 2C; Adams and Hudson, 2004; Gutenkunst et al., 2009; Li and Durbin,
 102 2011; Schiffels and Durbin, 2014).



103
 104 Figure 2. Toy example of three lines of evidence that can be used to identify a landscape change event from
 105 biological data. Patterns seen from A) inference of admixture proportions, B) identification of introgression events
 106 through phylogenetic network inference, and C) modeling of past demographic dynamics. Such patterns indicate
 107 carry a signal left by D) a river capture event that took place at time t_2 . The event resulted in secondary contact and
 108 admixture of two previously isolated plant populations (orange and green). IP: Inheritance probability or the
 109 proportion of genes contributed by each parental population to a reticulation node, if gene trees are used for network
 110 inference.

111
 112 The example in Fig. 2 shows a river capture event (i.e., the diversion of part of the course of a river into
 113 another river; Fig. 2D) taking place at time t_2 , and the patterns that could be identified with the
 114 approaches mentioned above. The river capture event resulted in secondary contact and admixture of two

115 previously isolated populations of river plants (green and orange), as identified in an admixture plot (Fig.
116 2A, top). However, signatures of admixture could be eroded by subsequent events such as genetic drift
117 (Fig. 2A, bottom), particularly in small populations (e.g., plant populations in savanna pockets within the
118 Amazon rain forest, sky islands like the páramos, or river-rapids). Exploration of non-bifurcating
119 relationships through phylogenetic network inference could also detect the branches of the tree involved
120 in the reticulation event (green and orange populations; Fig. 2B), as well as the inheritance probability
121 (IP; the proportion of genes contributed by each parental population to a reticulation node, if gene trees
122 are used for network inference; (Solís-Lemus et al., 2017). Ideally, the 3D demographic model in Fig. 2C
123 would be identified as the best fit for genomic data generated from the three populations (purple, green
124 and orange). Variations of this model would include changes in population size and asymmetric gene
125 flow. Concordant patterns across approaches would facilitate interpretation of the resulting patterns as
126 indicators of a specific past geomorphic process.

127 Constraining landscape and climatic hypotheses with the use of time-trees, admixture plots, phylogenetic
128 networks, and inferred demographic models as suggested above requires the joint application of at least
129 two of the four approaches for various reasons. For example, if patterns of admixture are identified, they
130 could be a result of retention of ancestral polymorphisms from shared ancestry (Lawson et al., 2018).
131 Also, the signatures of polymorphism in truly admixed populations are expected to be removed with time
132 (Clark, 1997). There are also practical limits to existing algorithms: phylogenetic network inference is
133 computationally feasible only for a small number of terminal taxa, does not identify the direction at which
134 gene flow took place, and may suffer from non-identifiability as complexity of the model increases (Yu et
135 al., 2014; Pardi and Scornavacca, 2015; Solís-Lemus and Ané, 2016; Solís-Lemus et al., 2017). In
136 addition, demographic models inferred from genetic data have been developed for a limited number of
137 populations (Farleigh et al., 2021), suffer from non-identifiability (Terhorst and Song, 2015), and require
138 *a priori* specifications of parameters and demographic events (unknown in most empirical systems; Li
139 and Durbin, 2011; Loog, 2021).

140 Special caveats specific but not exclusive to plant geogenomics include the scarcity of fossil evidence
141 useful for calibration for many plant groups (Crane et al., 2004; Schenk, 2016). Other caveats derive from
142 the prevalence of hybridization, polyploidization, and various mating systems in plants (Stebbins, 1950)
143 and the difficulty in developing analytical tools that accommodate these complexities (Blischak et al.,
144 2023). However, some standard genotyping tools that enable ploidy specification and algorithms for
145 genotype likelihood estimation are currently available (Garrison and Marth, 2012; Serang et al., 2012;
146 Blischak et al., 2018b; Gerard et al., 2018; Clark et al., 2019). Other challenges include the identification
147 of paralogous and orthologous loci prior to phylogenetic inference (Yang and Smith, 2014; Johnson et al.,
148 2016; Gardner et al., 2021; Morales-Briones et al., 2022; Freyman et al., 2023; Mendez-Reneau et al.,
149 2023), scarcity of resources for analysis of shared ancestry in polyploids (but see Kolář, 2021; Shastry et
150 al., 2021), and demographic modeling under various ploidy and mating systems scenarios (but see Roux
151 and Pannell, 2015; Blischak et al., 2023; Roux et al., 2023). Simulation studies testing the effect of
152 strategies and violations of model assumptions in plant geogenomics research (Schenk, 2016; Stift et al.,
153 2019) are critical to better constrain and propose plausible geological scenarios from biological data in
154 light of uncertainty.

155 *Selection of study systems and the problem of scale*— Figure 1B shows how inferred divergence
156 times of river plants could be used to propose a hypothesis for fluvial connections in the past. The
157 efficacy of this approach requires careful selection of model systems whose evolutionary histories reflect
158 that of rivers, as organismal responses to geomorphic barriers are highly heterogeneous (Araya-Donoso et
159 al., 2022; Rodriguez-Muñoz et al., 2022). A famous example of differing responses to a common event is
160 the variable timing of dispersal in terrestrial organisms (i.e., plants, birds, mammals, reptiles, amphibians,
161 and arthropods) across the isthmus of Panama (Bacon et al., 2015). The hypothesis-testing approach
162 shown in Figure 1B would necessitate that the plants studied live strictly in riverine habitats, is naturally
163 distributed across rivers, and has a low dispersal ability. Once taxa that meet these characteristics are

164 found, interpretation of analyses derived from genomic data for such study systems should consider other
165 aspects of their biology like population size, mating systems, life cycle, and ploidy.

166 Estimating the timing of geomorphic events from genomic data is scale-dependent and the limitations
167 mentioned above determine the type of research questions that are feasible. The example shown in Figure
168 1B is an oversimplification of a deep time (millions of years), large-scale problem, where the connections
169 of river drainages and not of single rivers are inferred. This is the scale at which research questions are
170 approached by classical biogeography. At a smaller-scale, the genomic signature of changes in river
171 connections (Fig. 2) could be erased by changes in population size, and by the possibly large effect of
172 genetic drift in removing polymorphisms in river plants with small population sizes (Bedoya et al., 2021).
173 Thus, river capture events may be reconstructed most reliably only if occurring at a shallow time scale.
174 Similarly, estimating the timing of connections across high-elevation ecosystems based on genomic data
175 might be limited to the most recent events. For example, páramo plants experienced demographic changes
176 due to the contraction and expansion of high-elevation ecosystems during the Pleistocene glaciations
177 (Flantua et al., 2019), the most recent of which are identifiable from genomic data (Nevado et al., 2018).

178 PLANT EVOLUTION IN GEOGENOMICS RESEARCH OR HOW ELSE ARE PLANTS 179 SPECIAL

180 The proposal of the geogenomics neologism (Baker et al., 2014) included the identification of major
181 geologic questions in the Neotropics to be addressed with biological data. These included the history of
182 mountain building (e.g., Andean uplift), river reconfiguration (e.g., the Amazon river), and formation of
183 bridges between terrestrial ecosystems (e.g., closure of the Isthmus of Panama). The ground work for
184 many of these problems consists of a model of the geologic history of northern South America since the
185 Paleogene by Hoorn et al., 2010, which provided a review of paleontological, sedimentological,
186 ecological, and geological data for the region. Over the past decade, research using evidence derived from

187 plants has helped constrain hypotheses laid out by Hoorn et al., 2010, by Baker et al., 2014, and by the
188 several prior studies that they were based on.

189 ***Fine-scale uplift history of different segments and Cordilleras of the Northern Andes—***

190 Landscape models indicate that the uplift of the Andes (which in northern South America are split into the
191 Western, Central, and Eastern Cordilleras) was an asynchronous process that started in the Late
192 Cretaceous (Horton, 2018; Gianni et al., 2020). A current challenge is to generate a finer-scale geological
193 model that reflects the asynchrony and spatial heterogeneity of the Andes, particularly for the Western
194 and Central cordilleras (Sanín et al., 2022). A recent study by Sanín et al., 2022 tackled this challenge and
195 revised the geologic history of uplift across the three mountain ranges that constitute the northern Andes,
196 using genomic data generated for the *Geonoma undata*–*G. orbignyana* species complex (Arecaceae).

197 The study integrated inference of genetic structure, phylogenetic relationships, divergence times, and
198 demographic modeling, with information from past geological studies in the region (Velandia et al., 2005;
199 Mora et al., 2008; Veloza et al., 2012; Anderson et al., 2016; Montes et al., 2019, 2021; Mora-Páez et al.,
200 2019; Noriega-Londoño et al., 2020). They provide evidence for a scenario where mountain segments
201 were already uplifted but disconnected until the Pliocene, when continuous cordilleras were formed.

202 Phylogeographic breaks coincide with the location of strike-slip faults (i.e., vertical fractures where
203 blocks have moved horizontally). The authors suggest that extensional and compressional features of the
204 faults may represent initial negative topography that is subsequently inverted (Dewey et al., 1998; Veloza
205 et al., 2012; Montes et al., 2019; Zapata et al., 2023). This resulted in the connection of previously
206 isolated mountain segments during the Pliocene. Phylogeographic inference in the *Tococa-Azteca* plant-
207 ant mutualism (Torres Jimenez et al., 2021) and biogeographic modeling in the Rubiaceae (Antonelli et
208 al., 2009), further support the existence of lowland passes across the Andean mountains, and the dispersal
209 through lowland portals in frogs, mammals, and non-avian reptiles (Rodríguez-Muñoz et al., 2022).

210 ***Drainage basin reconfiguration through time across the northern Andes—*** In addition to being
211 a major driver of Neotropical diversification (Antonelli et al., 2009; Antonelli and Sanmartín, 2011;

212 Sklenář et al., 2011; Smith et al., 2014; Hoorn et al., 2018), Andean uplift reconfigured watersheds across
213 South America (Albert et al., 2006; Hoorn et al., 2010; Ruokolainen et al., 2019). However, the
214 configuration of paleodrainages through time remains largely unknown. To propose a hypotheses for past
215 river connections in northern South America, I have used two species of strictly riverine plants
216 (*Marathrum*, Podostemaceae) with broad distributions across the Andes as a model system. I used
217 genome-scale data for populations from different drainage basins to infer population structure,
218 phylogenetic networks, divergence dating, and population summary statistics (Bedoya et al., 2021). After
219 establishing that gene flow is limited in populations across drainage basins, the timing and pattern of
220 population divergence across river drainages was used as evidence to propose that drainage basins
221 became separated at ~12 and ~4 Ma as a result of major pulses of Andean uplift (Gregory-Woodzicki,
222 2000; Garziona et al., 2008; Hoorn et al., 2010; Mora et al., 2010; Anderson et al., 2016; Boschman,
223 2021). This paleodrainage model was subsequently validated by a paleogeographic model of northern
224 South America through the Miocene, built from geochronological, petrographic, and thermochronological
225 data (Zapata et al., 2023).

226 Additional botanical studies have introduced promising study systems to constrain hypotheses for river
227 reconfiguration in the past, including *Mauritia flexuosa* (Arecaceae), whose populations are inferred to
228 have strong genetic structure across drainage basins in the Amazon (Sander et al., 2018). Another
229 example is *Amphirrox longifolia* (Violaceae), a species with limited gene flow across broad portions of
230 the Rio Negro in the Amazon basin (Nazareno et al., 2017). These plant examples complement studies in
231 animals that are constrained to river drainages (Albert et al., 2006; Picq et al., 2014; Tagliacollo et al.,
232 2015). Comparative phylogeographic studies at various geographic scales for these and other plant
233 systems with putative limited dispersal ability could constitute powerful tools to generate hypotheses for
234 river configurations in the past.

235 *Neogene and Pleistocene marine incursions in the Neotropics—*

236 Landscape change models developed from palynological, sedimentological, geochemical, and microfossil
237 data indicate that the Pebas Lake, a vast wetland-like system of > 1 million km², was in place in northern
238 and central South America from ca. 23–10 Ma (Hoorn, 1993; Wesselingh et al., 2002; Hoorn et al., 2010;
239 Linhares et al., 2017; Jaramillo, 2023). Isotopic and palynological evidence shows a marine influence on
240 the Pebas lake, but conflicting interpretations of the Miocene depositional environment and extent of
241 marine incursions at the time remain (Frailey et al., 1988; Hoorn, 1993, 2006; Räsänen et al., 1995;
242 Webb, 1995; Vonhof et al., 2003; Latrubesse et al., 2007, 2010; Jaramillo et al., 2017).

243 Based on distributional information, dispersal biology, and calibrated phylogenies of plant taxa, Bernal et
244 al., 2019 proposed the hypothesis that the current occurrence of plants from estuarine or coastal zones in
245 western Amazonia is the legacy of Miocene marine incursions in the region. Divergence times of
246 populations of *Pachira aquatica* (Malvaceae) and *Manicaria saccifera* (Arecaceae) suggest that
247 Amazonian and coastal populations differentiated as a result of the complete retreat of the marine-
248 influenced embayment. Appropriate testing of this hypothesis to reconstruct the extent of past estuarine
249 environments requires further geogenomics studies, for which Bernal et al., 2019 provide a list of 28
250 candidate species from deltaic or estuarine environments in the region. Evidence for past estuarine
251 connections from botanical data also include a study of a relict inland mangrove ecosystem (*Rhizophora*
252 *mangle*) in the Yucatán peninsula (Aburto-Oropeza et al., 2021). Population genetic analyses from
253 genotyping by sequencing data, together with floristics, sedimentological, paleontological evidence, and
254 sea-level modeling, point to the mangrove relict reaching its current inland location in the Yucatán
255 peninsula during last interglacial (ca. 120 K years).

256 FUTURE CONSIDERATIONS FOR PLANT GEOGENOMICS

257 Geogenomics research is a call to conduct research that specifically transfers information from biology to
258 geology, making use of the wealth of information that can be extracted from large-scale genomic data.
259 This sets it apart from other fields like comparative phylogeography, which in search of finding landscape

260 and climatic features that drive shared genomic splits across co-distributed taxa (Edwards et al., 2022),
261 could in fact inform geologic models without necessarily attempting to explicitly test models of landscape
262 or climatic evolution. In addition to a paradigm shift to explicitly constrain geomorphic hypotheses,
263 geogenomics research advocates for interdisciplinary networks. An example of this is the Baja
264 GeoGenomics consortium (BGGc; <https://bajageogenomics.org>), a research group of geologists,
265 conservation geneticists, and evolutionary biologists working with mammals, reptiles, and plants. The
266 group aims to constrain how the physical landscape of the Baja California Peninsula has changed over the
267 last 6 My, and its impact on the evolutionary trajectories of organisms in the region.
268 Given their limited dispersal, distribution patterns, and affinity to specific habitats, many plant groups are
269 excellent systems for geogenomics research. Recent botanical investigations demonstrate that the analysis
270 of data derived from plant DNA can improve our knowledge of paleotopography (Sanín et al., 2022),
271 river system structure (Sander et al., 2018; Bedoya et al., 2021), habitat connections (Bernal et al., 2019;
272 Aburto-Oropeza et al., 2021), or biome configuration through time (Särkinen et al., 2012). In times when
273 advancing technologies allow the generation of massive amounts of genomic data, classical botanical
274 knowledge is critical to identify good study systems to unveil long-standing questions on how the earth
275 evolved with the use of plant DNA.

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