

1 **Potential bioavailability of representative pyrogenic organic matter**  
2 **compounds in comparison to natural dissolved organic matter pools**

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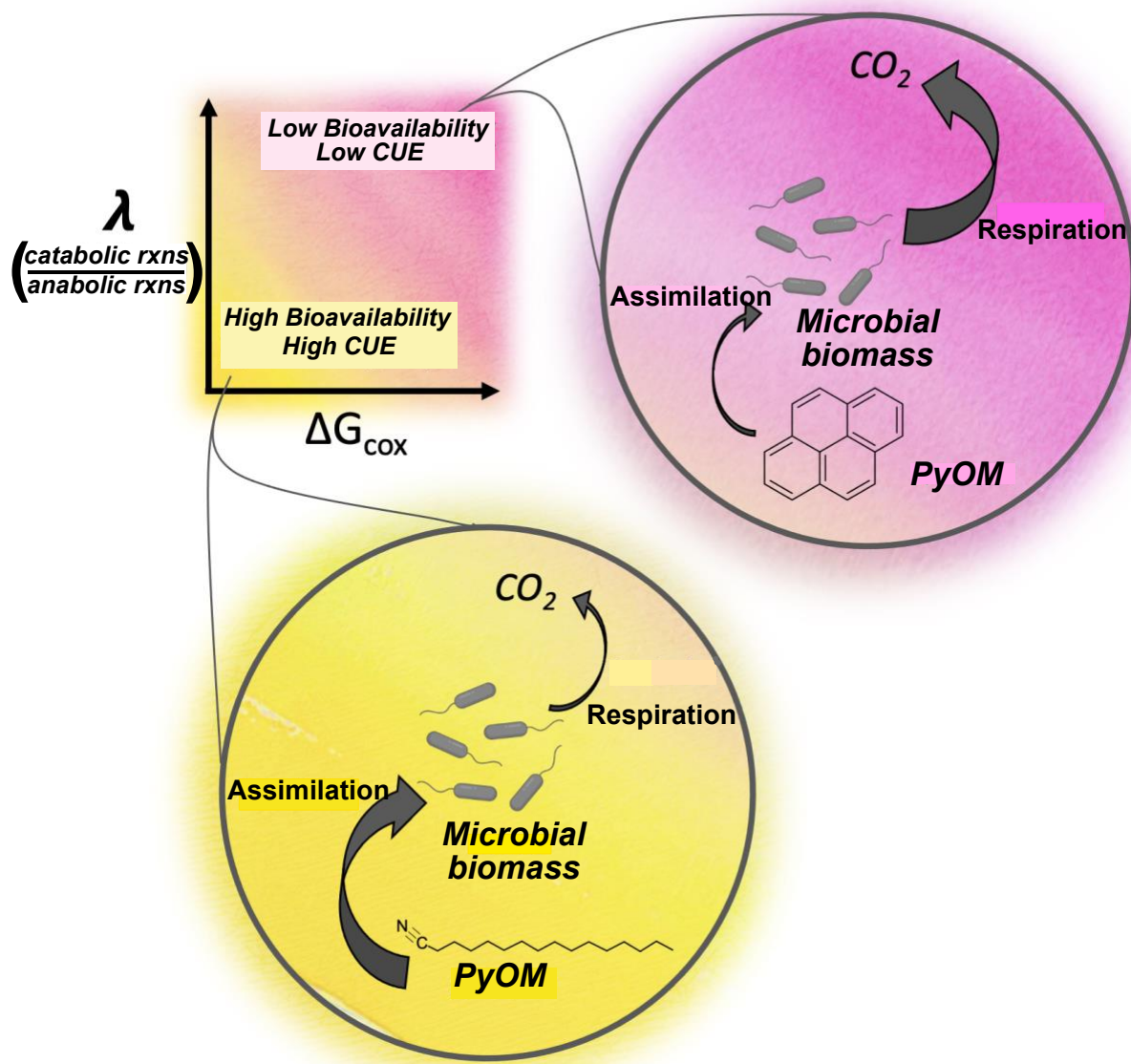
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## 20 GRAPHICAL ABSTRACT



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22

23 **Abstract.** Pyrogenic organic matter (PyOM) from wildfires impacts river corridors globally and  
24 is widely regarded as resistant to biological degradation. Though recent work suggests PyOM  
25 may be more bioavailable than historically perceived, estimating bioavailability across its  
26 chemical spectrum remains elusive. To address this knowledge gap, we assessed potential  
27 bioavailability of representative PyOM compounds relative to ubiquitous dissolved organic  
28 matter (DOM) with a substrate-explicit model. The range of potential bioavailability of PyOM  
29 was greater than natural DOM; however, the predicted thermodynamics, metabolic rates, and  
30 carbon use efficiencies overlapped significantly between all OM pools. Compound type (e.g.,  
31 natural vs. PyOM) had approximately 6-fold less impact on predicted respiration rates than  
32 simulated carbon and oxygen limitations. Within PyOM, the metabolism of specific chemistries  
33 differed strongly between unlimited and oxygen limited conditions—anhydrosugars, phenols,  
34 and polycyclic aromatic hydrocarbons degradations were more favorable under oxygen  
35 limitation than other molecules. Notably, amino sugar-, protein-, and lignin-like PyOM had  
36 lower carbon use efficiencies relative to natural DOM of the same classes, indicating potential  
37 impacts in process-based model representations. Overall, our work illustrates how similar PyOM  
38 bioavailability may be to that of natural DOM in the river corridor, furthering our understanding  
39 of how PyOM may influence riverine biogeochemical cycling.

## 40 **1 Introduction**

41 Wildfires have burned an average of 1.8-million ha year<sup>-1</sup> in the United States alone over  
42 the past 80 years, dramatically impacting terrestrial and aquatic ecosystems (Bladon et al., 2014;  
43 Shakesby and Doerr, 2006; Randerson et al., 2006; Verma and Jayakumar, 2012). As wildfire  
44 activity continues to increase in response to climate change (Pierce et al., 2004; Bowman et al.,  
45 2020; Flannigan et al., 2009), its impact on river corridor biogeochemistry is receiving  
46 significant attention (Wagner et al., 2018; Abney et al., 2019).

47 Pyrogenic organic matter (PyOM) generated by wildfires influences river corridor  
48 biogeochemistry due to the importance of organic matter as a carbon (C) and energy source.  
49 Though there is substantial uncertainty in the quantification of PyOM, in part due to its chemical  
50 diversity, estimates suggest that 116–385 Tg of condensed PyOM is generated per year. This  
51 amounts to 300 to 500 Tg C stored in sediments, soils, and waters (Jaffé et al., 2013; Dittmar et  
52 al., 2012; Hockaday et al., 2007; Santin et al., 2016) and ~10% of dissolved organic C pools in  
53 surface waters (Jaffé et al., 2013). Given that organic matter drives aquatic biogeochemical  
54 cycles, the loading of PyOM into river corridors has the potential to produce substantial impacts  
55 on ecosystem functions and downstream drinking water treatability (Emelko et al., 2011; Hohner  
56 et al., 2017).

57 While some estimates place aquatic residence times at thousands of years (Meyer and  
58 Wells, 1997; Elliott and Parker, 2001; Bigio et al., 2010; Kuzyakov et al., 2014), recent work has  
59 shown that PyOM may be more bioavailable (i.e., biolabile or able to be accessed and degraded  
60 by microorganisms) than previously thought (Myers-Pigg et al., 2015; Norwood et al., 2013;  
61 Zimmerman and Ouyang, 2019). This inference is also supported by research on biochars,  
62 highlighting the diverse reactivities of combustion byproducts (Sohi et al., 2010; Mia et al.,

63 2017). PyOM bioavailability may, therefore, play an unrecognized role in global biogeochemical  
64 cycles and climate feedbacks. Yet, we are just beginning to understand its potential  
65 bioavailability (Zimmerman and Mitra, 2017; Wozniak et al., 2020; Masiello, 2004).

66 We used a substrate-explicit model to assess the potential bioavailability of PyOM in  
67 comparison to dissolved organic matter (DOM) from global surface waters and sediments. We  
68 derived 16,971 representative PyOM compounds from the primary literature describing plant  
69 charcoals, biochar, smoke, burnt soil, and their leachates (Table S1). Natural DOM pool  
70 composition was derived from a recent high-resolution survey of river corridor sediments and  
71 surface waters (Garayburu-Caruso et al., 2020a, Figure S1). The substrate-explicit model used  
72 molecular formulae to predict energetic content, metabolic efficiency, and rates of aerobic  
73 metabolism for organic substrates. It has proven useful in linking DOM composition to aerobic  
74 metabolism in natural environments (Song et al., 2020; Graham et al., 2017; Garayburu-Caruso  
75 et al., 2020b). It provides a systematic way to formulate reaction kinetics and is agnostic of many  
76 factors that have complicated a universal understanding of OM bioavailability; including  
77 molecular structure, chemical inhibition, mineral-associations and physical protection, terminal  
78 electron acceptors, microbial community composition and accessibility, and abiotic reactions  
79 (reviewed in Arndt et al. (2013)). We hypothesize that PyOM compounds have a greater  
80 similarity in their potential bioavailability to natural DOM pools than expected based on  
81 historical literature.

82

## 83 **2 Results and discussion**

84 Our model-based approach enabled us to directly compare known combustion products to  
85 thousands of ubiquitous DOM compounds, which would have been unfeasible to directly assess

86 in a laboratory setting. Key outputs from the model include lambda ( $\lambda$ , the number of catabolic  
87 reactions that must occur to provide the energy required for the synthesis of one mole of biomass  
88 carbon), Gibbs free energy of C oxidation ( $\Delta G_{Cox}$ ), under standard conditions with a modification  
89 to pH 7, carbon use efficiency (CUE), and the rate of aerobic metabolism (as oxygen consumed  
90 per mol-C biomass produced) under three scenarios: (a) no limitations, (b) C-limitation and (c)  
91 oxygen ( $O_2$ ) limitation. Though the relative magnitude of these metrics will vary based on the  
92 specific stoichiometry of a molecule, highly bioavailable compounds are indicated by low  $\lambda$  and  
93  $\Delta G_{Cox}$  coinciding with high CUE and metabolic rates. Low  $\lambda$  values denote efficient energetics  
94 of catabolism in producing biomass through anabolism. Low  $\Delta G_{Cox}$  denotes high thermodynamic  
95 favorability in an electron donor half reaction associated with organic matter, and high CUE  
96 reflects more C assimilated into biomass per unit C respired.

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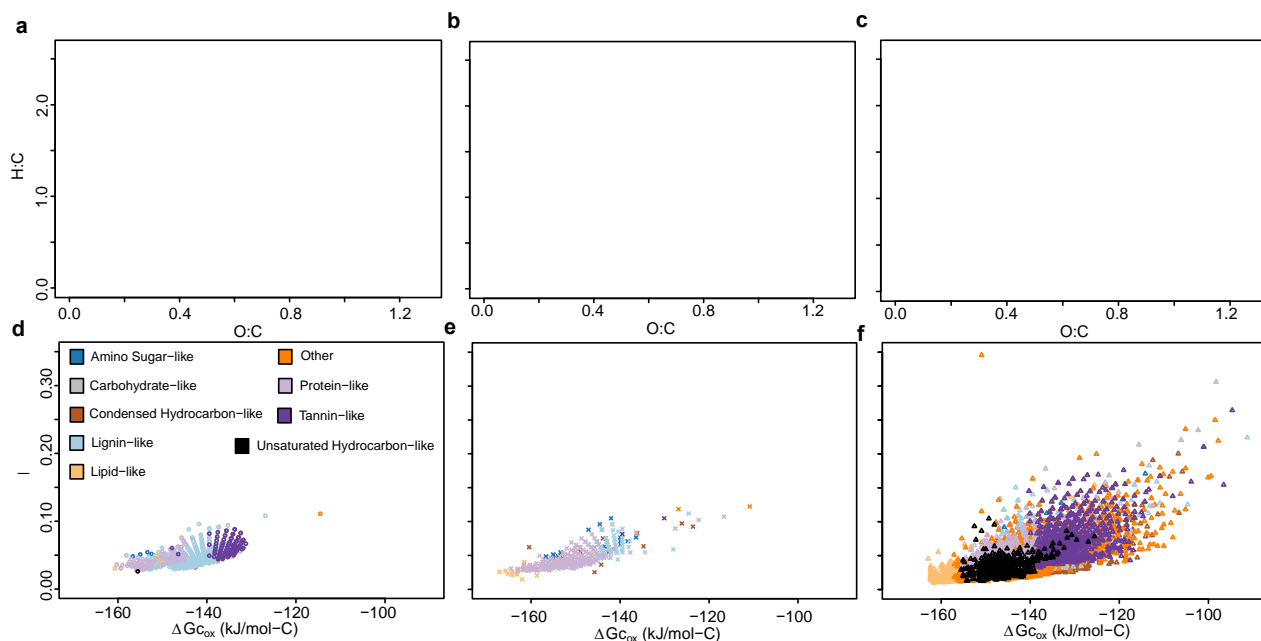
### 98 *2.1 Potential Bioavailability of Pyrogenic Organic Matter*

99        Though previous work has shown that sediment and surface water DOM is altered by  
100 wildfires (Cawley et al., 2018; Jaffé et al., 2013; Wagner et al., 2018), our results suggest that the  
101 chemically distinct pools of PyOM transported to aquatic systems may have similar potential  
102 bioavailability to DOM. We found that the ranges of  $\Delta G_{Cox}$ ,  $\lambda$ , and CUE overlapped substantially  
103 between PyOM and DOM in sediments and surface waters (Figure 1 and 2a-b, Figure S2).  
104 Predicted CUE of PyOM classes was also comparable to literature values reported by others  
105 (Saifuddin et al., 2019; Domeignoz-Horta et al., 2020; Pold et al., 2020). While  $\lambda$ ,  $\Delta G_{Cox}$  and  
106 CUE were significantly different when comparing all three groups (ANOVA,  $P < 0.001$ ,  $F =$   
107 16.12, 604.9, and 461.2 respectively),  $\lambda$  did not vary significantly between sediment DOM and  
108 PyOM (Tukey HSD,  $P = 0.21$ ) and was more similar between sediment DOM and PyOM than

109 between the surface water and sediment DOM pools (mean difference, surface water-sediment  
110 DOM = 0.0025, sediment DOM-PyOM = 0.0019) . The range of all PyOM metrics was greater  
111 than that of surface water or sediment DOM, reflecting the known heterogenous nature of PyOM  
112 chemistries. Though the  $\lambda$  of many PyOM compounds was higher than the maximum observed  
113 in either DOM pool, minimum values of  $\lambda$  were lowest for PyOM (surface water DOM = 0.026,  
114 sediment DOM = 0.015, PyOM = 0.010), indicating that a portion of wildfire-derived molecules  
115 may be more bioavailable than extant DOM pools. It is also interesting that surface water and  
116 sediment DOM had greater dissimilarity in  $\Delta G_{Cox}$  and CUE than any comparison involving  
117 PyOM. For example, the mean differences in  $\Delta G_{Cox}$  and CUE between surface water and  
118 sediment DOM were 7.34 kJ/mol-C and 0.058. The differences between PyOM and surface  
119 water were 4.28 kJ/mol-C for  $\Delta G_{Cox}$  and 0.024 for CUE. Overall, statistical differences were not  
120 surprising given an extremely large sample size (sediment  $n = 398$ , surface water  $n = 811$ , PyOM  
121  $n = 16,971$ ), and the low effect sizes denote that differences in bioavailability between PyOM  
122 and DOM were minimal despite statistical separation.

123         These results signal a strong overlap between the potential bioavailabilites of PyOM and  
124 DOM pools; however, within PyOM compounds, there was variability in  $\Delta G_{Cox}$ ,  $\lambda$ , and CUE  
125 consistent with heterogeneous chemistries (Figure 1 and 2). This is not surprising, given the  
126 diversity of PyOM molecules generated by wildfires of different burn severities and source  
127 materials (Wagner et al., 2015; Wagner et al., 2018; Neary et al., 2005), some of which are  
128 chemically similar to natural DOM.

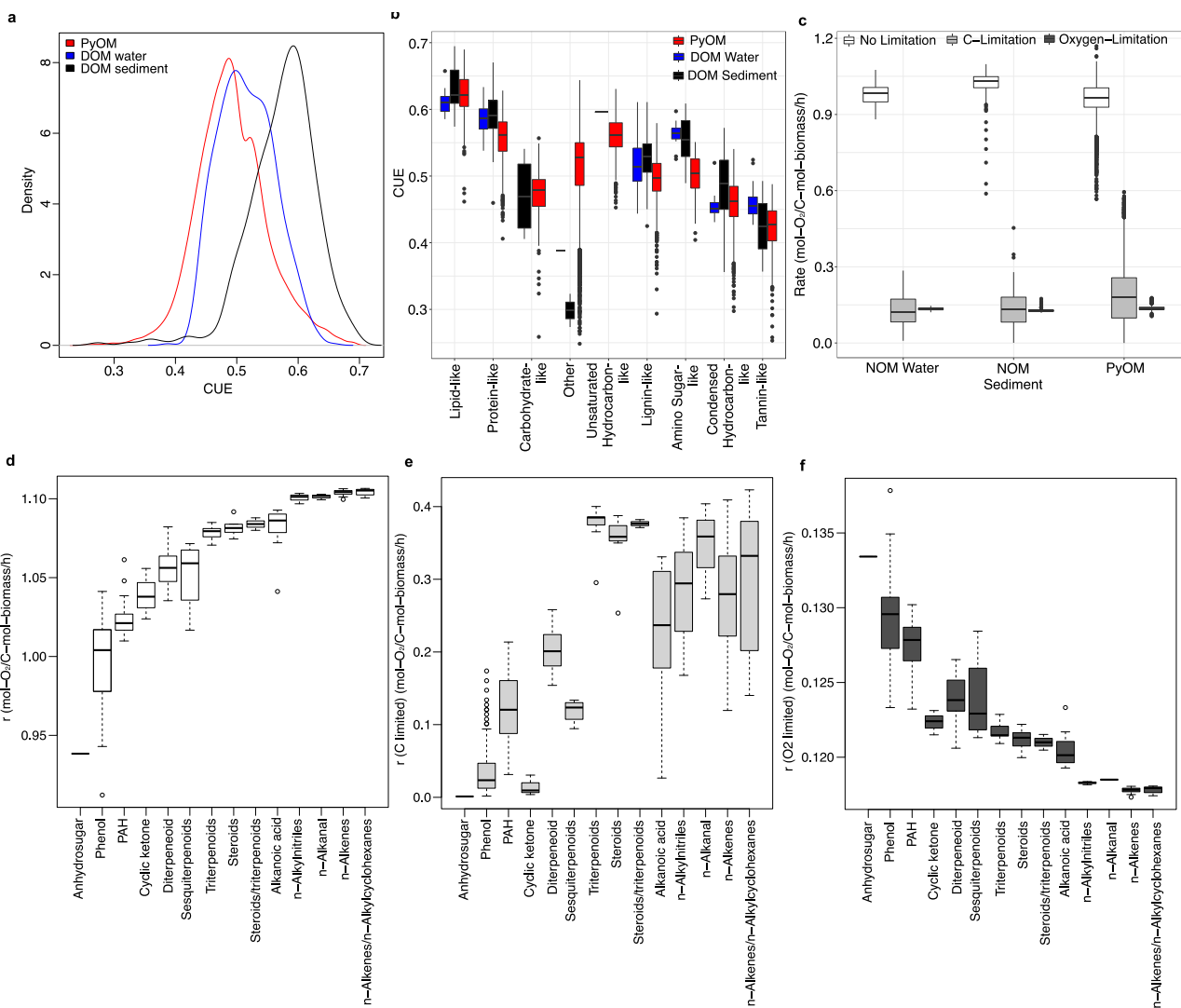
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131 **Figure 1.** Comparison of natural DOM and PyOM. (a-c) Van Krevelen diagrams of molecules  
 132 that were present in at least 95% of (a) surface water or (b) sediment samples versus (c)  
 133 representative PyOM. (d-f) Distribution of energy content ( $\Delta G_{CoX}$ ) vs. metabolic efficiency ( $\lambda$ ) in  
 134 (d) surface water and (e) sediment DOM vs (f) representative PyOM. Higher values on the x-axis  
 135 ( $\Delta G_{CoX}$ ) represent less energetically favorable molecules, while higher values on the y-axis ( $\lambda$ )  
 136 represent higher rates of catabolism relative to anabolism. Colors for all panels indicate inferred  
 137 chemical classes.





138

139 **Figure 2.** Predicted CUE and metabolism of DOM and PyOM. (a) shows the probability density  
 140 function (PDF) of CUE in surface water DOM (blue), sediment (black) DOM, and PyOM (red).  
 141 The PDF reflects the relative likelihood that value of a random sample drawn from a particular  
 142 group would equal the value on the x-axis. (b) depicts CUE of surface water DOM, sediment  
 143 DOM, and PyOM by inferred chemical class. (c) depicts the predicted metabolism of surface  
 144 water DOM, sediment DOM, and PyOM under no elemental limitations (no fill), C-limitation  
 145 (light grey fill), and oxygen-limitation (dark grey fill). (d-f) represent metabolic rates for PyOM  
 146 molecules with known chemical classes under (d) no elemental limitations, (e) C-limitation, and  
 147 (f) oxygen-limitation. For boxplots, median values are denoted by a bar, hinges correspond to the  
 148 first and third quartiles, and whiskers extend to the largest/smallest value no further than 1.5  
 149 times the inter-quartile range from the hinge. Data beyond the end of the whiskers are plotted  
 150 individually.

151

152

153            Interestingly, the large overlaps in predicted CUE across PyOM and DOM pools suggests  
154 that PyOM decomposition in rivers could emit proportionally similar amounts of CO<sub>2</sub> to natural  
155 DOM (Figure 2a-b). CUE is used in many microbially-explicit decomposition models to  
156 constrain organic matter bioavailability (reviewed in Graham and Hofmockel, 2022). Therefore,  
157 predicted CUE offers a path for assimilating PyOM in microbially-explicit models. Such an  
158 approach could be used to directly evaluate the impact of PyOM on global C cycles, and help  
159 lead to better incorporation of PyOM impacts in models (Santin et al., 2020). We specifically  
160 highlight lower CUE of amino sugar-like, protein-like, and lignin-like PyOM in comparison to  
161 corresponding classes of natural DOM, pointing to potential shifts needed in the representation  
162 of these classes in process-based models in ecosystems impacted by wildfires.

163

## 164 *2.2 Inferred Metabolism of Pyrogenic Organic Matter.*

165            Predicted rates of PyOM metabolism were also similar to DOM pools (Figure 2c),  
166 reinforcing comparable bioavailability. Elemental limitations had a ~6-fold larger impact on  
167 predicted metabolic rates than changes in compound type (e.g., surface water DOM, sediment  
168 DOM, and PyOM). Metabolic rates averaged across all compound types were 0.965, 0.179, and  
169 0.136 mol O/mol C biomass produced for unlimited, C-limited, and oxygen-limited conditions  
170 respectively. In contrast when metabolic rates were averaged across elemental limitations within  
171 each compound group, we found that surface water DOM, sediment DOM, and PyOM had  
172 nearly identical mean predicted rates of metabolism (ANOVA  $P = 0.14$   $F = 1.922$ , all pairwise  
173 comparisons  $P > 0.10$ ). While metabolic rates within a given limitation scenario were statistically  
174 different across compound types (ANOVA, all  $P < 0.001$ ,  $F = 197.4$ ,  $145.9$ , and  $265.9$ ), effect  
175 sizes were low in comparison to changes across elemental limitations. Still, PyOM displayed

176 statistically elevated rates of metabolism under C- and oxygen-limitation versus both surface  
177 water and sediment DOM (Tukey HSD,  $P < 0.001$ ). Given that riverbed sediments can reach  
178 anoxia within millimeters of the sediment-water interface, elevated rates of PyOM metabolism  
179 under oxygen limitation may be ecologically relevant at ecosystem scales.

180         When considering the impact of elemental limitations on PyOM metabolism, predicted  
181 rates were strongly inhibited under low C and oxygen conditions. Predicted PyOM metabolism  
182 was over 5 times lower when C or oxygen was scarce. Low decomposition rates under C and  
183 oxygen limitation could be one reason for the observed persistence of PyOM in depositional  
184 features that tend to be anoxic. Still, it is worthwhile to note that metabolism of all PyOM classes  
185 under low C or oxygen was predicted to be substantially slower than without elemental  
186 limitations, indicating PyOM compounds may both actively cycle in well-oxygenated surface  
187 waters with fresh C inputs and persist over long periods of time in O<sub>2</sub>-limited sediments.

188         When examining PyOM with known chemistries (e.g., instead of inferred chemical  
189 classes from FTICR-MS), anhydrosugars, phenols, and polycyclic aromatic hydrocarbons  
190 (PAHs) were less negatively impacted by oxygen limitation than any other group (Figure 2 d-f).  
191 Previous work has demonstrated that microorganisms are capable of decomposing chemically  
192 complex organic molecules, such as long-chained and/or aromatic hydrocarbons under low  
193 oxygen availability (Bushnell and Haas, 1941; Pozdnyakova, 2012; Rabus et al., 2016; Coates et  
194 al., 1997). Similar microbial metabolic pathways may also be capable of degrading pyrogenically  
195 modified phenols, and PAHs in natural settings and could be investigated with future laboratory  
196 work.

197         Notably, our work also supports the notion that black nitrogen could be more bioavailable  
198 than other PyOM pools. While we only examined one class of PyOM molecules containing

199 nitrogen in detail (n-alkylnitriles), it had among the highest predicted CUE and metabolic rates.  
200 We posit this may be due, in part, to its chemical structure that includes pyrrole-type moieties,  
201 which are relatively biodegradable (Knicker, 2010; De La Rosa and Knicker, 2011).

202

### 203 *2.3 Correspondence to Empirical Investigations.*

204 While the substrate-explicit modelling approach used here has been validated in natural  
205 settings, its underlying assumptions preclude accounting for DOM structure and size, abiotic  
206 reactions, and chemical complexation with minerals and particulates. Some aspects of model  
207 predictions are inconsistent with experimental evidence, highlighting the role of laboratory  
208 studies in evaluating PyOM bioavailability.

209 For instance, n-alkenes and related compounds tended to have high modelled  
210 bioavailability despite being relatively stable in the environment (Wiesenberg et al., 2004;  
211 Smittenberg et al., 2004). These compounds are characterized by carbon-carbon double bonds,  
212 which are not considered by the model and may decrease bioavailability. However, n-alkanes  
213 generated through combustion tend to have reduced chain length in comparison to their unburned  
214 counterparts (Knicker et al., 2013), and thus may be relatively bioavailable compared to  
215 unburned n-alkanes. Additionally, we note that previous work has shown fast degradation of  
216 combustion-derived lipids in soils (Knicker et al., 2013); as well as high n-alkene metabolism in  
217 anaerobic sediments and high lipid reactivities at the sediment-water interface (Grossi et al.,  
218 2008; Wilkes et al., 2016; Yongdong et al., 2015; Mbadinga et al., 2011; Canuel and Martens,  
219 1996). While work on aerobic n-alkene metabolism is limited, the comparative bioavailability of  
220 n-alkenes and known degradation pathways suggests that sediment microbiomes may metabolize  
221 them as part of natural biogeochemical cycles.

222 Another notable discrepancy is the low potential bioavailability of anhydrosugars when  
223 compared to other PyOM compounds. Experimentally, anhydrosugars are highly bioavailable in  
224 oxic conditions, with a half-life of less than seven days (Norwood et al., 2013). The model may  
225 therefore not adequately account for some enzyme-catalyzed reactions such as levoglucosan  
226 kinase or levoglucosan dehydrogenase that may be common enzymes in aquatic microorganisms  
227 (Bacik and Jarboe, 2016; Suciu et al., 2019).

228 Because of these nuances, the analysis presented here is best used as bounding estimates  
229 for experimental validation and as a holistic comparison to DOM bioavailability. Still, the span  
230 of compounds investigated here, and their comparison to DOM pools, provides a breadth of  
231 investigation that is unfeasible without model-based approaches.

232

### 233 **3 Conclusions**

234 Our work supports the hypothesis that PyOM may have similar overall bioavailability as  
235 compared to natural sources of DOM and provides a foundation for targeted experiments  
236 investigating specific components of the PyOM continuum. Globally intensifying wildfires are  
237 increasing the production of PyOM with potential implications for water supplies, which are  
238 critical for domestic, industrial, agricultural, and ecological needs. Yet, many fundamental  
239 questions such as “how much” PyOM exists in ecosystems, “how fast” it cycles, and “how old”  
240 it is remain largely unknown (Abiven and Santin, 2019). Our work provides the first  
241 comprehensive computational assessment of the potential bioavailabilities of various PyOM  
242 chemistries in comparison to natural DOM pools. The comparable potential bioavailability to  
243 DOM revealed that PyOM may be actively transformed in aquatic ecosystems and may be an  
244 increasing source of C emissions to the atmosphere as the prevalence of wildfires increases.

245

**246 Code and Data Availability**

247 Code is available at: <https://github.com/hyunseobsong/lambda>. Data describing DOM pool  
248 chemistry are published as a data package (Goldman et al., 2020) (available at:  
249 doi:10.15485/1729719) and are discussed in more detail by Garayburu-Caruso et al. (2020a).

250

**251 Author Contributions**

252 EBG conceived of the manuscript and was responsible for writing the manuscript and generating  
253 all figures. HSS performed all modelling. SG determined PyOM compounds for modelling based  
254 on extensive literature review, with guidance from AMP. VGC and JCS contributed data and  
255 insight on DOM pool chemistry. All authors contributed to revisions.

256

**257 Competing Interests**

258 The authors declare that they have no conflict of interest.

259

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## 270 **APPENDIX. Materials and methods**

271 An extended version of our methods is available in the Supporting Information.

272 To assess the potential bioavailability of PyOM, we searched primary literature for  
273 representative compounds of the PyOM continuum. Specifically, we targeted characteristic  
274 organic compounds from controlled burns of various fuel types representing a range of moisture,  
275 temperature, and oxygen conditions (Table S1). The chosen compounds focused on biomass  
276 burning alteration products derived from plant charcoals, biochars, smoke, burnt soil and their  
277 leachates, which are often used to characterize PyOM in the environment. This included  
278 compounds such as theoretical black carbon-like compounds (defined here as condensed  
279 aromatic core structures polysubstituted with O-containing functionalities (Wagner et al., 2017),  
280 anhydrosugars, and polycyclic aromatic hydrocarbons (PAHs). The list also included compounds  
281 created and/or transformed from biomass burning, such as those derived from biopolymers like  
282 lignin (e.g., methoxyphenols), waxes (e.g., n-alkenes from thermal dehydration of n-alkanols),  
283 and resins (e.g., thermally oxidized diterpenoids) (Oros and Simoneit, 2001b, a). We also include  
284 molecular formula assigned via FTICR-MS spectra collected on biochar leachates to increase  
285 data type comparability with common methods used in natural organic matter studies. The  
286 selected set of compounds spans the chemical continuum of PyOM but is not intended to be  
287 exhaustive. In total, our literature search for PyOM chemistries yielded 16,971 unique molecular  
288 formulae. When a formula was found in more than one study or material type, this was noted in  
289 Table S1, however only unique formulae were considered for the model.

290 After generating a set of representative compounds, we used a substrate-explicit  
291 modelling framework developed by Song et al. (2020) to characterize the potential  
292 bioavailability of each compound and predict its rate of decomposition. The model uses  
293 molecular formulae to predict energetic content, metabolic efficiency, and rates of aerobic  
294 metabolism, while it does not account for structural components of organic molecules (e.g.,  
295 double bonds, folding patterns, cross-linkages). This enables flexibility in application to high-  
296 throughput mass spectrometry techniques that yield chemical formulae but not structural  
297 information (e.g., FTICR-MS) for comparison to environmental DOM. Despite its limitations,  
298 the substrate-explicit model used here has proven useful in linking DOM composition to aerobic  
299 metabolism in natural environments (Song et al., 2020; Graham et al., 2017; Garayburu-Caruso  
300 et al., 2020b), and its structure is consistent with Harvey et al. (2016) who argued for the  
301 importance of thermodynamic estimates of PyOM bioavailability that underlie this model. It was  
302 chosen to allow for comparison of PyOM to the most comprehensive assessment of global  
303 aquatic DOM pools to date (Garayburu-Caruso et al., 2020a).

304 Briefly, the substrate-explicit model uses the elemental stoichiometry of organic  
305 molecules, based on molecular formulae, to predict the number of catabolic reactions that must  
306 occur to provide the energy required for the synthesis of one mole of biomass carbon. This  
307 quantity is described by the parameter lambda ( $\lambda$ ) in which lower  $\lambda$  values denote more efficient  
308 energetics of catabolism in producing biomass through anabolism. The model also predicts the  
309 Gibbs free energy of C oxidation ( $\Delta G_{Cox}$ ), under standard conditions with a modification to pH 7  
310 adjusted from LaRowe and Van Cappellen (2011) by Song et al. (2020), as well as C use  
311 efficiency (CUE) as defined by Saifuddin et al. (2019). Lower  $\Delta G_{Cox}$  denotes higher  
312 thermodynamic favorability in an electron donor half reaction associated with organic matter,



313 and higher CUE reflects more C assimilated into biomass per unit C respired. We also predicted  
314 the rate of aerobic metabolism (as oxygen consumed per mol-C biomass produced) under three  
315 scenarios commonly observed in aquatic ecosystems: (a) C-limitation, (b) oxygen ( $O_2$ )  
316 limitation, and (c) both C and  $O_2$ -limitation. For more details of the substrate-explicit modelling  
317 approach used, please see Song et al. (2020). Each metric ( $\lambda$ ,  $\Delta G_{Cox}$ , CUE, metabolic rates)  
318 denotes a different aspect of potential bioavailability. Though the relative magnitude of the  
319 metrics in comparison to each other will vary based on the specific stoichiometry of a molecule,  
320 highly bioavailable compounds are indicated by low  $\lambda$  and  $\Delta G_{Cox}$  coinciding with high CUE and  
321 metabolic rates.

322 Three sets of organic molecules were used as model inputs: measured global dissolved  
323 (1) surface water and (2) sediment DOM pools, extracted in  $H_2O$  and analyzed by FTICR-MS as  
324 per Garayburu-Caruso et al. (2020a); and (3) literature-derived PyOM compounds as described  
325 above. Inputs to the model from the PyOM compounds were unique molecular formulae,  
326 grouped in subsequent analysis by their corresponding compound classes (Table S1). If one  
327 molecular formula was represented by several PyOM compounds (e.g.,  $C_{10}H_{16}O_2$ , which  
328 corresponds to the sesquiterpenoid cis-Thujan-10-oic acid and 3-, 4- substituted methylcatechol  
329 phenols), we assigned multiple compound classes to that molecular formula. Surface water and  
330 sediment DOM pools were filtered to compounds occurring in 95% of samples to yield a dataset  
331 of globally ubiquitous DOM (surface water  $n = 811$  molecules, sediment  $n = 398$  molecules).  
332 Formulae assignment and inferred chemical classes via van Krevelen diagrams in DOM pools  
333 are described by Garayburu-Caruso et al. (2020a). As a whole, 16,332 compounds were found  
334 only in PyOM, 197 were found only in surface water DOM, and 167 were found only in  
335 sediment DOM. We compared modelling outputs from representative PyOM to outputs of

336 ubiquitous DOM pools to infer relative bioavailability using ANOVA and Tukey HSD statistical  
337 tests with R software v 4.1.0 (R Core Team, 2021). Visualizations were produced using base R  
338 and ‘ggplot2’ (Wickham, 2011). All model outputs are available in Tables S2–S4.

339

340 **REFERENCES**

- 341 Abiven, S. and Santin, C.: From Fires to Oceans: Dynamics of Fire-Derived Organic Matter in  
342 Terrestrial and Aquatic Ecosystems, *Frontiers in Earth Science*, 7, 31, 2019.
- 343 Abney, R. B., Kuhn, T. J., Chow, A., Hockaday, W., Fogel, M. L., and Berhe, A. A.: Pyrogenic  
344 carbon erosion after the Rim Fire, Yosemite National Park: The role of burn severity and slope,  
345 *Journal of Geophysical Research: Biogeosciences*, 124, 432-449, 2019.
- 346 Arndt, S., Jørgensen, B. B., LaRowe, D. E., Middelburg, J., Pancost, R., and Regnier, P.:  
347 Quantifying the degradation of organic matter in marine sediments: a review and synthesis,  
348 *Earth-science reviews*, 123, 53-86, 2013.
- 349 Bacik, J. P. and Jarboe, L. R.: Bioconversion of anhydrosugars: emerging concepts and strategies,  
350 *IUBMB life*, 68, 700-708, 2016.
- 351 Bigio, E., Swetnam, T. W., and Baisan, C. H.: A comparison and integration of tree-ring and  
352 alluvial records of fire history at the Missionary Ridge Fire, Durango, Colorado, USA, *The*  
353 *Holocene*, 20, 1047-1061, 2010.
- 354 Bladon, K. D., Emelko, M. B., Silins, U., and Stone, M.: Wildfire and the future of water supply,  
355 2014.
- 356 Bowman, D. M., Kolden, C. A., Abatzoglou, J. T., Johnston, F. H., van der Werf, G. R., and  
357 Flannigan, M.: Vegetation fires in the Anthropocene, *Nature Reviews Earth & Environment*, 1,  
358 500-515, 2020.
- 359 Bushnell, L. and Haas, H.: The utilization of certain hydrocarbons by microorganisms, *Journal of*  
360 *Bacteriology*, 41, 653, 1941.
- 361 Canuel, E. A. and Martens, C. S.: Reactivity of recently deposited organic matter: Degradation of  
362 lipid compounds near the sediment-water interface, *Geochimica et cosmochimica acta*, 60,  
363 1793-1806, 1996.
- 364 Cawley, K. M., Hohner, A. K., McKee, G. A., Borch, T., Omur-Ozbek, P., Oropeza, J., and Rosario-  
365 Ortiz, F. L.: Characterization and spatial distribution of particulate and soluble carbon and  
366 nitrogen from wildfire-impacted sediments, *Journal of soils and sediments*, 18, 1314-1326,  
367 2018.
- 368 Coates, J. D., Woodward, J., Allen, J., Philp, P., and Lovley, D. R.: Anaerobic degradation of  
369 polycyclic aromatic hydrocarbons and alkanes in petroleum-contaminated marine harbor  
370 sediments, *Applied and Environmental Microbiology*, 63, 3589-3593, 1997.
- 371 de la Rosa, J. M. and Knicker, H.: Bioavailability of N released from N-rich pyrogenic organic  
372 matter: an incubation study, *Soil Biology and Biochemistry*, 43, 2368-2373, 2011.
- 373 Dittmar, T., De Rezende, C. E., Manecki, M., Niggemann, J., Ovalle, A. R. C., Stubbins, A., and  
374 Bernardes, M. C.: Continuous flux of dissolved black carbon from a vanished tropical forest  
375 biome, *Nature Geoscience*, 5, 618-622, 2012.
- 376 Domeignoz-Horta, L. A., Pold, G., Liu, X.-J. A., Frey, S. D., Melillo, J. M., and DeAngelis, K. M.:  
377 Microbial diversity drives carbon use efficiency in a model soil, *Nature communications*, 11, 1-  
378 10, 2020.
- 379 Elliott, J. G. and Parker, R.: Developing a post-fire flood chronology and recurrence probability  
380 from alluvial stratigraphy in the Buffalo Creek watershed, Colorado, USA, *Hydrological*  
381 *Processes*, 15, 3039-3051, 2001.

382 Emelko, M. B., Silins, U., Bladon, K. D., and Stone, M.: Implications of land disturbance on  
383 drinking water treatability in a changing climate: Demonstrating the need for “source water  
384 supply and protection” strategies, *Water research*, 45, 461-472, 2011.

385 Flannigan, M. D., Krawchuk, M. A., de Groot, W. J., Wotton, B. M., and Gowman, L. M.:  
386 Implications of changing climate for global wildland fire, *International journal of wildland fire*,  
387 18, 483-507, 2009.

388 Garayburu-Caruso, V. A., Danczak, R. E., Stegen, J. C., Renteria, L., McCall, M., Goldman, A. E.,  
389 Chu, R. K., Toyoda, J., Resch, C. T., and Torgeson, J. M.: Using Community Science to Reveal the  
390 Global Chemogeography of River Metabolomes, *Metabolites*, 10, 518, 2020a.

391 Garayburu-Caruso, V. A., Stegen, J. C., Song, H.-S., Renteria, L., Wells, J., Garcia, W., Resch, C. T.,  
392 Goldman, A. E., Chu, R. K., Toyoda, J., and Graham, E. B.: Carbon limitation leads to  
393 thermodynamic regulation of aerobic metabolism, *Environmental Science & Technology*  
394 *Letters*, 2020b.

395 Goldman, A. E., Chu, R. K., Danczak, R. E., Daly, R. A., Fansler, S., Garayburu-Caruso, V. A.,  
396 Graham, E. B., McCall, M. L., Ren, H., and Renteria, L.: WHONDRS Summer 2019 Sampling  
397 Campaign: Global River Corridor Sediment FTICR-MS, NPOC, and Aerobic Respiration,  
398 *Environmental System Science Data Infrastructure for a Virtual Ecosystem ...*, 2020.

399 Graham, E. B. and Hofmockel, K.: Ecological stoichiometry as a foundation for omics-enabled  
400 biogeochemical models of soil organic matter decomposition, *Biogeochemistry*, 157, 31-50,  
401 2022.

402 Graham, E. B., Tfaily, M. M., Crump, A. R., Goldman, A. E., Bramer, L. M., Arntzen, E., Romero,  
403 E., Resch, C. T., Kennedy, D. W., and Stegen, J. C.: Carbon inputs from riparian vegetation limit  
404 oxidation of physically bound organic carbon via biochemical and thermodynamic processes,  
405 *Journal of Geophysical Research: Biogeosciences*, 122, 3188-3205, 2017.

406 Grossi, V., Cravo-Laureau, C., Guyoneaud, R., Ranchou-Peyruse, A., and Hirschler-Réa, A.:  
407 Metabolism of n-alkanes and n-alkenes by anaerobic bacteria: A summary, *Organic*  
408 *Geochemistry*, 39, 1197-1203, 2008.

409 Harvey, O. R., Myers-Pigg, A. N., Kuo, L.-J., Singh, B. P., Kuehn, K. A., and Louchouart, P.:  
410 Discrimination in degradability of soil pyrogenic organic matter follows a return-on-energy-  
411 investment principle, *Environmental science & technology*, 50, 8578-8585, 2016.

412 Hockaday, W. C., Grannas, A. M., Kim, S., and Hatcher, P. G.: The transformation and mobility of  
413 charcoal in a fire-impacted watershed, *Geochimica et Cosmochimica Acta*, 71, 3432-3445, 2007.

414 Hohner, A. K., Terry, L. G., Townsend, E. B., Summers, R. S., and Rosario-Ortiz, F. L.: Water  
415 treatment process evaluation of wildfire-affected sediment leachates, *Environmental Science:*  
416 *Water Research & Technology*, 3, 352-365, 2017.

417 Jaffé, R., Ding, Y., Niggemann, J., Vähätalo, A. V., Stubbins, A., Spencer, R. G., Campbell, J., and  
418 Dittmar, T.: Global charcoal mobilization from soils via dissolution and riverine transport to the  
419 oceans, *Science*, 340, 345-347, 2013.

420 Knicker, H.: “Black nitrogen”—an important fraction in determining the recalcitrance of charcoal,  
421 *Organic Geochemistry*, 41, 947-950, 2010.

422 Knicker, H., Hilscher, A., De la Rosa, J., González-Pérez, J. A., and González-Vila, F. J.:  
423 Modification of biomarkers in pyrogenic organic matter during the initial phase of charcoal  
424 biodegradation in soils, *Geoderma*, 197, 43-50, 2013.

- 425 Kuzyakov, Y., Bogomolova, I., and Glaser, B.: Biochar stability in soil: decomposition during eight  
426 years and transformation as assessed by compound-specific <sup>14</sup>C analysis, *Soil Biology and*  
427 *Biochemistry*, 70, 229-236, 2014.
- 428 LaRowe, D. E. and Van Cappellen, P.: Degradation of natural organic matter: a thermodynamic  
429 analysis, *Geochimica et Cosmochimica Acta*, 75, 2030-2042, 2011.
- 430 Masiello, C. A.: New directions in black carbon organic geochemistry, *Marine Chemistry*, 92,  
431 201-213, 2004.
- 432 Mbadanga, S. M., Wang, L.-Y., Zhou, L., Liu, J.-F., Gu, J.-D., and Mu, B.-Z.: Microbial communities  
433 involved in anaerobic degradation of alkanes, *International Biodeterioration & Biodegradation*,  
434 65, 1-13, 2011.
- 435 Meyer, G. A. and Wells, S. G.: Fire-related sedimentation events on alluvial fans, Yellowstone  
436 National Park, USA, *Journal of Sedimentary Research*, 67, 776-791, 1997.
- 437 Mia, S., Dijkstra, F. A., and Singh, B.: Long-term aging of biochar: a molecular understanding  
438 with agricultural and environmental implications, *Advances in agronomy*, 141, 1-51, 2017.
- 439 Myers-Pigg, A. N., Louchouart, P., Amon, R. M., Prokushkin, A., Pierce, K., and Rubtsov, A.:  
440 Labile pyrogenic dissolved organic carbon in major Siberian Arctic rivers: Implications for  
441 wildfire-stream metabolic linkages, *Geophysical Research Letters*, 42, 377-385, 2015.
- 442 Neary, D. G., Ryan, K. C., and DeBano, L. F.: Wildland fire in ecosystems: effects of fire on soils  
443 and water, Gen. Tech. Rep. RMRS-GTR-42-vol. 4. Ogden, UT: US Department of Agriculture,  
444 Forest Service, Rocky Mountain Research Station. 250 p., 42, 2005.
- 445 Norwood, M. J., Louchouart, P., Kuo, L.-J., and Harvey, O. R.: Characterization and  
446 biodegradation of water-soluble biomarkers and organic carbon extracted from low  
447 temperature chars, *Organic Geochemistry*, 56, 111-119, 2013.
- 448 Oros, D. R. and Simoneit, B. R.: Identification and emission factors of molecular tracers in  
449 organic aerosols from biomass burning Part 2. Deciduous trees, *Applied Geochemistry*, 16,  
450 1545-1565, 2001a.
- 451 Oros, D. R. and Simoneit, B. R.: Identification and emission factors of molecular tracers in  
452 organic aerosols from biomass burning Part 1. Temperate climate conifers, *Applied*  
453 *Geochemistry*, 16, 1513-1544, 2001b.
- 454 Pierce, J. L., Meyer, G. A., and Jull, A. T.: Fire-induced erosion and millennial-scale climate  
455 change in northern ponderosa pine forests, *Nature*, 432, 87, 2004.
- 456 Pold, G., Domeignoz-Horta, L. A., Morrison, E. W., Frey, S. D., Sistla, S. A., and DeAngelis, K. M.:  
457 Carbon use efficiency and its temperature sensitivity covary in soil bacteria, *Mbio*, 11, 2020.
- 458 Pozdnyakova, N. N.: Involvement of the ligninolytic system of white-rot and litter-decomposing  
459 fungi in the degradation of polycyclic aromatic hydrocarbons, *Biotechnology research*  
460 *international*, 2012, 2012.
- 461 Rabus, R., Boll, M., Heider, J., Meckenstock, R. U., Buckel, W., Einsle, O., Ermler, U., Golding, B.  
462 T., Gunsalus, R. P., and Kroneck, P. M.: Anaerobic microbial degradation of hydrocarbons: from  
463 enzymatic reactions to the environment, *Journal of molecular microbiology and biotechnology*,  
464 26, 5-28, 2016.
- 465 Randerson, J. T., Liu, H., Flanner, M. G., Chambers, S. D., Jin, Y., Hess, P. G., Pfister, G., Mack, M.,  
466 Treseder, K., and Welp, L.: The impact of boreal forest fire on climate warming, *science*, 314,  
467 1130-1132, 2006.

- 468 Saifuddin, M., Bhatnagar, J. M., Segrè, D., and Finzi, A. C.: Microbial carbon use efficiency  
469 predicted from genome-scale metabolic models, *Nature communications*, 10, 1-10, 2019.
- 470 Santin, C., Doerr, S. H., Jones, M. W., Merino, A., Warneke, C., and Roberts, J. M.: The relevance  
471 of pyrogenic carbon for carbon budgets from fires: insights from the FIREX experiment, *Global*  
472 *Biogeochemical Cycles*, 34, e2020GB006647, 2020.
- 473 Santin, C., Doerr, S. H., Kane, E. S., Masiello, C. A., Ohlson, M., de la Rosa, J. M., Preston, C. M.,  
474 and Dittmar, T.: Towards a global assessment of pyrogenic carbon from vegetation fires, *Global*  
475 *Change Biology*, 22, 76-91, 2016.
- 476 Shakesby, R. and Doerr, S.: Wildfire as a hydrological and geomorphological agent, *Earth-*  
477 *Science Reviews*, 74, 269-307, 2006.
- 478 Smittenberg, R. H., Hopmans, E. C., Schouten, S., Hayes, J. M., Eglinton, T. I., and Sinninghe  
479 Damsté, J.: Compound-specific radiocarbon dating of the varved Holocene sedimentary record  
480 of Saanich Inlet, Canada, *Paleoceanography*, 19, 2004.
- 481 Sohi, S. P., Krull, E., Lopez-Capel, E., and Bol, R.: A review of biochar and its use and function in  
482 soil, *Advances in agronomy*, 105, 47-82, 2010.
- 483 Song, H.-S., Stegen, J. C., Graham, E. B., Lee, J.-Y., Garayburu-Caruso, V., Nelson, W. C., Chen, X.,  
484 Moulton, J. D., and Scheibe, T. D.: Representing Organic Matter Thermodynamics in  
485 Biogeochemical Reactions via Substrate-Explicit Modeling, *Frontiers in Microbiology*,  
486 <https://doi.org/10.3389/fmicb.2020.531756>, 2020.
- 487 Suciu, L. G., Masiello, C. A., and Griffin, R. J.: Anhydrosugars as tracers in the Earth system,  
488 *Biogeochemistry*, 146, 209-256, 2019.
- 489 Team, R. C.: R: A language and environment for statistical computing, R Foundation for  
490 Statistical Computing [code], 2021.
- 491 Verma, S. and Jayakumar, S.: Impact of forest fire on physical, chemical and biological  
492 properties of soil: A review, *Proceedings of the International Academy of Ecology and*  
493 *Environmental Sciences*, 2, 168, 2012.
- 494 Wagner, S., Ding, Y., and Jaffé, R.: A new perspective on the apparent solubility of dissolved  
495 black carbon, *Frontiers in Earth Science*, 5, 75, 2017.
- 496 Wagner, S., Dittmar, T., and Jaffé, R.: Molecular characterization of dissolved black nitrogen via  
497 electrospray ionization Fourier transform ion cyclotron resonance mass spectrometry, *Organic*  
498 *Geochemistry*, 79, 21-30, 2015.
- 499 Wagner, S., Jaffé, R., and Stubbins, A.: Dissolved black carbon in aquatic ecosystems, *Limnology*  
500 *and Oceanography Letters*, 3, 168-185, 2018.
- 501 Wickham, H.: ggplot2, *Wiley interdisciplinary reviews: computational statistics*, 3, 180-185,  
502 2011.
- 503 Wiesenberg, G. L., Schwarzbauer, J., Schmidt, M. W., and Schwark, L.: Source and turnover of  
504 organic matter in agricultural soils derived from n-alkane/n-carboxylic acid compositions and C-  
505 isotope signatures, *Organic Geochemistry*, 35, 1371-1393, 2004.
- 506 Wilkes, H., Buckel, W., Golding, B. T., and Rabus, R.: Metabolism of hydrocarbons in n-alkane-  
507 utilizing anaerobic bacteria, *Journal of molecular microbiology and biotechnology*, 26, 138-151,  
508 2016.
- 509 Wozniak, A. S., Goranov, A. I., Mitra, S., Bostick, K. W., Zimmerman, A. R., Schlesinger, D. R.,  
510 Myneni, S., and Hatcher, P. G.: Molecular heterogeneity in pyrogenic dissolved organic matter  
511 from a thermal series of oak and grass chars, *Organic Geochemistry*, 148, 104065, 2020.

- 512 Yongdong, Z., Yaling, S., Zhengwen, L., Xiangchao, C., Jinlei, Y., Xiaodan, D., and Miao, J.: Long-  
513 chain n-alkenes in recent sediment of Lake Lugu (SW China) and their ecological implications,  
514 *Limnologica*, 52, 30-40, 2015.
- 515 Zimmerman, A. R. and Mitra, S.: Trial by fire: on the terminology and methods used in  
516 pyrogenic organic carbon research, *Frontiers in Earth Science*, 5, 95, 2017.
- 517 Zimmerman, A. R. and Ouyang, L.: Priming of pyrogenic C (biochar) mineralization by dissolved  
518 organic matter and vice versa, *Soil Biology and Biochemistry*, 130, 105-112, 2019.
- 519