

1 Developing systems theory in soil agroecology:
2 Incorporating heterogeneity and dynamic instability

3 Nicholas Medina ^{1*} John Vandermeer ¹

4 ¹ Ecology and Evolutionary Biology, University of Michigan, Ann Arbor, MI USA

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12 *Correspondence: Nicholas Medina, nmedina@umich.edu, [https://orcid.org/0000-0001-](https://orcid.org/0000-0001-5465-3988)
13 5465-3988

14 Abstract

15 Ecosystem management is integral to the future of soils, yet anthropogenic drivers represent
16 a key source of uncertainty in ecosystem models. First- and new-generation soil models
17 formulate many soil pools using first-order decomposition, which tends to generate simpler
18 yet numerous parameters. Systems or complexity theory, developed across various scientific
19 and social fields, may help improve robustness of soil models, by offering consistent assump-
20 tions about system openness, potential dynamic instability and distance from commonly
21 assumed stable equilibria, as well as new analytical tools for formulating more generalized
22 model structures that reduce parameter space and yield a wider array of possible model
23 outcomes, such as quickly shrinking carbon stocks with pulsing or lagged respiration. This
24 paper builds on recent perspectives of soil modeling to ask how various soil functions can be
25 better understood by applying a complex systems lens. We synthesized previous literature
26 reviews with concepts from non-linear dynamical systems in theoretical ecology and soil sci-
27 ences more broadly to identify areas for further study that may help improve the robustness
28 of soil models under the uncertainty of human activities and management. Three broad dy-
29 namical concepts were highlighted: soil variable memory or state-dependence, oscillations,
30 and tipping points or hysteresis. These themes represent less intuitive yet key dynamics that
31 can emerge after assuming nuanced observations, such as reversibility of organo-mineral as-
32 sociations, dynamic aggregate- and pore hierarchies, persistent wet-dry cycles, higher-order
33 microbial community and predator-prey interactions, cumulative legacy land use history,
34 and social management interactions and/or cooperation. We discuss how these aspects may
35 contribute useful analytical tools, metrics, and frameworks that help integrate the uncer-
36 tainties in future soil states, ranging from micro- to regional scales, including those indirectly
37 affected by human activities and management decisions. Overall, this study highlights the
38 potential benefits of incorporating spatial heterogeneity and dynamic instabilities into future
39 model representations of whole soil processes. Additionally, it advocates for transdisciplinary
40 collaborations between natural and social scientists, extending research into anthropedology
41 and biogeosociochemistry, to better integrate and understand longer-term anthropogenic
42 drivers of soil processes, potentially from soil structural dynamics to microbial community
43 and food web ecology.

44 Introduction

45 Soils represent the basis of recurring civilization (*Montgomery, 2007; Marris, 2022*), and
46 models depicting their structure and dynamics may help improve generalized understand-
47 ing of their behavior and ecology. New generation models of soil nutrient cycles (*Sulman*
48 *et al., 2018; Zhang et al., 2021*) certainly improve on older generation ones (*Coleman and*
49 *Jenkinson, 1996; Powlson et al., 1996; Berardi et al., 2020*), adding foci on organic matter
50 stabilization by mineral association that considers the efficiency of microbial degradation
51 (*Cotrufo et al., 2013*), molecular and pool stochasticity (*Sierra et al., 2018; Waring et al.,*
52 *2020; Azizi-Rad et al., 2021*), and increasingly large datasets (*Todd-Brown et al., 2022*).
53 However, both handling large parameter spaces and making predictions across microbial to
54 global scales remain difficult (*Wieder et al., 2015; Vereecken et al., 2016*), explaining com-
55 mon decisions to make design tradeoffs between model generality, or qualitative diversity of
56 dynamical output and potential case applications, with numerical precision and realism at a
57 particular scale (e.g. global) (*Levins, 1966; Livingston, 1985*). Additionally, the increasing
58 recognition of soil habitat structure for organic matter storage and stability (*Cotrufo et al.,*
59 *2013; Kravchenko et al., 2019a; King, 2020*), as well as for community assembly and biodi-
60 versity maintenance (*Erktan et al., 2017; Charlotte et al., 2022; Schweizer, 2022; Vogel et*
61 *al., 2022*), highlights the potential utility of re-conceptualizing how the soil environment is
62 modeled and formulated. Given the breadth of services that soils offer, adopting modeling
63 strategies that are transferable across soil ecology sub-disciplines (*Buchkowski et al., 2017*)
64 indeed helps move toward addressing the most general of goals in soil ecology, such as how
65 soil biotic and abiotic spheres or networks interact over time to confer bulk soil properties.
66 Fortunately, an old yet increasingly studied field of complex systems focuses on gaining
67 generalized insights from large multi-component systems, from social and neural networks
68 (*Marder and Calabrese, 1996*) to oscillating chemical reactions (*Epstein et al., 1983; Pacault*
69 *et al., 1987; Epstein and Showalter, 1996*), that offer potentially useful analytical perspec-
70 tives and strategies for generalizing about how heterogeneous soil environments can behave.
71 This synthesis builds on previous efforts to highlight how soil ecology may benefit from inte-
72 grating tools and concepts from complex systems (*Baveye et al., 2000; Young and Crawford,*
73 *2004; Lavelle et al., 2016; Pachepsky and Hill, 2017; Bennett et al., 2019*), elaborating on
74 how specific modeling principles may help gain insight into a variety of soil processes in more
75 generalized ways.

76 Systems perspective

77 Early studies of complex systems tended to include explicitly recognized spatial extent and/or
78 separation among modules (*Turing, 1952; Levins, 1969; Mandelbrot, 1983*), and others recog-
79 nized explicit state-dependence or time lags (*Rutherford and Do, 1997*). Mathematical roots
80 of systems and complexity theory built upon nonlinear dynamics (*Lorenz, 1963; Winfree
81 and Strogatz, 1984*) and chaos theory (*May, 1974; Li and Yorke, 1975; Rogers et al., 2022*),
82 which was facilitated by computational technologies allowing simulation. In soil science,
83 early reports on fractal dimensions of particle size distributions also represent narratives
84 working toward scale-invariant descriptions of the physical habitat (*Young and Crawford,
85 1991; Kravchenko et al., 1999*), though fractal methods can also be useful in describing
86 non-physical domains. Given how large and complicated some systems can be, analytical
87 research questions have tended to switch focus from valuing the precision of model predictive
88 ability to its generality (*Levins, 1966*), or how consistent a result is across example systems.
89 This strategy of finding intersecting truths across case studies is already somewhat familiar
90 to research in the form of useful reviews, meta-analyses, and model comparison projects
91 (*Sulman et al., 2018*), but it can also take a more computational form. More useful analysis
92 metrics for complex models with variable outputs tend to shift from distribution centrality
93 (e.g. median, mean) to variance and/or key single exponents in cases of high skewness. Ac-
94 cordingly, research questions about similar complex systems also benefit in switching from
95 precise future values to the probability of certain types of events occurring at any future
96 time or location, depending on input data. Focusing on variance as output also promotes
97 including realistic inherent variation or randomness as input, known as stochasticity, which
98 has been a necessary part of explaining and reproducing natural time series of experimental
99 populations (*Henson et al., 2001*). Embracing and allowing for variability in time series
100 also offers the potential to predict and observe temporal autocorrelation in anticipation of
101 critical transitions of tipping points (*Scheffer et al., 2012*), such as during accelerated soil
102 degradation or restoration, or naturally unpredictable chaotic fluctuations with predictable
103 bounds (*Schaffer and Kot, 1985; Hastings et al., 1993*), such as in response to human drivers
104 (*Berryman and Millstein, 1989*).

105 Key examples across fields of study have become increasingly emblematic of complex sys-
106 tems theory. For example, in geophysics, measuring irregular structures in nature using
107 relative units instead of absolute ones (*Mandelbrot, 1983*) has widely uncovered power law
108 frequencies of system components sub-sets, famously used in geophysics to study regional
109 earthquake magnitudes (*Bak et al., 2002; Christensen et al., 2002*). This tool from the frac-
110 tal geometry branch of mathematics was later applied to ecology, both implicitly (*Macarthur*

111 *and Wilson, 1963*) and explicitly (*Harte, 1999; Ostling, 2000*), as well as to soil structure,
112 including aggregation and porosity, with implications for predator-prey (i.e. invertebrate-
113 microbe) interactions (*Baveye et al., 2000*). Although current applications of nonlinear
114 dynamical systems principles to soils remain somewhat siloed in geophysics, potential appli-
115 cations to soil ecology remain. Another exemplary analysis tool is re-framing focal metrics
116 from distribution centrality to variance measures, as mentioned above. This shift has ad-
117 vanced understanding of consistent spatial and temporal fluctuations (*Taylor, 1961*) to focus
118 on sub-module synchronization potential, notably formulated as coupled pendulum dynam-
119 ics in physics (*Kuramoto, 1984*), and has recently inspired various applications in ecology,
120 including to predator-prey dynamics on interaction networks (*Vandermeer, 2021*), dispersal
121 shifts across urban agricultural landscapes (*Ong et al., 2020*), and to biodiversity collapse
122 broadly. These analytical methods tend to uncover internal consistencies or rules among
123 heterogeneous components that predict relatively new patterns at the whole system level,
124 advancing the specific goal of using theory to advance hypothesis testing, compared to other
125 assumed goals like precision forecasting.

126 **Early complexity in soils**

127 In ecology generally, there appears to be a consensus that micro- and meso-habitats are com-
128 plex, and indeed natural soil systems are regarded as the most complex habitats we know
129 of. Yet formal modeling frameworks remain relatively simplified, in part due to common
130 statistical limitations and disciplinary influences. For example, Michaelis-Menten kinetics
131 generated a strong biochemical influence on representing soil organic carbon cycling and de-
132 composition, which has over time been modified to be “reversed” to focus on total enzyme
133 production rather than individual catalysis (*Schimel and Bennett, 2004*), and has maintained
134 wide use as a core model structure for fine-scale carbon and nutrient transformations. This
135 was even applied to population ecology (*Volterra, 1928*), where population growth equations
136 analogous to enzyme kinetics have been well analyzed to yield technically unstable but no-
137 ticeably consistent oscillatory dynamics, also observable in soil data empirically (*Reijneveld,*
138 *2013; Kuzyakov and Zamanian, 2019*) and modeled (*Wang et al., 2014*). Other modeling
139 approaches can also yield counter-intuitive results, which may contradict existing empiri-
140 cal data, or pre-date supporting data from future long-term studies. One case of this was
141 the incorporation of adaptive or environmentally-responsive (i.e. non-linear) enzyme produc-
142 tion, also interpretable as emphasizing a positive biological feedback based on phenotypic
143 plasticity, and formulated as a flux- or interaction modification (*Ludington, 2022*). This
144 early modeling study (*Schimel, 2003*) predicted, in part contrasting contemporary evidence

145 supporting decomposition as primarily nitrogen-limited while aligning with others , that ni-
146 trogen additions to soil would suppress rather than stimulate decomposition, a result that
147 preceded future validating evidence of this from longer-term nitrogen fertilization studies . In
148 this case, a key multiplicative (i.e. non-linear) model term associating organic matter decay
149 with both its current pool size (first-order decay) and also a dynamically-responsive enzyme
150 pool, ultimately improved model generality, specifically across time scales from minutes
151 when proteins turnover to multiple years and seasons. More broadly, mathematical ecol-
152 ogy has provided very interesting insights based on linear stability analysis tools to focus
153 analyses (*King and Schaffer, 1999*), but recent pursuits have expanded to include what was
154 otherwise considered unstable outcomes, like transient dynamics (*Hastings, 2004; Hastings*
155 *et al., 2018*) with implications for introduced species establishment (*Armstrong and McGe-*
156 *hee, 1976; McGehee and Armstrong, 1977; Wilson and Abrams, 2005; Xiao and Fussmann,*
157 *2013*). These approaches indeed provided new insight on consistently observable patterns,
158 based on the increasing use and acceptance of graphical analytical methods like cobweb-
159 bing following the rise of personal computing and simulation power. Recently, comparable
160 tools and concepts from community ecology have been synthetically presented to potentially
161 help address questions in soil carbon cycling and soil community ecology (*Buchkowski et al.,*
162 *2017*).

163 Accordingly, this perspective now draws from theoretical ecology to further extend recent
164 efforts to integrate complex systems principles into soil sciences (*Baveye et al., 2000; Young*
165 *and Crawford, 2004; Sierra and Müller, 2015; Lavelle et al., 2016; Buchkowski et al., 2017;*
166 *Pachepsky and Hill, 2017; Bennett et al., 2019*), and thereby generalize our understanding
167 of the soil habitat, from micro- to macro- scales, notably embracing persistent variation and
168 heterogeneity as notably critical to fundamental soil processes (*O’Leary et al., 2018*). Specif-
169 ically, this paper includes non-exhaustive discussion of various aspects of soil research that
170 may benefit from integrating perspectives from other complex systems research: including
171 memory, oscillations, tipping points and hysteresis. Relevant supporting phenomena ob-
172 served in soils include reversible organo-mineral associations, aggregate- and pore hierarchy
173 and dynamics, persistent wet-dry cycles, multi-year gas fluxes (e.g. respiration) and nutrient
174 availability, higher-order microbial community and predator-prey interactions, and cumula-
175 tive legacy land use history. Overall, adopting tools from systems and complexity theory
176 primarily offers ways of reducing model structural uncertainties (*Bradford et al., 2016*),
177 thereby also potentially facilitating model-data integration efforts, for example by reducing
178 parameter space (*Bennett et al., 2019*).

Memory

Iteration and hierarchy

The concept of memory represents a form of self-referential dynamics specifically based on an event occurring at least one time step in the past, also referred to as time lags, which in some real systems can result in hierarchical structures. An example from the field of physics is nuclear fusion, in which heavier atoms with more protons are made not by the instant fusion of many lighter atoms, but the simpler merging of few medium atoms previously built, interestingly, also with an apparent critical drop in stability at higher atomic numbers after iron (Pfützner *et al.*, 2012). This shows not only a hierarchical building process resulting from the dependence of the existence of larger nuclei on past states (i.e. previously formed smaller nuclei), but this also shows related consequences, namely a skewed distribution of nuclear stability across the spectrum of existing nucleus sizes, where the nuclei of smaller atoms are more tightly bound together than larger ones.

Highly skewed or long-tailed distributions, like the power law family, based on component or module attributes like cluster size, are increasingly recognized as (Gillespie, 2015) properties of complex systems (Clauset *et al.*, 2009; Locey and Lennon, 2016). This contrasts the normal “bell curve” distribution, which is often assumed for classic statistical procedures, from simple averages to within-group variance during linear regression. As a result, observing power laws in data has served as initial support for hypothesizing hierarchical or self-organizing processes, offering a novel path to infer process from pattern. Although various underlying processes can yield similar distributions (Bashkirov and Vityazev, 2000; Curado *et al.*, 2018), observing non-normal distributions may at least help de-emphasize selection-like mechanisms that tend to produce bell curves.

Somewhat surprisingly, the nuclear fusion example parallels soil aggregation almost exactly, where larger aggregates tend to be weaker and less structurally stable than smaller aggregates. In this case soil aggregation is in part affected by the amount of surface area allotted to binding agents, as well as the strengths of the various binding agents themselves, ranging from fine-scale organo-mineral associations to fungal mycelial networks (Tisdall and Oades, 1982). Ultimately, this can be tied back to early discussions of fractal dimension parameters of power law distributions observed in porous media and landscapes (Burrough, 1981), which simultaneously converged with empirical support for soil aggregation as a hierarchical processes (Tisdall and Oades, 1982), and later exploration (Nortcliff, 1984; Armstrong, 1986; Tyler and Wheatcraft, 1989; Perrier *et al.*, 1996; Assouline *et al.*, 1998). As a result, while perhaps controversial (Baveye *et al.*, 2000; Pachepsky and Hill, 2017), one approach

213 for theoreticians and modelers may be to further attempt the study of simple models (*Vi-*
214 *tousek et al., 2022*) to explore ideas and test underlying hypotheses about the implications of
215 hierarchical structural dynamics for soil properties (*Stamati et al., 2013*). This may be espe-
216 cially important for processes that remain near impossible to measure empirically (*Bennett*
217 *et al., 2019*). Some studies seem to align with this process-based modeling focus (*Waring et*
218 *al., 2020*), while others tend to prioritize more output precision after long-term simulation
219 (*Coleman and Jenkinson, 1996; Powlson et al., 1996; Cong et al., 2014*). Additionally, there
220 is also supporting evidence for various hierarchical processes in soils, including those under-
221 lying clay flocculation (*Brostow et al., 2007; Cuthbertson et al., 2018*), aggregation of solids
222 (*Tisdall and Oades, 1982*) (mentioned above) as a result of physical mixing (*Klaminder et*
223 *al., 2013*) and fungal enmeshment (*Rillig and Mummey, 2006*) amplified by wet-dry cycles
224 (*Denef et al., 2001*), and for pore cluster networks (*Quigley and Kravchenko, 2022; Vogel*
225 *et al., 2022*), all of which represent interesting modules of the soil environment to explore
226 with hypotheses about hierarchical structural dynamics and their implications. Specifically,
227 hierarchical aggregation has been previously analyzed by binning all aggregates into either
228 micro- or macroaggregate functional groups, a fractionation that has served as a basic struc-
229 ture for some models specific to solid aggregation (*Segoli et al., 2013; Stamati et al., 2013*),
230 though it may also be relevant for other soil models, such as ones specific to greenhouse gas
231 production (*Kravchenko et al., 2019b; Wang et al., 2019*) or microbial diversity.

232 Overall, these hierarchical patterns emerging from system-wide memory (i.e. past state-
233 dependence) tend to generally reflect natural solutions for coarser level system limitations.
234 Endogenous solutions to system limits can in some cases be formulated and analyzed as an
235 issue of optimizing information flow (*Czaplicka et al., 2013*). If applied to modeling soil pore
236 dynamics, for example, this general mechanism or principle could help explain empirical soil
237 pore structure data specifically as a single naturally-resulting geometric solution (stable or
238 temporary/unstable) for a set of constraints. Biological constraints might be microbial cross-
239 feeding rates fueled by metabolite-carrying water flowing through pore networks. Physical
240 constraints could be spatial, such as topsoil depth, and/or temporal, such as growing season
241 length, both of which would affect pore-forming processes including microbial secretions and
242 gas exchange (micro-pores) and invertebrate burrowing activity (macro-pores). Importantly,
243 real soils also have historical constraints, which state where pores have been and intuitively
244 predict where new ones can potentially form, which ultimately holistically frames pore net-
245 work structures as ongoing solutions, with varying robustness, for both past and present
246 ecological activity. Similarly, hierarchical aggregation of solids may also elongate the effects
247 of environmental changes enough to minimize compositional variance of microbial commu-
248 nities living in the smallest habitat pockets (*Rillig et al., 2017; Wilpieszski et al., 2019*).

249 System memory, in the form of iterative structural dynamics, can also be environmentally
250 adaptive, such as when system components are dynamic or continuously dismantled and
251 re-assembled into new yet familiar structures, which range in scope from soil micro-habitats
252 to landscape profiles over months and years. Fine-scale positive feedbacks may help explain
253 the diversity of soil profiles and pedons across and within order-level taxonomic soil classi-
254 fications (*Jenny, 1961; Phillips, 2017*). In addition to hierarchies, soil systems' internal or
255 temporal memory can also be a regular source of heterogeneity that increases uncertainty
256 about how soil behaves. In soils, this hypothesis of increasing heterogeneity along hierarchical
257 trajectories could, for example, predict greater variance in diversity or simply beta diver-
258 sity among microbial communities observed in macro-aggregates compared to those among
259 micro-aggregates. However, there remains little evidence addressing this topic, although
260 technological advances may facilitate future studies (*Bailey et al., 2012*).

261 **Soil depth and history**

262 Studying subsoils and soil depth also represents an axis that integrates soil profile memory
263 and time lags. Often only topsoils (e.g. to 10 or 20 cm depth) are studied due to high nutrient
264 concentrations there, yet subsoil horizons store more total carbon (*Hicks Pries et al., 2017*)
265 and can influence topsoil microbial activity, ultimately highlighting their relevance to whole
266 profile soil functions. Pedological studies have long recognized that land use history, in addi-
267 tion to the classic five state factors of soils, affects current soil function (*Turley et al., 2020*),
268 especially tillage and fertilization via changes in soil structure and soil fertility (*Weitzman et*
269 *al., 2022*). Additionally, microbes may decompose stable organic matter reserves when new
270 labile organic matter is added, known as soil priming (*Kuzyakov, 2006; Bastida et al., 2019;*
271 *Liu et al., 2020*), to which subsoils may be more sensitive (*Li et al., 2022*), and thus overall
272 priming may offset any expected new carbon storage in topsoils. Similarly, a multi-year
273 whole profile warming experiment recently showed that soil overall lost carbon mostly from
274 subsoil even though topsoil accumulated carbon (*Soong et al., 2021*), suggesting that oppo-
275 site patterns in subsoils may require re-shaping fundamental understanding of soil systems
276 at the profile and pedon levels. In contrast, subsoils at over one meter depth in agricultural
277 systems may be better posed to accumulate carbon in the long-term compared to their top-
278 soils, due to existing degradation from deeper tillage and the potential for added fertility
279 from added root inputs by perennial plants with deep roots (*Button et al., 2022*) and other
280 biological subsoiling methods (*Ning et al., 2022*). Since soil pedogenesis is now understood
281 to follow complex trajectories (*Phillips, 2017*), future study of subsoils should help develop
282 explanations for underlying processes simultaneously affecting different soil horizons.

283 Ideas of memory also relate to group-level patterns and processes which are increasingly
284 reported, alongside individual component-level processes (*Kerr and Godfrey-Smith, 2002;*
285 *Traulsen and Nowak, 2006*), and can have important implications for overall soil processes.
286 Cooperation often manifests itself as synchrony among individuals, as in early examples of
287 tree seed masting (*Ostfeld et al., 2006; Victor et al., 2016*) as well as disease transmission
288 (*Ostfeld et al., 2005*), with similar principles extended to apply to forests (*Filotas et al.,*
289 *2014*) and soil rhizospheres to describe nutrient exchange (*Simard et al., 1997*) along with
290 mycorrhizal symbioses (*van der Heijden and Horton, 2009; Simard et al., 2012*). Similarly,
291 population quorum sensing by soil bacteria has affected antibiotic production (*Li et al., 2021*)
292 and other benefits at critical population sizes (*Heilmann et al., 2015*) and likely with weak
293 time delays in reciprocity (*Alfaro and Sanjuan, 2022*), and quorum sensing has also been
294 hypothesized to affect enzyme production with implications for nitrogen cycling (*DeAngelis*
295 *et al., 2008; Wang et al., 2014*). Time delays also affecting interactions between partners
296 changing at different rates or temporal scales – like between monthly root turnover and
297 exudation affecting momentary microbial gene expression and predatory grazing below- and
298 aboveground – could affect reciprocity of symbioses in soil. Ultimately, the combinations
299 of spatial and temporal variation generate the fluctuations that make treating dynamic
300 instability as a somewhat unavoidable and thus inherent property of natural soil systems.

301 **Oscillations**

302 **Focus on variance**

303 Diel cycles in soil temperature and respiration are commonly observed, yet few studies ana-
304 lyze the implications of natural cycles for modeling soil responses to environmental changes,
305 which in some cases can lead to hysteretic irreversibility (*Phillips et al., 2011*). Complexity
306 and systems theory offer generalized tools and perspectives to better incorporate variance
307 (i.e. by soil depth) into more generalized models of soil processes. The tendency of a systems
308 perspective to shift analysis toward variance is ultimately more inclusive of a diversity of
309 model outcomes, such as regular oscillations or constrained chaotic fluctuations, and more
310 generally validates informative model outputs that are not precise single point solutions
311 and otherwise considered unstable by linear stability analysis. One method of incorporating
312 variances has been to re-formulate dynamical systems using trigonometric functions, which
313 are unique for producing repeated symmetrical curves, from bounded measures of relative
314 distance from a fixed line segment (*Coolidge, 1952*). These first principles already reflect

315 modern principles now understood about many complex systems, like relativity and sym-
316 metry with modification, which likely increase the generality and applicability of modeled
317 output especially when processes and questions are newly framed and formulated in tractable
318 ways.

319 **Soil cycles**

320 Oscillations have been predicted by novel and widely cited model structures in agroecology
321 (*Vandermeer and Perfecto, 2017*) and soil ecology (*Baveye et al., 2018*), and supported by
322 empirical data across fine- and coarse scales. At fine scales, soil carbon molecular turnover
323 has been recently proposed to depend on functional group complexity that also hypothesizes
324 spatial modularity or hotspots in activity, which could produce oscillations of broader soil
325 properties over time (*Lehmann et al., 2020*). Spatial structuring also strengthens positive
326 feedbacks in decomposition between exo-enzyme activity and assembly or production, which
327 has influenced some early soil models toward predicting consistent variance, in the form
328 of limit cycles, in soil nutrient availability (*Schimel, 2003; Wang et al., 2014*). Wet-dry
329 cycle frequency has also been cited as affecting soil aggregate stability and as a determinant
330 of patterns in microbial activity (*Evans et al., 2022*). Various types of limit cycles have
331 also been a classic prediction of predator-prey models, but are rarely applied to describe
332 soil faunal grazer food webs (*Baveye et al., 2000; Buchkowski et al., 2017; Erktan et al.,*
333 *2020; McCary et al., 2021*) especially involving viruses and their traits (*Emerson et al.,*
334 *2018; Trubl et al., 2018*). When these low-dimensional ecological models are explicitly
335 extended in space, for example as metacommunities, a wide variety of mosaic landscapes
336 can be generated with some sensitivity to model formulation or structure (*Vandermeer and*
337 *Yitbarek, 2012; Vandermeer, 2013; Yitbarek and Vandermeer, 2017*), suggesting that flexible
338 model structures are likely important for robust understanding of causes and consequences of
339 soil heterogeneity. Spatial explicitness also emphasizes studying dispersal processes among
340 soil modular populations more than internal community dynamics shaped by particular
341 parameterizations, which have been increasingly studied (*Chaudhary et al., 2020; Hajian-*
342 *Forooshani and Vandermeer, 2020*). Spatial or temporal separation of populations has also
343 inspired other model structures proposing competitive hierarchies resulting in intransitive
344 loops (*Vandermeer, 2013; Vandermeer and Jackson, 2018*), which is already supported
345 by strong evidence (*Kerr et al., 2002; Lozano et al., 2019*). Again, this is a case where
346 new and diverse model predictions precede empirical supporting evidence, which ultimately
347 highlights the potential value of general models, with a wide range of output, over precise
348 ones, specifically for more basic than applied research. At coarser scales, soil respiration has

349 also experienced regular variance in magnitude over multiple years, explained by regime shifts
350 among various dominant stages of community-level decomposer activity (*Sihi et al., 2016;*
351 *Melillo et al., 2017*). Various soil nutrients have also shown regular oscillations over decades
352 (*Reijneveld, 2013*), although common statistical analyses remain linear which remains an
353 analytical limitation. Some nutrient oscillations are expected from direct harvests, but if
354 nutrient cycles are coupled (i.e. by microbial metabolisms) while oscillating, unintended and
355 unintuitive synergies or conflicts may emerge, as oscillations align either in- or out of phase
356 (*Vandermeer, 2006*), leading to either robust, or more likely quickly degrading, soil nutrient
357 availabilities. The appearance and prediction of oscillations among a range of soil variables
358 from both advanced first principles and several empirical studies justifies further study of
359 non-linear models of soil behavior and ecology, with example approaches listed in Table 1.

360 **Tipping points and hysteresis**

361 Hysteresis can be framed as a specific kind of non-linear transition and tipping point occur-
362 ring in soils. Hysteresis is revealed when model stability analyses include outcomes across
363 range of parameters, and underlying component processes also change pace at different speeds
364 (*Ong and Vandermeer, 2018*). In soils, this occurs with overall water content and its matric
365 pressure, or availability; this is because water always moves through wide pores first regard-
366 less of whether the process is drying or wetting, and wide pores have a weaker relationship
367 between these two water variables than in narrow pores, which always mediate water move-
368 ment second. Here the phenomenon is explained by temporal consistency in process across
369 spatial heterogeneity, which can be another cause of dynamical asymmetry more generally.
370 For example, the relationship between soil temperature and total soil respiration is also
371 hysteretic in that soil respiration responds only after considerable changes in temperature,
372 making respiration stay low as soil warms yet stay high while it cools, even across the same
373 middle temperatures (*Riveros-Iregui et al., 2007; Phillips et al., 2011; Zhang et al., 2015*),
374 which can also be described as state-dependence or short-term memory for a system. A
375 widely recognized implication of hysteresis is the practical and sometimes permanent (*Ong*
376 *and Vandermeer, 2018*) irreversibility of past ecological states, like permanent wilting points
377 for roots under drought, or more internally regulated predator or pest populations.

378 Soil carbon saturation is a related concept, which is supported by field experiments and
379 models, and is inspiring useful model structures for describing and understanding non-linear
380 processes in soil. Appropriately, the concept of soil carbon saturation has already been
381 formulated and incorporated into new generation soil models like MEMS (*Zhang et al., 2021*).

Table 1: Example cases where employing modeling strategies based on systems theory could be useful for generalizing about soil ecology and agroecology.

| Goal | Limitation | Systems concept | Model approach | Pro | Con | Related Refs |
|---|--|---|--|---|--|---|
| Forecast soil C for general agricultural or climate change management decisions | Reversible organo-mineral associations (priming); Enzyme production plasticity | Oscillatory dynamics; strong positive feedbacks | Simpler implicitly oscillatory pools, with different winding frequencies | Fewer pool-specific parameters; Computational efficiency allows more stochastic simulations for system accuracy | Less precise final SOM pool sizes | Schimmel & Weinttraum 2003; Wang et al 2014 Biogeosci; Sulman et al 2018; Kuramoto 1975 |
| Estimate soil biodiversity for conservation and sustainable development | Large nested food webs, with trophic cascades | Indirect / higher-order interactions | Collection of coupled oscillator predator-prey pairs and/or synchronous community clusters | Fewer taxon-specific parameters; Computational efficiency allows more taxon pairs | Less precise species population sizes | Potapov 2022; Buchkowski et al 2017; Vandermeer et al 2021 |
| Assess persistent pathogen risk from soil, compost, or municipal sludge | High heterogeneity in microbial pathogen population sizes | Chaos, bifurcations, period-doubling | Meta-population with chaotic (high) discrete growth and/or dispersal rates | Higher statistical confidence compared to plate colony count estimates | Output states more parameter dependent | Levins 1969; May 1974; Feigenbaum 1978 |

382 An early proposed model (*Stewart et al., 2007*) simulated decadal field data by mixing two
383 pools with different local equilibria: a stable mineral-bound carbon pool that saturates,
384 and a labile pool that mostly decomposes but with no growth limit. This mixed structure
385 produced flexible output – slow linear increases in soil carbon at high concentrations, and
386 quicker but saturating increases at low concentrations – making it more widely applicable
387 and representing a more generalized understanding of long-term soil carbon dynamics.

388 In contrast, older-generation soil models like CENTURY (*Powlson et al., 1996*) and RothC
389 (*Coleman and Jenkinson, 1996*) did not incorporate pool saturation, but instead offered a
390 simpler initial models offering widespread practical use. The relative simplicity of the model
391 structure is broadly observed in the overall linear successive flow from one decomposition
392 pool to the next. Furthermore, each individual pool was formulated to observe first-order
393 decomposition kinetics. While first-order decomposition represents exponential decay, which
394 is curvi-linear over time, the non-linearities and complex dynamics under investigation by
395 systems theory instead stem from multiplicative associations or interactions among dynamic
396 variables. In many cases, multi-pool soil systems tend to be represented with an attempted
397 thoroughness, resulting in a list of coupled equations that influence each other by additive
398 (or subtraction) terms, which usually consist of a key variable multiplied by a corresponding
399 abiotic parameter, which is static relative to carbon or the nutrient of interest (*Zhang et al.,*
400 *2021*). This style of formulation acknowledges that soils are complicated habitats, but leaves
401 room to lean into the complexities of the habitat with various nonlinear additions.

402 Fundamentally, because parameters and dynamic variables in an ordinary differential equa-
403 tion often represent concepts formulated to be operating at relatively distinct time scales,
404 the re-formulation to incorporate multiplicative variable associations can also be interpreted
405 as a change in descriptive timescale. For example, a representative modeling study may test
406 the hypothesis that particulate organic matter and mineral-associated organic matter, even
407 as distinct pools, might each observe transfer rates that could be affected by both its own
408 and the other’s size, within a model time step, rather than between them. More specifically,
409 rather than modeling the primed loss of existing organic matter implicitly in the form of a
410 saturating stable pool, soil priming could be formulated as an interaction modification that
411 implicates the more stable pool’s size back into its incoming transfer rate from the labile
412 pool. Generally, this style of formulation may offer a new class of hypothesis testing, espe-
413 cially for soil processes that are currently difficult to test empirically (*Bennett et al., 2019*),
414 representing a wider array of model outcomes with just a few key soil descriptor variables,
415 in line with recent qualitative syntheses (*Phillips, 2017; Kuzyakov and Zamanian, 2019*).
416 Modeling soil systems using tools from complexity theory appears to offer an antidote to the

417 increasingly-large soil simulations that have become more popular, in part alongside increas-
418 ing computing power in cloud systems. The approach using complexity theory can improve
419 conceptual efficiency by reformulating soil models into fewer modular components with more
420 inter-linked process rates (*Lehmann et al., 2020*). In this way, soil modeling studies may
421 emerge as useful analyses not only for understanding soils themselves, but also for aiding
422 other academic disciplines studying complex systems more generally.

423 Additional recent studies improved generality of understanding by explicitly modeling bio-
424 logical (*Craig et al., 2021*) and microbial (*Wang et al., 2014*) processes underlying trans-
425 formation rates, including with saturating enzyme activity (*Buchkowski et al., 2017*). Even
426 more general understanding can be added by expanding similar model structures to describe
427 effects of soil fauna and invertebrates on soil micro-habitats including predatory and dis-
428 persive influences on microbial communities (*Grandy et al., 2016; Creamer et al., 2022*).
429 However, integrating models of short-term microbial processes with long-term carbon dy-
430 namics remains incomplete (*Todd-Brown et al., 2013*), in part because they operate on very
431 different timescales. Each model class improves understanding of soil, but for generalized
432 predictions, integrating or coupling models with the fewest explicit dimensions that yield
433 the most diverse qualitative output (*Levins, 1966; Lane, 2018*) may help at least bound
434 possible outcomes, such as for soil carbon exchange and net storage, or other nutrient cycles
435 (*Manzoni and Porporato, 2009*). In addition there are many aspects of soil ecology and func-
436 tioning such as biodiversity maintenance or specific soil pathogen suppression that warrant
437 continued understanding through modeling the soil environment.

438 **Restorative agropedogenesis**

439 Together, several non-linear dynamics or functions among soil variables may then combine
440 to reveal critical transitions (Fig 1). A recent synthesis (*Kuzyakov and Zamanian, 2019*)
441 revealed many non-linear relationships among key soil variables across several decades, such
442 as between soil bulk density and organic carbon. Their analysis showed distinct phases
443 of pedogenesis, or soil development, under human management, which they termed agro-
444 pedogenesis. Ultimately, they presented a convergence of these processes into an attractor,
445 which they considered to span a narrower range of values for key soil descriptors than would
446 be found under natural pedogenesis. However, underlying non-linearities between the same
447 key soil descriptors, such as bulk density and soil organic carbon, suggest potential future
448 attractor instability and more potential divergence among local pedons, especially under
449 human management. Relatedly, another recent review (*Phillips, 2017*) highlighted how nat-
450 ural pedogenesis can be described with concepts from non-linear dynamics and complexity

451 theory, notably revealing how pedogenesis can show unstable trajectories, or many possible
452 alternative end states. These studies suggest that while globally, recent anthropedogene-
453 sis has converged mostly toward degradation, locally, future anthropedogenesis could still
454 diverge depending on regional management strategies. This inference that helps maintain
455 agro-pedogenesis as a potentially regenerative force for soil fertility, rather than inherently
456 degrading, especially when distinguishing by industrial vs. small-scale agricultural land man-
457 agement. Accordingly, a recent global synthesis of soil microbial biomass carbon showed
458 evidence of higher variance in percent change where existing stocks were lower, also show-
459 ing relative increases in some tropical regions such as the Caribbean (near southeast Ayiti)
460 (*Patoine et al., 2022*). In this region and others including southern Africa and central Asia,
461 even where soil microbial biomass carbon was lost in sum due to climate changes, land-use
462 change effects on soil microbial biomass carbon were often much closer to positive. This
463 higher variance in soil microbial biomass carbon, in part driven by positive land-use effects,
464 together with global cropland analyses (*Padarian et al., 2022*), points to a potential for
465 land management to increase soil carbon, rather than necessarily degrade it (*Dynarski et al.,*
466 *2020*). These insights highlight that increasing soil carbon globally may still be still feasible
467 especially by sustainable local or regional management coordination.

468 Accordingly, even social processes that affect regional land management, such as in agricul-
469 ture, represent a source of uncertainty affecting soil organic matter dynamics. However, for
470 soil ecology, the understanding offered by systems and complexity theory comes less from
471 understanding farmers' individual decision-making processes, which is still relevant (*Mestre*
472 *et al., 2020*), but instead comes from how information spread and collective decision-making
473 processes may affect regional agricultural management regimes and related soil processes.
474 Given the modern corporate food regime (*Campbell, 2009; McMichael, 2009*) and reliance
475 on plantation land systems (*Wolford, 2020*), widespread deficit narratives about the ecology
476 of agroecosystems have been somewhat normalized. However, a liberal interpretation of a
477 systems approach to soil ecology, by virtue of normalizing and aiming to integrate hard-
478 to-anticipate critical transitions (*Scheffer et al., 2012; Scheffer et al., 2015*), motivates the
479 vision for a counter-narrative – one that focuses on supporting the world's small-scale farm-
480 ers. This results from future global changes to, for example, soil carbon, emerging from the
481 sum of local and regional cooperative efforts toward ecosystem restoration and regenerative
482 agriculture .

483 Rather than continuing to degrade, regional soil organic matter levels could potentially in-
484 crease rapidly, assuming most soils are far from carbon (or other nutrient) saturation. This
485 could be directly supported by local social movements that have formed to protect small-scale

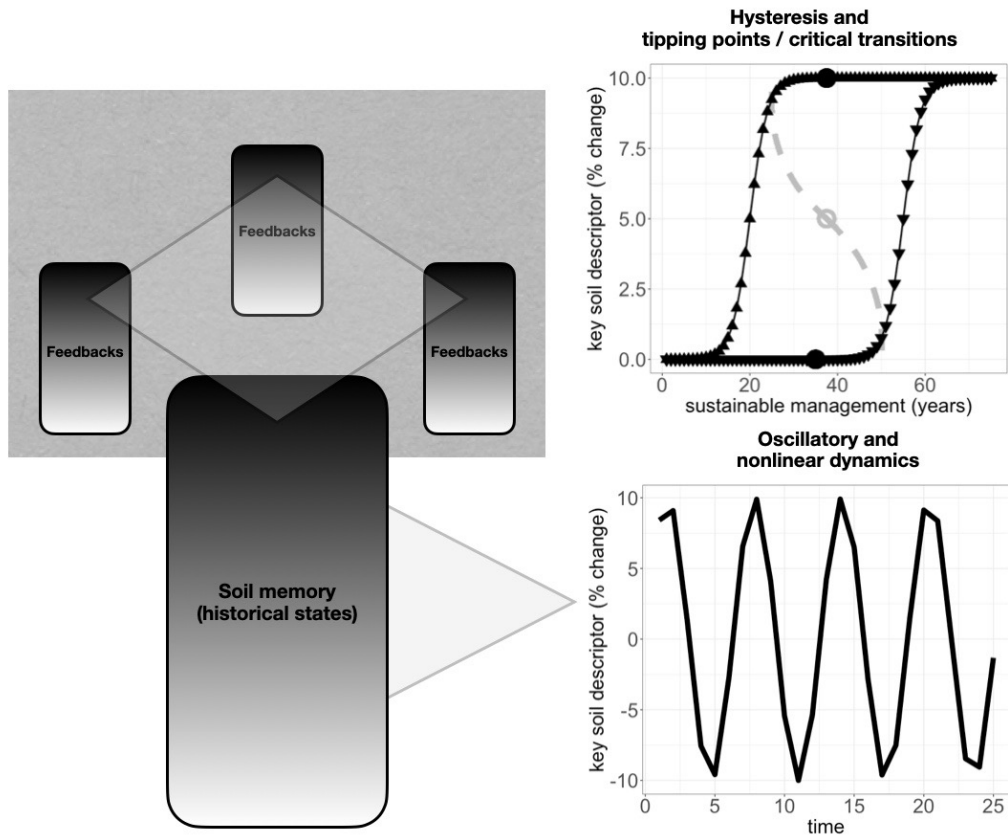


Figure 1: Conceptual diagram highlighting (bottom) that past states incorporated down a soil profile, like buried or subsoil horizons that are plowed, primed and nutrient-poor, or instead deeply fertile (black gradient, bottom-left), may induce oscillatory dynamics (bottom-right) in key soil variables, like moisture and microbial biomass, and that (top) similarly continuous feedbacks across many soil profiles and pedons in a landscape (top-left) may produce nonlinear regional patterns including tipping points (top-right) for key soil variables, like soil organic matter, not only during land degradation, but also during coordinated collective management toward regeneration.

486 agriculture and land tenure via local neighboring cooperation and collaborative governance.
487 Studies in Brazil near activity of the landless workers movement, or MST, show that soil
488 fertility is positively associated with the recent adoption of and transition to sustainable land
489 management practices (*Stratton et al., 2022*) like cover crop diversity, which could also be
490 promoted by stronger social recognition of benefits (*Williams et al., 2021*). This perspective,
491 together with other studies directly tying soil organic matter to crop yield (*Oldfield et al.,*
492 *2022*), ultimately tie soil regeneration to environmental justice issues of land tenure. Inte-
493 grating socio-ecological processes thus offers an additional means by which key soil variables
494 and anthropogenesis can show critical non-linear transitions, especially in positive direc-
495 tions, challenging common narratives that depict humans as forces of soil degradation. Some
496 ideas of how social processes affect nutrient cycling have emerged as biogeo-socio-chemistry
497 especially for urban settings (*Pataki et al., 2011; Kaushal et al., 2014*), and others may help
498 address additional soil ecological dimensions of multi-functionality (*Creamer et al., 2022*)
499 beyond nutrients like spatial patterns of faunal diversity and soil food web network struc-
500 tures. Recent studies of urban ecology already point to interesting patterns that challenge
501 deficit narratives of societal relationships with local soils and agriculture (*Bonilla-Bedoya et*
502 *al., 2022; Nugent and Allison, 2022; Pindral et al., 2022; Zhang et al., 2022*). Novel insights
503 on soil socio-ecological dynamics may help guide how to tailor sustainable development ini-
504 tiatives by individual countries to achieve international soil governance initiatives (*Farnese,*
505 *2022; García et al., 2022*) like through the UN FAO Global Soil Partnership, Global Soil
506 Biodiversity Initiative (*Wall et al., 2015*), and other working groups generally addressing
507 UN sustainable development goals of combating soil and habitat degradation to enhance
508 ecosystem services via dynamic key soil ecological indicators (*Pradhan et al., 2017; Bennich*
509 *et al., 2020*).

510 Conclusion

511 This synthesis applies a complex systems framework to analyzing key uncertainties about soil
512 processes and habitats, drawing from various analytical tools used across interdisciplinary
513 fields, and presenting how they have and will better address key research questions in the field
514 of soil agroecology. Concepts highlighted include soil memory, or legacy effects of manage-
515 ment history and past ecological states; oscillations over time, which are observed in many
516 key soil descriptive variables; and, tipping points and hysteresis, when several correlated
517 variables change at different times and/or rates. Together, these three principles should help
518 decrease uncertainty around soil model structures (*Bradford et al., 2016*) by pointing toward

519 how to improve model generality for key soil processes of interest, such as soil respiration and
520 particulate organic matter storage via occlusion by microbial community byproducts pro-
521 duced during soil aggregation. The complexity perspective on soil agroecology also inspires
522 a positive outlook on the potential feasibility of collective societal solutions to soil degrada-
523 tion crises (*Montgomery, 2007; Richter, 2021*), by including and better anticipating drastic
524 unexpected changes that often emerge from the combined effects of many interacting pro-
525 cesses. Because farms are social-ecological systems, forward steps include collaborating with
526 social sciences, humanities such as history or literature (*Schloss and Handelsman, 2007*), or
527 non-equilibrium or statistical physics (*Bak et al., 2002*) including geophysics, to expand and
528 improve the set of potentially useful analytical tools and perspectives needed to inclusively
529 and robustly describe the extreme heterogeneity and complexity of soil habitats.

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534 Authors declare no conflicts of interest.

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539 **Author contributions**

540 NM and JV contributed to paper conceptualization, funding acquisition, literature review,
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543 **Data statement**

544 Not applicable; All references to literature data are made in-text.

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