

1 Ecological stoichiometry as a foundation for omics-enabled biogeochemical models of soil
2 organic matter decomposition

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4 Emily B. Graham^{1,2*} (0000-0002-4623-7076) and Kirsten S. Hofmockel^{1,3} (0000-0003-1586-
5 2167)

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7 ¹Earth and Biological Sciences Directorate, Pacific Northwest National Laboratory

8 ²School of Biological Sciences, Washington State University, Richland, WA, USA

9 ³Department of Agronomy, Iowa State University, Ames, IA, USA

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11 *Corresponding author: emily.graham@pnnl.gov

12 **Abstract.**

13 Coupled biogeochemical cycles drive ecosystem ecology by influencing individual-to-
14 community scale behaviors; yet the development of process-based models that accurately capture
15 these dynamics remains elusive. Soil organic matter (SOM) decomposition in particular is
16 influenced by resource stoichiometry that dictates microbial nutrient acquisition ('ecological
17 stoichiometry'). Despite its basis in biogeochemical modeling, ecological stoichiometry is only
18 implicitly considered in many high-resolution microbial investigations and the metabolic models
19 they inform. State-of-science SOM decomposition models in both fields have advanced largely
20 separately, but they agree on a need to move beyond pool-based models. This presents an
21 opportunity and a challenge to maximize the strengths of various models across different scales
22 and environmental contexts. To address this challenge, we contend that ecological stoichiometry
23 provides a framework for merging biogeochemical and microbiological models, as both
24 explicitly consider substrate chemistries that are the basis of ecological stoichiometry as applied
25 to SOM decomposition. We highlight two gaps that limit our understanding of SOM
26 decomposition: (1) understanding how individual microorganisms alter metabolic strategies in
27 response to substrate stoichiometry and (2) translating this knowledge to the scale of
28 biogeochemical models. We suggest iterative information exchange to refine the objectives of
29 high-resolution investigations and to specify limited dynamics for representation in large-scale
30 models, resulting in a new class of omics-enabled biogeochemical models. Assimilating
31 theoretical and modelling frameworks from different scientific domains is the next frontier in
32 SOM decomposition modelling, and advancing technologies in the context of stoichiometric
33 theory provides a consistent framework for interpreting molecular data, and further distilling this
34 information into tractable SOM decomposition models.

- 35 **Key words:** carbon use efficiency, carbon cycling, nitrogen cycling, ecosystem models, soil
- 36 nutrients, microbiome, stoichiometry

37 **Introduction.**

38 The world's soil ecosystems contain a pool of carbon that is larger than vegetation and
39 atmospheric stocks combined, with 1500 Gt C stored in the top one meter alone (Batjes 2016). A
40 large proportion of belowground C is contained in soil organic matter (SOM) whose
41 destabilization has the potential to dramatically alter future climates (Crowther et al. 2019). Soil
42 organic matter decomposition is regulated by physical, chemical, hydrologic, and biological
43 controls (collectively called biophysical controls) on resource availability that determine
44 microbial energy generation and carbon and nutrient acquisition (Robertson et al. 2019; Wang &
45 Houlton 2009; Zhang et al. 2014). These processes are notoriously difficult to measure and
46 predict beyond the scale of experimental plots (Bond-Lamberty et al. 2016; Naylor et al. 2020).
47 Because of this, SOM decomposition at scales most relevant to climate change largely continues
48 to be predicted through generalized environmental proxies such as moisture, temperature,
49 minerology, and total soil C or nutrient pool sizes (Bailey et al. 2018) — parameters that leave a
50 substantial amount of uncertainty surrounding model predictions (Todd-Brown et al. 2014;
51 Todd-Brown et al. 2013). The shortcomings of predicting SOM decomposition beyond localized
52 scales point to a need for scalable spatial and biogeochemical processes in SOM decomposition
53 models (Todd-Brown et al. 2013).

54 This topic has received considerable attention from multiple scientific domains over the
55 past few decades, resulting in parallel advances between biogeochemical and microbial research
56 (Fig. 1) and calls for greater diversity of measurements used within biogeochemical models
57 (Billings et al. 2021). Collectively, we have made outstanding progress in our understanding of
58 global C cycles; however, a key opportunity remains in integrating theoretical and modelling
59 frameworks from different scientific domains. Existing classes of models each have strengths at

60 different scales and environmental contexts — where one model fails, another excels. Recent
61 work has improved conceptual, empirical, and numerical representations of mineral protection
62 and hydrologic mechanisms that regulate microbial access to SOM (Blankinship et al. 2018;
63 Waring et al. 2020; Woolf & Lehmann 2019). In contrast, detailed representations of the
64 microbial metabolisms that directly convert SOM into carbon dioxide are still missing from
65 predictive models at the ecosystem-scale (Stone et al. 2021). Soil organic matter decomposition
66 (particularly unprotected SOM) is often rate-limited by microbial metabolism (Dwivedi et al.
67 2019; Hunter et al. 1998; Lehmann & Kleber 2015). In turn, microbial decomposition is coupled
68 to necromass stabilization that is now thought to be the primary mechanism of belowground C
69 storage (Cotrufo et al. 2013; Kallenbach et al. 2015; Kallenbach et al. 2016; Liang et al. 2019;
70 Robertson et al. 2019; Zhang et al. 2021; Fig. 2). Microbial processes and soil physical
71 properties are therefore central considerations in SOM modelling, but they are still coarsely
72 represented in cutting edge SOM decomposition models. Due to recent technological
73 achievements, we now have the ability to measure soil biophysical properties with high
74 molecular resolution (i.e., genes, transcripts, proteins, and metabolites, and soil mineralogy,
75 porosity, chemistry). Our current challenge lies in interpreting and scaling these vast molecular
76 data types into information useful for biogeochemical models (Hall et al. 2018; Malik et al.
77 2020). Next-generation models that are able to link across molecular to ecosystem scales should
78 enable more robust predictions of SOM decomposition (Chowdhury et al. 2019), with follow-on
79 improvement in global C cycling predictions.

80

81 **Research Opportunities for Integrating Microbial and Biogeochemical Modelling**

82 **Approaches**

83 The importance of various controls on decomposition rates differs with the state of SOM
84 at a given point in time and space, the portion of the soil system that is being represented in a
85 given model, and/or with the predictive goal of the model. Biophysical attributes that regulate
86 SOM decomposition collectively vary with soil type and horizon (Patel et al. 2021; Rumpel &
87 Kögel-Knabner 2011); and SOM pool size and composition is dynamic within a single
88 environment. SOM is continuously added to soils through plant litter production, root turnover,
89 and rhizodeposition, while simultaneously being metabolized through microbial decomposition
90 that leads to secondary metabolites and biomass synthesis, all of which have multiple possible
91 routes through soil C cycles (Cotrufo et al. 2013; Liang et al. 2019; Sulman et al. 2014; Wieder
92 et al. 2014). Small changes in the balance of various SOM inputs and subsequent decomposition
93 processes impact the overall fate of C within the Earth system. Embedded in this is an
94 understanding of rate limitation, as soil biology interacts with physiochemistry to constitute a
95 holistic soil ecosystem. In some cases, SOM decomposition may be regulated by abiotic
96 interactions with soil biology — relationships between SOM substrates and microbial
97 decomposition are perhaps most important in the soil organic horizon and within the rhizosphere
98 in mineral layers, where decomposition ultimately regulates C flow to persistent SOM pools

99 These ideas have been formalized into a biogeochemical model that centers on mineral
100 stabilization while also highlighting the controls of microbial metabolic efficiency on rates of
101 SOM formation and accumulation (Robertson et al. 2019, Fig. 2). While there may be physical
102 limits on the amount of SOM that can accumulate in mineral layers (the C-saturation hypothesis,
103 Six et al. 2002; Stewart et al. 2007), Craig et al. (2021) recently used a simulation model to show
104 that microbial community dynamics which impact biomass C flow are also strong regulators of
105 mineral-stabilized SOM. As such, microorganisms impact SOM cycling across all soil horizons.

106 Understanding the strength and mechanisms of this regulation across different portions of the
107 soil system (e.g., size fractions and horizons) will further help constrain SOM model
108 development but is beyond the scope of this paper.

109 Key foci of biogeochemical modelling efforts to improve estimates of SOM
110 decomposition and/or accumulation include mineral-SOM interactions, soil hydrologic
111 properties, chemical speciation of SOM, and microbial biomass and/or microbial use efficiency
112 (Dungait et al. 2012; Robertson et al. 2019; Wieder et al. 2014; Zhang et al. 2021).
113 Physicochemical protection (e.g., mineral-associated 'MAOM' or 'protected pool(s)' vs.
114 'unprotected pool(s)' including rhizosphere and particulate organic matter) and SOM chemistry in
115 particular are pervasive across most classes of biogeochemical models (Robertson et al. 2019;
116 Sulman et al. 2014; Zhang et al. 2021). For instance, the PROMISE framework centers on SOM
117 accessibility as a key determinant of decomposition, encompassing the known roles of clay
118 mineral composition and heterogeneity in residence times of identical molecules (Müller &
119 Höper 2004; Waring et al. 2020). State-of-science biogeochemical frameworks, even those with
120 an emphasis on physical processes, acknowledge chemical diversity within all SOM pools by
121 using chemical parameters to constrain the fate of various classes SOM and/or generate
122 distribution of decomposition rates for a given pool (Azizi-Rad et al. 2021; Robertson et al.
123 2019; Waring et al. 2020; Zhang et al. 2021).

124 While the rate-limiting biophysical factor(s) fluctuates with environmental context,
125 microbial metabolism is the proximate control of bioavailable SOM decomposition in all cases.
126 Because the connectivity between protected and unprotected SOM is already encompassed by
127 state-of-science biogeochemical models, we largely consider processes occurring within
128 unprotected SOM pools when considering avenues for consolidating microbial and

129 biogeochemical models. While we recognize that pore-scale spatial data is critical to
130 understanding SOM decomposition, these data types remain difficult to obtain and are not yet
131 widely used. In contrast, omics data are becoming widespread, and information on microbial
132 communities reflect their underlying physical and chemical habitat. Process-based models
133 arising from microbiology generally represent much smaller spatial and temporal scales than
134 biogeochemical models, with inputs comprised of higher molecular resolution data. They tend to
135 focus on genomic pathways or microbial traits that lead to decomposition (Borer et al. 2019;
136 Borer & Or 2021; Heinken et al. 2021; Malik et al. 2020; Wang & Allison 2021).

137 Each class of microbial and biogeochemical models provides valuable information on
138 SOM decomposition while also acknowledging considerable error and seeking opportunities for
139 improvement. Therefore, a pressing question remains: how do we merge disparate model
140 structures and underlying theoretical assumptions such that we maximize our predictive power of
141 SOM decomposition, mineralization, and stabilization?

142 We see a ripe opportunity for harmonizing microbial and biogeochemical SOM models
143 within and across all scales through substrate (plant and microbial) chemistries, which determine
144 SOM composition (from a biogeochemical perspective) as well as the rate and pathway of
145 microbial metabolism (from a microbial perspective). Billings et al. (2021) have identified SOM
146 chemistry as a measurement for the next generation of soil C research, highlighting the
147 increasing usage of molecular advances such as Fourier-Transformed Infrared Spectroscopy
148 (FTIR, Cheng et al. 2006; Keiluweit et al. 2010), mid-infrared spectral libraries (Dangal et al.
149 2019), and Diffuse Reflectance Fourier-Transformed Infrared (DRIFTS, Kaiser & Ellerbrock
150 2005; Leue et al. 2010). Substrate chemistry is represented in fine-scale biogeochemical models
151 (e.g., reaction networks and reactive transport models) by influencing reactions with specific

152 SOM molecules, extracellular enzymes, or inorganic nutrients (Porta et al. 2018; Riley et al.
153 2014; Wilson et al. 2019). Microbiology also considers substrate chemistry at the individual
154 level (e.g., single cell metabolic models), albeit with different terminology, in the framework of
155 metabolomics. Metabolic models consider substrate chemistry in the context of a model
156 organism's genome to make predictions of growth dynamics or specific metabolomic products
157 (e.g., GEMS, Duarte et al. 2004; Loira et al. 2012; Lu et al. 2019). Biogeochemical models at
158 intermediate scales consider the composition of elemental and SOM pools, often through
159 relatively coarse parameters such as fast vs. slow cycling SOM pools (unprotected vs. protected),
160 mineral SOM, chemical classes or C:N ratios of SOM, and/or organic vs. inorganic nutrient
161 pools. Many recent advances to consider SOM composition have been incorporated in MEND
162 (Wang et al. 2020a; Wang et al. 2015), RESOM (Tang & Riley 2015), CORPSE (Sulman et al.
163 2014) and MIMICS-CN (Kyker-Snowman et al. 2020), yet none leverage molecular chemical
164 information or stoichiometry of compounds or elements other than C and N. Microbial models at
165 the intermediate-scale simulate interactions between individuals, limited sets of individuals (or as
166 cumulative "mixed-bag" organisms)(Henry et al. 2016), or their characteristic traits to predict
167 metabolite chemistry or changes in the growth of specific organisms or functional guilds
168 (Bouskill et al. 2012; Follows & Dutkiewicz 2011; Malik et al. 2020; Reed et al. 2014). Recent
169 research has also implicitly used the intersection of biogeochemistry and microbiology through
170 substrate chemistry in order to move towards integration by representing fine- (individual
171 microbial metabolisms and metabolites), intermediate- (groups of microorganisms and/or
172 elemental pool composition), and coarse-resolution (multiple interacting microbial groups,
173 "mixed bag" communities, and/or elemental pool sizes) connectivity of microbial drivers and
174 biogeochemical outcomes in predictive models at a particular level of resolution (Fig. 1, 3, 4).

175 Most of the aforementioned biogeochemical models also include some aspects of mineral-OM
176 interactions and/or soil hydrologic properties.

177 Despite their coarse representation of microbial processes, biogeochemical models with
178 explicit microbial dynamics represent an improvement over seminal biogeochemical models
179 (Wieder et al. 2013). This has led to soil microbiology research aimed at improving predictions
180 of SOM decomposition through microbial genomics. For instance, Stone et al. (2021) recently
181 demonstrated that the efficiency of microbial SOM decomposition co-varies with taxonomy and
182 soil nutrient status. Other efforts include the prediction of microbial carbon use efficiency from
183 microbial metagenomes (Saifuddin et al. 2019), microbial function under warming (Guo et al.
184 2020), and SOM decomposition from genome-derived microbial traits (Malik et al. 2020; Wang
185 & Allison 2021) — all of which emerge from decades of work in microbial ecology aiming to
186 understand the role of microbiome composition in SOM decomposition (Graham et al. 2016).
187 State-of-science models in both biogeochemistry and microbiology arise out of a need to move
188 beyond pool-based models towards continuous relationships between microbial metabolism and
189 SOM decomposition and persistence (Tang & Riley 2020; Waring et al. 2020), but they rely on
190 different underlying frameworks such that there are few common concepts, parameters, and
191 equations with which to connect biogeochemical and microbial models.

192 Ecological stoichiometry (i.e., the role of elemental composition in regulating ecological
193 interactions, Sterner & Elser 2002; Van de Waal et al. 2018) is a useful framework for cross-
194 scale and cross-domain integration. While many biogeochemical models are built on the
195 principles of ecological stoichiometry, they are largely absent from microbial models. Ecological
196 stoichiometry uses the elemental (e.g., C:H:N:O:P:S) imbalance between organisms and their
197 substrates to predict that nutrient limitations universally control biological activity (Elser et al.

198 2000; Elser et al. 1996; Sterner 1995). It has been leveraged to decipher overarching patterns in
199 the processes governing SOM decomposition (Billings & Ballantyne IV 2013; Sinsabaugh et al.
200 2013; Sinsabaugh & Shah 2011) and is the guiding framework for decades of biogeochemical
201 investigations that spawned process-based models.

202 Biogeochemical modelling research has begun to explore the usage of ecological
203 stoichiometry across a broad range of scales. At the largest scale (pool models), a recent update
204 to the CENTURY model (Parton 1996) demonstrates increased accuracy from a tighter coupling
205 of C and N dynamics (Berardi et al. 2020). At higher molecular resolution, Buchkowski et al.
206 (2019) have improved predictions of SOM decomposition by more explicitly accounting for the
207 stoichiometries of microbial biomass and SOM. Fatichi et al. (2019) have proposed direct
208 representation of microbial communities involved in coupled SOM and nutrient cycling, an
209 approach that could not only aid in increasing the accuracy of SOM decomposition predictions
210 but also become a useful tool for empiricists to interpret multidimensional microbial data.

211 By using an ecological stoichiometry framework to combine new microbiological and
212 computational tools with biogeochemistry, we have an unexploited opportunity to combine
213 predictive models across measurement scales and to better understand global patterns in
214 microbial SOM decomposition. The use of top-down (ecosystem to molecular) and bottom-up
215 (molecular to ecosystem) information exchange guided by ecological stoichiometry can
216 iteratively (1) refine the objectives of high molecular resolution investigations and (2) specify a
217 limited set of dynamics for representation in large-scale models through omics-enabled reaction
218 networks (Fig. 4). Top-down approaches leverage bulk biogeochemical pool sizes to identify
219 nutrient limitation(s) that then guide parameters selection and calibration of reduced complexity
220 models. Bottom-up approaches use stoichiometric principles to guide the analysis and

221 interpretation of large molecular datasets by establishing predictive rules that are translatable
222 across scales. We propose that coupling these methods can specifically improve representations
223 of SOM decomposition within the soil organic layer that ultimately regulate C transport and
224 stabilization in mineral horizons. By using ecological stoichiometry as a conceptual and
225 modelling framework to transfer knowledge across scales and couple existing models, we have
226 the potential to develop a new generation of reduced complexity omics-enabled models.

227

228 **Ecological Stoichiometry as a Guiding Framework in Soil Organic Matter Decomposition.**

229 Microbial (i.e., fungal, bacterial, archaeal, and viral) degradation of chemically diverse
230 SOM is reliant on broader organismal and soil stoichiometry (Buchkowski et al. 2015;
231 Buchkowski et al. 2019; van Groenigen et al. 2006). Nutrient availabilities in soil determine the
232 metabolic pathways (and their efficiencies) and substrates that are used during SOM
233 decomposition (Ge et al. 2020; Sinsabaugh et al. 2016; Wei et al. 2020). In turn, microbial
234 decomposition and necromass products fuel a cycle of SOM decomposition, formation, and
235 stabilization. Much belowground nitrogen (N) and sulfur (S) is stored within organic matter that
236 is mineralized by soil microbial communities (Freney 1986; Jenkinson 1990; Kirkby et al. 2011).
237 While phosphorus (P) availability is often associated with mineral weathering, the decomposition
238 of organic matter is a key source of P in many ecosystems (Heuck et al. 2015; Margalef et al.
239 2017). Microbial decomposition increases the bioavailability of plant inputs and SOM by
240 cleaving polymeric compounds into simple C substrates that can be directly assimilated by
241 microorganisms. This process is often catalyzed by extracellular enzymes that release nutrients
242 from organic matter, connecting C, N, S, and P cycling (Allison et al. 2014). Viral predation can
243 also influence these dynamics by lysing specific clades of microorganisms and by influencing

244 organic matter accessibility to microorganisms (Kuzyakov & Mason-Jones 2018; Weinbauer
245 2004; Wilhelm & Suttle 1999). Global patterns of mechanisms governing SOM formation and
246 cycling and the specific environmental conditions under which certain metabolic pathways are
247 favored by microbial communities are active areas of research.

248 Identifying parameters that regulate decomposition across different ecosystems is
249 challenging because heterogeneity in nutrient statuses, substrate energetics, mineral and
250 hydrologic properties, and microbial communities can lead to different apparent controls in
251 different soils (Milcu et al. 2011; Sullivan et al. 2014). In some cases, rates of decomposition are
252 linked to microbial adjustments in the acquisition of specific elements in response to differences
253 between microbial biomass and substrate stoichiometries (i.e., nutrient limitations, Billings &
254 Ballantyne IV 2013; Billings & Ziegler 2008; Frost et al. 2005; Manzoni 2017; Manzoni et al.
255 2008; Spohn 2016; Sterner & Elser 2002). This research is rooted in the paradigm that organisms
256 exhibit stoichiometric homeostasis in which they maintain stable biomass elemental ratios
257 regardless of substrate stoichiometry. It leads to the prediction that nutrient limitations
258 universally regulate biological activity (Redfield 1958; Spohn 2016). Specifically, the relative
259 strength of the relationship between microbial metabolisms involved in N, S, or P acquisition and
260 SOM decomposition should be strongest in soils with the lowest underlying proportion of that
261 element. Indeed, the stoichiometry of microbial biomass features relatively little global variation
262 (60:7:1 C:N:P) (Cleveland & Liptzin 2007) and substrate stoichiometries are highly variable
263 (Frost et al. 2005; Manzoni 2017; Manzoni et al. 2008; Spohn 2016; Sterner & Elser 2002),
264 supporting a dynamic in which microorganisms must adapt to account for soil nutrient
265 concentrations.

266 The dependency of SOM decomposition on soil nutrients has been investigated most
267 frequently using a suite of biogeochemical techniques in observational or fertilization studies
268 such as CO₂ flux, mass loss, bulk chemistry (e.g., C:N), microbial biomass measurements, and/or
269 potential enzyme activity measurements (Sullivan et al. 2014). Each of these approaches
270 provides indirect evidence for nutrient-regulated microbial decomposition using comparatively
271 low molecular resolution. Additionally, the majority of researchers quantify total or coarsely-
272 defined lumped C pools with minimal characterization of specific SOM compound classes or
273 their stoichiometry. For instance, the classical pool and flux model focuses on three or four
274 operationally defined C pools that are connected by biological drivers of decomposition, namely
275 extracellular enzymes (Abramoff et al. 2018). As such, current estimates of nutrient limitations
276 on SOM decomposition vary widely (Cleveland et al. 2006; Grandy et al. 2008; Kirkby et al.
277 2013; Kirkby et al. 2014; Sullivan et al. 2014; Takriti et al. 2018), and the majority of research
278 has focused on a few elements, generally on linkages between C and N cycling (Cleveland et al.
279 2006; Grandy et al. 2008; Khan et al. 2016; Kirkby et al. 2013; Kirkby et al. 2014).

280 Microbial potential enzyme activity remains the primary tool in biogeochemical research
281 for measuring the decomposition potential of specific compound classes in SOM pools.
282 Extracellular enzymes catalyze the depolymerization of complex SOM compounds into simpler
283 molecules for direct uptake and are ubiquitous in soils (Allison et al. 2010). The standard suite of
284 enzyme assays target organic substrates including proteins, carbohydrates, amino sugars, organic
285 phosphates, and lignins (Allison et al. 2007; Burns 1982; Sinsabaugh & Shah 2011). However,
286 several microbially-mediated steps in SOM decomposition—notably those catalyzed by phenol
287 oxidases, peroxidases, and sulfatases which are more constrained by methodological limitations
288 in enzyme activity assays—are under investigated than others (e.g., widely-assayed NAG, BG,

289 phosphatases). Additionally, while the ratios of enzyme activities have been used to assess
290 nutrient limitations (Grandy et al. 2008; Hill et al. 2014; Jing et al. 2020; Moorhead et al. 2013;
291 Sinsabaugh et al. 2008), these assays can be influenced by pH and substrate availability and only
292 provide potential rates. They are low molecular resolution compared to many omics-based
293 approaches because they target chemical bonds contained by classes of SOM rather than specific
294 molecules, and they can have long turnover times that decouple them from SOM cycling at a
295 specific point in time (Schimel et al. 2017).

296 Microbial investigations into nutrient-regulated SOM decomposition have traditionally
297 used a complementary set of tools, revealing differences in metabolic strategies to maintain
298 stoichiometric homeostasis using lumped traits (e.g. community-level variation between bacteria
299 and fungi or copiotrophic and oligotrophic metabolisms, Elser et al. 2003; Fierer et al. 2007;
300 Strickland & Rousk 2010). Changes in microbial heterotrophy has also been suggested to vary
301 microbial preferences for C vs. N globally, an inference drawn from primarily biogeochemical
302 data (Taylor & Townsend 2010). Understanding how changes in metabolic strategies of
303 individual microorganisms respond to changes in substrate stoichiometry scale to a microbial
304 community's elemental use efficiency and capacity for decomposition is a key unknown in
305 process-based modelling.

306 Omics-based characterization of microbial communities (i.e., metagenomic,
307 metatranscriptomic, metaproteomic, and metabolomic) and other high molecular resolution
308 approaches (e.g., chemical probes) provide direct means to evaluate how stoichiometry regulates
309 the ecology of decomposition. These approaches have been highlighted as emerging tools in
310 ecological stoichiometry but are not yet widely implemented (Van de Waal et al. 2018).
311 Microbial genes and their expression patterns can reveal the genetic potential (genes), expression

312 (transcripts), and translation (proteins) of specific enzymes involved in decomposition and more
313 directly evaluate if biological N, S, and P acquisition vary predictably with SOM stoichiometry.
314 While these modern microbial approaches have their own limitations (e.g., data annotation and
315 discrepancies in turnover time when compared to rates), they provide a more complete
316 representation of biogeochemical cycles than enzyme potential assays.

317 Several potential analysis targets for stoichiometric regulation of SOM decomposition
318 have been identified in omics pipelines. For example, Finn et al. (2020) and Wilhelm et al.
319 (2019) proposed lists of relevant Kyoto Encyclopedia of Genes and Genomes (KEGG)
320 Orthologies and Carbohydrate-Active enzyme database (CAZy) annotations. These contain
321 enzymes involved in the decomposition of organic C as well as N, S, P, and iron cycles,
322 including commonly measured enzymes. They span many classes of enzymes involved in soil C
323 and nutrient cycling, including those targeting cellulose, hemi-celluloses, lignin, cellobiose,
324 mineral and organic N, mineral and organic P, and mineral and organic S. These provide a
325 tangible and comprehensive set of microbial genes, transcripts, proteins, and metabolites that is
326 consistent with biogeochemical methods used to investigate SOM decomposition. New
327 annotations for molecular markers involved in SOM decomposition continue to be discovered,
328 and our ability for molecular resolution into stoichiometry-based processes, if viewed within a
329 consistent framework to biogeochemical applications, can help illuminate patterns where
330 biogeochemical models lack mechanistic resolution.

331

332 **Opportunities for Ecological Stoichiometry and High-Resolution Measurements in SOM**
333 **Decomposition Modelling.**

334 Stoichiometric relationships underlie the kinetics of SOM decomposition (for instance,
335 through nutrient limitations) and therefore are essential to developing and incorporating omics-
336 enabled models into ecosystem scale models. The enormous amount of soil microbial and
337 chemical diversity provided by new molecular data types complicates these efforts (Jansson &
338 Hofmockel 2018) such that most current SOM decomposition models cannot yet account for
339 cutting-edge molecular approaches to represent detailed microbial processes. Additionally, there
340 is a need to move beyond models that separate SOM by chemical, physical, and functional
341 properties (e.g., pool-based models) to more fluid relationships between microorganisms and
342 SOM decomposition (Abramoff et al. 2018; Tang & Riley 2020; Waring et al. 2020). Microbial
343 data mining approaches have made incremental advances to this effort, but remaining obstacles
344 underline an opportunity for broader use of ecological frameworks, such as stoichiometric
345 theory, to guide targeted investigations particularly with respect to unprotected pools of SOM
346 (Fig. 2). This approach can help determine which pathways to include in any metabolic model
347 and further scale high molecular resolution predictions to soil ecosystems (Zakem et al. 2020).
348 Below, we highlight opportunities for ecological stoichiometry to inform the development and
349 merging of high-resolution models such as omics-enabled reaction networks with
350 biogeochemical models at larger scales (Fig. 4).

351 State-of-science models already implicitly use ecological stoichiometry to unify
352 microbiology and biogeochemistry at intermediate scales, as they constrain rates of SOM
353 decomposition with parameters that represent nutrient availability, stoichiometries of organic
354 substrates or microbial biomass, and carbon and nutrient use efficiencies [e.g., MEND (Wang et
355 al. 2020a; Wang et al. 2015), RESOM (Tang & Riley 2015), and MIMICS-CN (Kyker-Snowman
356 et al. 2020)]. These models arise from biogeochemical research and are based on C pools with

357 lumped characteristics. However, they generally include two elements at most, in part due to
358 computational barriers to representing all processes in all elemental cycles as well as their
359 interconnectivity (Abramoff et al. 2018; Moorhead et al. 2012; Sulman et al. 2019; Wang et al.
360 2015; Wieder et al. 2015).

361 A small set of coarse parameters is often used to denote microbial processes within state-
362 of-science biogeochemical models, in particular C and nutrient use efficiencies (CUE, NUE) that
363 describe C and nutrient conversion into microbial biomass (Allison et al. 2010; Cleveland &
364 Liptzin 2007; Geyer et al. 2016; Min et al. 2016; Sihi et al. 2016; Sinsabaugh et al. 2013;
365 Sinsabaugh et al. 2016; Wang et al. 2015; Wieder et al. 2014). CUE and NUE incorporate the
366 principles of ecological stoichiometry into microbial-explicit biogeochemical models by serving
367 as proxies for generalized microbial activity. They are most often estimated from empirical
368 measurements of organic matter stoichiometry and/or microbial biomass, or from the ratios of C-
369 vs. nutrient-acquiring enzymes (Manzoni et al. 2012; Sinsabaugh et al. 2016). Under nutrient
370 limitation, microorganisms reduce their C uptake, acquire nutrients, and/or respire or excrete
371 excess C as enzymes or metabolites such that C assimilation into biomass often declines
372 (Anderson et al. 2005; Del Giorgio & Cole 1998; Manzoni 2017; Manzoni et al. 2017; Manzoni
373 et al. 2008; Middelboe & Sørensgaard 1993; Milcu et al. 2011; Mooshammer et al. 2014b; Smith
374 & Prairie 2004; Sterner & Elser 2002). Thus, CUE typically decreases when nutrients are
375 limiting. In contrast, when nutrients exist in excess, microorganisms adjust their uptake or rate of
376 respiration/excretion to decrease NUE (Milcu et al. 2011; Mooshammer et al. 2014a).

377 The balance of stoichiometric impacts on decomposition is difficult to predict due in part
378 to microbial interactions (i.e., stoichiometry influences individual behavior rather than
379 community-level behavior that ultimately regulates decomposition, Manzoni 2017; Manzoni et

380 al. 2012; Sterner & Elser 2002), highlighting an opportunity for improvement through omics-
381 enabled research. There is also large variation in the relationship of microbial CUE and NUE to
382 SOM decomposition due to spatial and temporal changes in resource availability, microbial
383 community structure, and soil physical properties (Frey et al. 2013; Herron et al. 2009; Malik et
384 al. 2018; Manzoni et al. 2012; Qiao et al. 2019; Sinsabaugh et al. 2013; Sinsabaugh et al. 2016).
385 Biogeochemical models have shown that small changes in CUE can have large impacts on SOM
386 decomposition (Li et al. 2014; Six et al. 2006; Wieder et al. 2013), offering an intriguing
387 possibility for model improvement through detailed representations of nutrient-regulated
388 microbial metabolic pathways to develop scalable omics-enabled reaction networks.

389 At the finest molecular resolution where microbial information is best-suited (Fig. 3),
390 there is a limited set of modelling approaches that currently predict rates of SOM decomposition,
391 largely through by representing genomic and/or metabolic processes. Omics-enabled metabolic
392 models, in particular, infer function from molecular data types by placing annotations within the
393 context of a cell's biochemical abilities to consume substrates, grow, and produce energy
394 (Cuevas et al. 2016). They are capable of including thousands of genes, metabolites,
395 transcriptomes, proteomes, and associated reactions (Lieven et al. 2020; Seaver et al. 2020).
396 However, they are limited by a set of biochemical reactions selected *a priori* that define the
397 possible functions of microorganisms and provide a structure in which to root functional
398 annotations (Seaver et al. 2020). In tandem, flux balance analysis is widely used within omics-
399 enabled models to predict fluxes through the reactions in a metabolic network (Cuevas et al.
400 2016; Orth et al. 2010). These omics-enabled models can also be coupled to other methods of
401 decoding high-resolution molecular data, such as thermodynamic theory (Garayburu-Caruso et

402 al. 2020; Song et al. 2020), deep learning (Zampieri et al. 2019), or network modelling (Kessell
403 et al. 2020; McClure et al. 2020) to further extend their applicability.

404 We see opportunity for ecology stoichiometry to provide this framework by guiding
405 representations of microbial metabolism in biogeochemical models, particularly where
406 stoichiometric regulation already exists in the unprotected or biologically-available SOM pools
407 of many microbial-explicit biogeochemical models (e.g., MEND-CN, MIMICS-CN), and as a
408 complement to abiotic controls in mineral soil (Fig. 4). Ecological stoichiometry influences the
409 behavior of individual microorganisms, as they act to maintain stoichiometric balance through
410 adjusting growth rate or biomass stoichiometry and/or mining or excreting nutrients. This
411 understanding could be incorporated into individual-scale models by adjusting active metabolic
412 processes based on environmental nutrient context, for instance through metabolic flux analysis.
413 Increasing in scale towards trait-based models (e.g., Malik et al. 2020; Wang & Allison 2021),
414 ecological stoichiometry operates by impacting the growth rates of organisms with particular
415 traits – for example, microorganisms with specific nutrient acquisition pathways, r- and K- life
416 strategies, stress tolerance, or other competitive advantages. Incorporation of stoichiometric
417 principles such as substrate stoichiometry and/or environmental nutrient status at this scale can
418 guide an individual's expressed metabolisms in the context of other microorganisms, as well as
419 the interactions of their metabolites and extracellular enzymes, and provides a more
420 computationally favorable model structure. In turn, these omics-enabled representations can
421 provide resolution into SOM cycling that currently aggregates microbial activity via lumped
422 SOM pools and complement existing representations of mineral-sorption processes (Fig. 4).

423 Ecological stoichiometry can also aid in merging models from microbiology and
424 biogeochemistry by constraining the need for detailed representations of specific SOM

425 decomposition pathways based on environmental context. In a bottom-up approach, ecological
426 stoichiometry provides an overarching hypothesis and sets *a priori* expectations for pathways of
427 interest to assist in the interpretation of omics measurements (Fig. 4). Using either hypothesis-
428 driven experiments or observational data from ecological networks, data mining efforts could
429 focus on nutrient acquisition metabolic pathways (e.g., peptidase-, phosphatase-, and sulfatase-
430 including metabolic pathways) in order of hypothesized nutrient limitations in order to compress
431 a multi-omic data set of nearly any size to a manageable amount of data. Data from multiple
432 sources across ecoregions can be leveraged in this way to move towards transferrable principles
433 of SOM decomposition. Metabolic pathways with the most predictive power of SOM
434 decomposition (inferred from transcripts, proteins, and metabolites mapping to metabolic
435 pathways) then guide the portions of intermediate-scale models in which greater resolution of
436 these pathways could lead to more accurate predictions. This could result in inclusion and/or
437 calibration of specific portions of SOM pools that are most tightly coupled to decomposition
438 rates [e.g., protein or necromass pools] or of specific metabolic pathways instead of lower
439 resolution measurements such as CUE and NUE. Dynamically activating omics-enabled reaction
440 networks for a specific nutrient status could provide the maximum benefit of high molecular
441 resolution while being computationally tractable. In turn, parameters and rate estimates made at
442 the intermediate scale can then influence the relative importance of variables in low molecular
443 resolution models such as CUE/NUE, bacterial:fungal biomass, and/or biomass C:N:S:P.

444 In parallel, top-down approaches use coarsest measurement scale (e.g., bulk
445 biogeochemistry) to first identify nutrient limitation(s) that then guide model development at
446 higher molecular resolutions. Patterns in nutrient concentrations and pool sizes can inform the
447 selection of limited high-resolution parameters for inclusion in reduced complexity models

448 thereby optimizing added value from molecular techniques. For instance, soil nutrient status at a
449 given point in time can be inferred from bulk measurements of soil biogeochemistry and nutrient
450 amendment experiments, and subsequently used to define a set of reactions that should most
451 strongly impact rates of SOM decomposition. Examples might include representing organic N
452 cycles (or amino acid metabolisms therein) for N limitation, microbial biomass recycling
453 (necromass decomposition) for P limitation, and/or organosulfatases for S limitation—or
454 combinations therein in the case of nutrient co-limitation—in omics-enabled reaction networks
455 (Fig. 4). Outcomes from models with enhanced representation of specific SOM decomposition
456 pathways can then further point to specific microbial genes/transcripts/proteins in pathways
457 associated with the rate limiting step in SOM decomposition (if biological). Using this approach,
458 we can narrow down molecular data to specific microbial metabolisms or omics markers of
459 interest to build omics-enabled models including only the most relevant parameters that are
460 scalable to the level of soil ecosystems. This approach reduces the computational cost of high
461 resolution reaction networks by only representing portions of a soil ecosystem that most strongly
462 impact SOM decomposition in detail.

463 When used in combination, *the result of these iterative approaches is a new suite of*
464 *models that represent a limited set of microbial metabolisms, guided by stoichiometric principles*
465 *and high-resolution molecular measurements, that are computationally feasible and more*
466 *accurately predict SOM decomposition, termed ‘omics-enabled biogeochemical models’* (Fig. 4).
467 Such models address a major challenge to the next generation of predictive models—deriving
468 omics-enabled reaction networks that can improve assessments of soil C storage and emissions
469 across molecular-to-ecosystem scales.
470

471 **Opportunities afforded by investments in new molecular and computational technologies**

472 The power of ecological stoichiometry to guide model development across scales of
473 biological complexity can increase even further when leveraged in combination with new
474 advances in molecular technologies and machine learning algorithms. Research that
475 simultaneously investigates microbial roles in C, N, S, and P cycling is rare, and enzyme activity
476 assays are imprecise relative to omics or probe-based assays. Investments in untargeted
477 molecular approaches are generating new SOM cycling analyses by allowing detection of active
478 microbial metabolisms. These are nascent technologies that have the potential to improve our
479 understanding of SOM decomposition but need greater usage and development to surmount
480 remaining challenges.

481 We especially highlight stable isotope probing (SIP, Bernard et al. 2007; Dumont &
482 Murrell 2005; Pepe-Ranney et al. 2016), new ultrahigh-resolution metabolomics (Tfaily et al.
483 2017), activity-based protein profiling (ABPP, Killinger et al. 2019), and machine learning as
484 new classes of technologies deserving of further development and broader consideration. Briefly,
485 microbial applications of SIP use substrates enriched in heavy isotopes (e.g., ^{13}C , ^{18}O , ^{15}N)
486 amended to microbial communities to reveal isotope-labeled biomarkers of active microbial
487 populations (Bernard et al. 2007; Dumont & Murrell 2005; Pepe-Ranney et al. 2016). While SIP
488 has long been used in biogeochemistry, recent advances are increasing the tractability of SIP
489 within amplicon profiles, metagenomes, metatranscriptomes, and metaproteomes (Barnett &
490 Buckley 2020; Jameson et al. 2017; Pepe-Ranney et al. 2016; Wilhelm et al. 2019; Youngblut et
491 al. 2018a; Youngblut et al. 2018b). In parallel, ultrahigh-resolution metabolomics, such as
492 Fourier transform ion cyclotron resonance mass spectrometry (FTICR-MS, Tfaily et al. 2015;
493 Tfaily et al. 2017), can illuminate new aspects of SOM chemistry by enabling the determination

494 of elemental stoichiometry in thousands of SOM molecules simultaneously (Tfaily et al. 2017).
495 Formulas assigned via ultrahigh-resolution metabolomics are a promising tool for evaluating
496 metabolisms influenced by ecological stoichiometry by providing information on molecules with
497 specific stoichiometries (e.g., C:N ratios of SOM molecules in a standing pool) or changes in
498 these molecules during the course of experiments. Finally, new targeted approaches can improve
499 the accuracy of existing enzyme-based approaches. Activity-based protein profiling (ABPP,
500 Killinger et al. 2019), for example, leverages chemical probes that react irreversibly with protein
501 families to identify enzymes binding to specific molecules, and when paired with mass
502 spectrometry-based approaches, can also quantify rates of enzyme production (Killinger et al.
503 2019). In contrast to enzyme activity assays that reveal potential enzyme rates, ABPP has the
504 potential to dramatically improve investigations of SOM decomposition by revealing the true
505 expression of decomposition enzymes; however, it has never been used in soils (Sadler & Wright
506 2015; Whidbey & Wright 2018; Zegeye et al. 2020).

507 Even with technological advances, the interpretation of untargeted molecular approaches
508 is complicated by the inadequate description of metabolic pathways for soil microorganisms in
509 existing databases and the unsuitability of standard statistical approaches in ecology for
510 combining multidimensional data types. Unsupervised machine learning is emerging as a tool
511 both to decipher hidden patterns in complex data and to eliminate the needs for *a priori*
512 relationships between genes (e.g., via metabolic pathways or metagenome-assembled genomes)
513 and for gap filling to infer absent data. There is an enormous diversity of machine learning
514 algorithms with varying levels of complexity that are well-suited to discern patterns in data
515 (Rana et al. 2020; Were et al. 2015). Techniques such as K-nearest-neighbor (Peterson 2009),
516 artificial neural network (ANN, Sarle 1994), support vector machines (SVM, Wang 2005),

517 neuro-fuzzy (Nauck et al. 1997), decision tree classifiers (Safavian & Landgrebe 1991), and
518 random forests (Liaw & Wiener 2002) are most commonly employed in microbiology and
519 environmental disciplines (e.g., Cai et al. 2019; De Clercq et al. 2019; Dong & Chen 2019; Qdais
520 et al. 2010; Rahimian Boogar et al. 2019; Thompson et al. 2019; Wang et al. 2020b). Still, we
521 lack rigorous machine learning investigations in which sampling and experimental designs move
522 beyond feature identification to translating selected features into meaningful ecological
523 outcomes. Pairing machine learning with ecological theory for experimental hypothesis testing is
524 needed to push fundamental biology forward (Rana et al. 2020; Were et al. 2015). Using new
525 high-resolution molecular data types in spatially- and temporally-distributed sampling, paired
526 with machine learning, is the gold standard for future research in SOM decomposition and can
527 offer unprecedented untargeted resolution into the pathways by which microorganisms
528 decompose SOM (Manzoni et al. 2017; Mooshammer et al. 2014b).

529

530 **Conclusions.**

531 We have made significant advances in understanding global C cycles and now face the
532 challenge of merging concepts and modelling frameworks from different scientific fields. To
533 facilitate this effort, we identify substrate chemistry as a linkage point between microbiology and
534 biogeochemistry, and we propose using ecological stoichiometry as a unifying framework that
535 can help overcome a major limitation to understanding global SOM decomposition —
536 interpreting the massive amount of data generated by molecular approaches and translating them
537 to the ecosystem scale. Our hope is that this common foundation can maximize the strengths of
538 models across scales and environmental contexts, and we encourage better communication
539 among empiricists and modelers within and across domains. Investing in and deploying new

540 technologies in the context of stoichiometric theory provides an untapped and promising avenue
541 for interpreting the vast amount of data we can now generate and further distilling this
542 information into a new generation of omics-enabled biogeochemical models.

543

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550

551 **Declarations.**

552

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558

559 *Conflicts of interest/Competing interests.* We declare no conflicts of interest/competing interests.

560

561 *Availability of data and material (data transparency).* Not applicable.

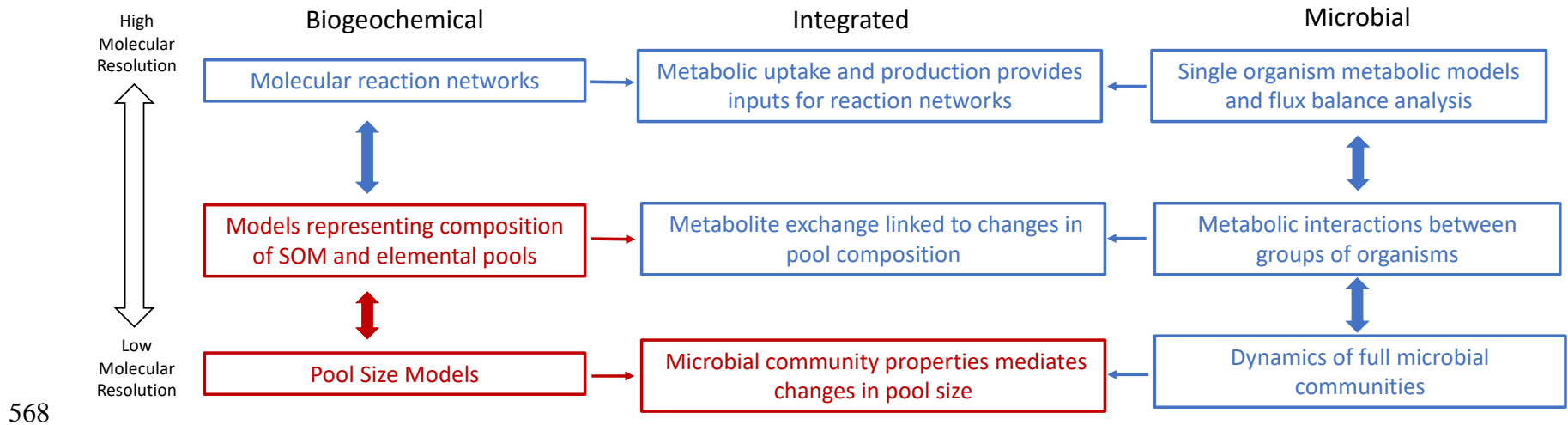
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563 *Code availability (software application or custom code)*. Not applicable.

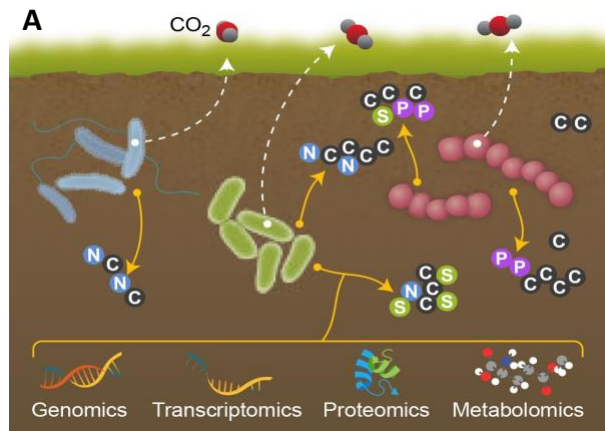
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565 *Authors' contributions*. EBG conceived and wrote the manuscript with significant contributions

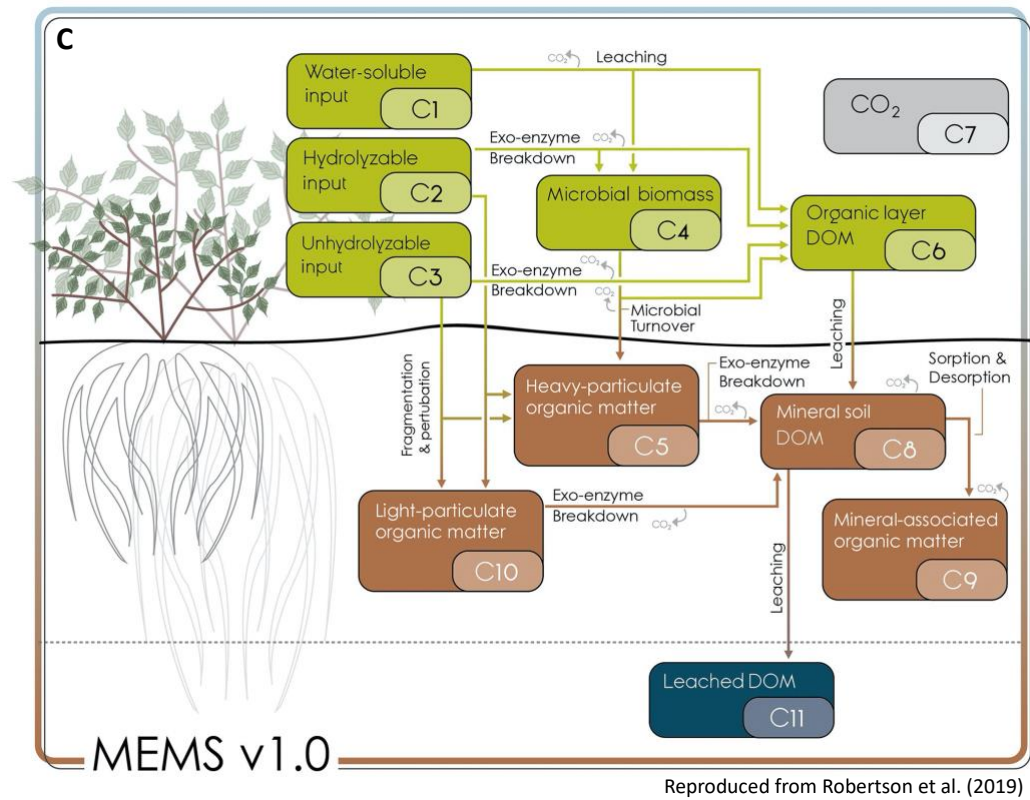
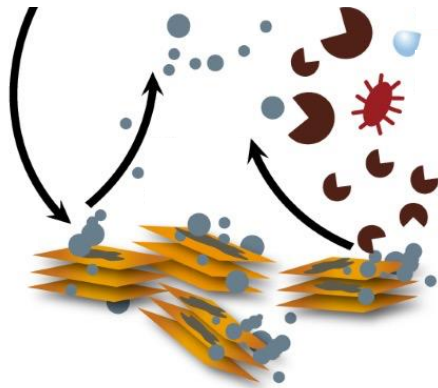
566 from KSH.

567 **Figures.**

569 **Fig 2. Parallel development of processed-based biogeochemical and microbial models across scales of molecular resolution.** A
570 short description of a representative class of models from biogeochemical (left) vs. microbiological (right) research domains is listed
571 within each box. Model types that bridge between microbial and biogeochemical research domains are in the center column. High
572 molecular resolution models are on the top row and low molecular resolution models are on the bottom row. Modelling frameworks
573 that leverage ecological stoichiometry are depicted in red, and opportunities to use these principles to guide multiscale model
574 development are depicted in blue. We present opportunities for stoichiometry to serve as an underlying framework to integrate state-
575 of-science SOM decomposition models across ecological hierarchies – from individual microorganisms to bulk pools and fluxes – as
576 well biogeochemical and microbial disciplines. Ecological stoichiometry is implicitly used to guide both biogeochemical and
577 microbiological models and therefore presents a natural linkage point for assimilating and scaling models from both research domains.



B Reproduced from Jilling et al. (2018)

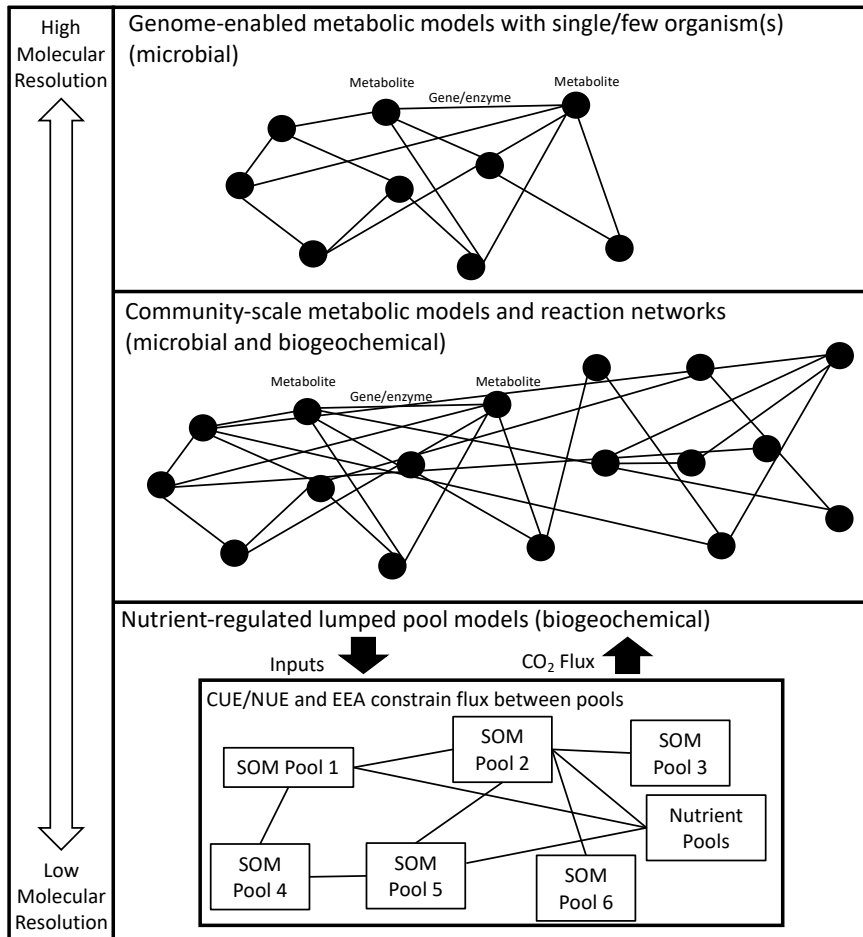


Reproduced from Robertson et al. (2019)

578

579 **Fig. 2. Overview of Factors Regulating Soil Carbon Cycles.** (A) Microbial metabolism is the proximate control over bioavailable C
 580 cycling. These controls dominate in unprotected SOM pools in which accessibility is not limited and provide an avenue for
 581 stoichiometry to improve model predictions. (B) Mineral-association and hydrologic variables are also key components of SOM

582 cycling, particularly in deeper soil horizons with high mineral content. (C) These conceptualizations are consolidated within the
583 MEMS framework. In MEMS, SOM is more readily bioavailable in the organic horizon where growth of microbial biomass and to a
584 lesser extent dissolved organic matter support C flow to and stabilization within mineral horizons. Panels (B) and (C) are reproduced
585 from Jilling et al. (2018) and Robertson et al. (2019).

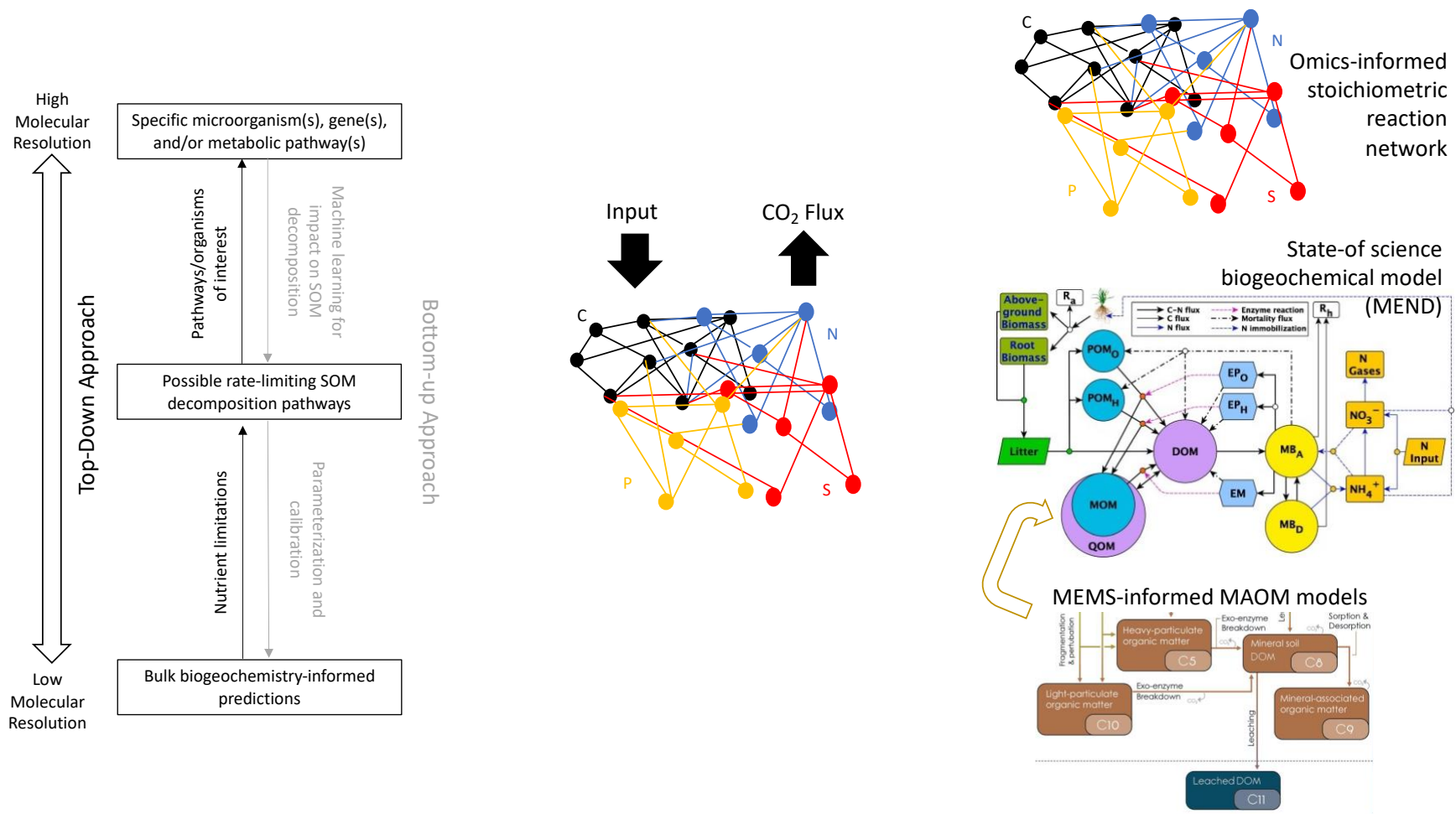


586

587 **Fig 3. Representation of state-of-science SOM decomposition models across scales of molecular resolution.** At the highest level
 588 of molecular resolution (top), genome-enabled models predict the production and consumption of specific metabolites by specific
 589 microorganisms or groups of microorganisms. This class of models is rooted in fundamental microbiology. At intermediate scales,

590 community-scale (mixed-bag) metabolic models and reaction networks have been largely separately developed in microbiology and
591 biogeochemistry, respectively. Finally, at coarser levels of molecular resolution, biogeochemical models predict SOM decomposition
592 using nutrient-regulated exchange between lumped SOM pools with specific chemical attributes.

593



594

595 **Fig 4. Schematic of opportunities for ecological stoichiometry to guide iterative approaches that generate omics-enabled**596 **biogeochemical models.** Ecological stoichiometry can guide a cycle of top-down and bottom-up approaches for more efficient

597 generation of next generation models. Top-down approaches use nutrient limitations inferred from bulk biogeochemical measurements

598 to guide the determination of rate-limiting pathways in SOM decomposition. In turn, rate-limiting pathways can reveal specific
599 metabolic pathways and/or microorganism for detailed representation in predictive models. Bottom-up approaches use ecological
600 stoichiometry as a framework for guiding the interpretation of high molecular resolution data streams with the help of new machine
601 learning algorithms, for example by focusing on metabolic pathways that tend to be involved in nutrient acquisition. Using these
602 approaches to identify the rate-limiting steps of SOM decomposition can then aid in the parameterization and calibration of larger-
603 scale models by informing the tuning of substrate use efficiencies and/or by revealing pathways for more detailed representation. By
604 using both approaches iteratively, we can identify specific portions of genome-enabled reaction networks that add predictive power
605 with detailed representation in large-scale SOM decomposition models. A schematic of iterative top-down and bottom-up approaches
606 is shown on the left. This leads to the development of a stoichiometry-informed reaction networks (middle) that can be merged into
607 existing state-of-science SOM decomposition models (right).

608

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