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4	Summer temperature - but not growing season length - influences radial growth of Salix arctica in coastal Arctic
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## 25 Abstract

26 Arctic climate change is leading to an advance of plant phenology (the timing of life history events) with uncertain 27 impacts on tundra ecosystems. Although the lengthening of the growing season is thought to lead to increased plant 28 growth, we have few studies of how plant phenology change is altering tundra plant productivity. Here, we test the 29 correspondence between 14 years of Salix arctica phenology data and radial growth on Qikiqtaruk – Herschel Island, 30 Yukon Territory, Canada. We analysed stems from 28 individuals using dendroecology and linear mixed-effect models 31 to test the statistical power of growing season length and climate variables to individually predict radial growth. We 32 found that summer temperature best explained annual variation in radial growth. We found no strong evidence that leaf 33 emergence date, earlier leaf senescence date or total growing season length had any direct or lagged effects on radial 34 growth. Radial growth was also not explained by interannual variation in precipitation, MODIS surface greenness 35 (NDVI), or sea ice concentration. Our results demonstrate that at this site, for the widely-distributed species S. arctica, 36 temperature – but not growing season length – influences radial growth. These findings challenge the assumption that 37 advancing phenology and longer growing seasons will increase the productivity of all plant species in Arctic tundra 38 ecosystems.

39

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52

# 53 Author Contributions

54 JB and IMS conceived and designed research. SAB collected the samples and JB conducted lab work, with all authors

55 contributing to the analysis and additional data. The main script was adapted by JB and IMS from code written by SAB.

56 JB and IMS wrote the manuscript, with JJA creating Figure 1. All authors read and approved the manuscript.

# 58 Data Availability Statement

- 59 The code and data used for this study and generated during our analyses are available at the following GitHub
- 60 repository: <u>https://github.com/ShrubHub/ShrubRingPhenoHub</u>

## 61 Introduction

62 The Arctic is warming three to four times faster than the rest of the planet (Meredith et al. 2019; You et al. 2021) and 63 tundra plant communities are particularly sensitive to that warming (Elmendorf et al. 2015; Bjorkman et al. 2020). 64 Climate change is resulting in a longer snow- and ice-free season, potentially facilitating longer growing seasons 65 (Cleland et al. 2007; Khorsand Rosa et al. 2015; Prevéy et al. 2021; Frei and Henry 2021). Concurrent with these 66 changes, shifts in distribution and abundance (Sturm et al. 2001; Elmendorf et al. 2012a), biomass (Hudson and Henry 67 2009), and phenology (timing of life history events) (Oberbauer et al. 2013) have been observed for species across the 68 tundra. Previous research assumes that the altered phenology will correspond directly with increased growth of tundra 69 plants (Myneni et al. 1997; Piao et al. 2007; Ernakovich et al. 2014; Park et al. 2016; Arndt et al. 2019; Kim et al. 70 2021). Arctic spectral greening trends from satellite vegetation indices, such as the Normalised Differential Vegetation 71 Index (NDVI), are used as proxy metrics of tundra plant phenology (Piao et al. 2007; Zeng et al. 2013; Park et al. 2016) 72 and plant productivity (Myneni et al. 1997; Kim et al. 2021). However, remote sensing studies of tundra phenology may 73 not be capturing on-the-ground plant phenological and growth dynamics and may instead be influenced by other land-74 surface changes such as community compositional shifts and snow cover (Helman 2018) particularly in Arctic systems 75 (Arndt et al. 2019; Myers-Smith et al. 2020; Cao et al. 2020). Satellite (Myers-Smith et al. 2020) and in situ (Oberbauer 76 et al. 2013) studies indicate widespread but variable phenological shifts in the Arctic. Remote sensing studies have 77 linked phenology change with increased plant productivity in tundra ecosystems (Myneni et al. 1997; Park et al. 2016; 78 Kim et al. 2021); however, *in situ* studies directly linking plant phenology change to plant growth are rare.

79

80 Plant phenology is changing throughout the tundra (Myers-Smith et al. 2019; Bjorkman et al. 2020; Prevéy et al. 2021), 81 but the consequences on plant growth remain unclear. Phenology defines the bounds for plant activity, including 82 photosynthesis, and has shifted around the Arctic due to warming (Assmann et al. 2019; Myers-Smith et al. 2020). The 83 snow-free season across the Arctic has extended by two to four days per decade of warming (Piao et al. 2007; 84 Barichivich et al. 2013; Park et al. 2016; Myers-Smith et al. 2019). Seasons are starting earlier and finishing earlier or 85 later depending on the location and study metrics investigated (Piao et al. 2007; Zeng et al. 2011, 2013; Keenan and 86 Richardson 2015; Park et al. 2016; Myers-Smith et al. 2019). Two key points in deciduous species' phenology are leaf 87 emergence and leaf senescence; the time between being the entire growing season. Leaf emergence and senescence are 88 both shifting across the Arctic, leading to a longer, earlier growing season at many sites, though changing phenology is 89 not uniform across sites or species (Oberbauer et al. 2013; Assmann et al. 2019; Myers-Smith et al. 2020). Earlier leaf 90 emergence is associated with earlier snowmelt (Assmann et al. 2019; Myers-Smith et al. 2019) and declining sea ice 91 (Post et al. 2009; Bhatt et al. 2010; Kerby and Post 2013), though some studies have identified trends toward later 92 phenology in some species and locations usually aligning with later snowmelt (Wipf and Rixen 2010; Bjorkman et al.

2015). Earlier leaf emergence may expose individuals to late spring frost events (Wheeler et al. 2015) or other harsh
conditions. Early senescence through deterministic leaf age (Oberbauer et al. 2013; Keenan and Richardson 2015),
nutrient availability (Lim et al. 2007), or photoperiod (Arft et al. 1999) may also undermine any growth benefits of
earlier leaf emergence. Whether plants can take advantage of an extended growing season to increase productivity and
accumulate biomass is therefore uncertain.

98

99 At mid-latitudes of the Arctic, shrub growth can be particularly sensitive to climate (Myers-Smith et al. 2015a) and 100 there is ground-based and satellite evidence for rapid shrub expansion in the region (Fraser et al. 2011; Moffat et al. 101 2016; Myers-Smith et al. 2019). Dendroecology allows for the exploration of the growth history of shrubs based on the 102 width of rings formed during seasonal woody tissue deposition (Myers-Smith et al. 2015b). Individual annual growth 103 ring chronologies can be compared with environmental variables to reveal the climate sensitivity of radial growth over 104 time. Through dendroecology, we can directly observe how changing conditions affect shrub growth, validating 105 assumptions and models. Individual growth is a key element in our understanding of shrub expansion throughout the 106 Arctic (Tape et al. 2006; Myers-Smith et al. 2011a, 2019; Elmendorf et al. 2012b; García Criado et al. 2020). Increasing 107 shrub cover and canopy height alters ecosystem processes and species interactions (Myers-Smith et al. 2011a; Tape et 108 al. 2016, 2018; Way and Lapalme 2021) through snow-trapping (Myers-Smith and Hik 2013), shading (Blok et al. 109 2010), hydrology and albedo (Pomeroy et al. 2006; Sturm et al. 2005), food webs (Ravolainen et al. 2014)-including 110 soil assemblages (DeMarco et al. 2014)-and habitat provision for wildlife (Boelman et al. 2015). Shrub encroachment 111 has been linked to warming in studies using dendrochronology (Forbes et al. 2010), remote sensing (Myneni et al. 1997; 112 Myers-Smith et al. 2020), field observations (Hudson and Henry 2009; Myers-Smith et al. 2011b), and experiments 113 (Elmendorf et al. 2012a, 2015; Khorsand Rosa et al. 2015; Frei and Henry 2021). To accurately predict the future 114 structure and function of northern ecosystems, we must understand how plant growth is changing (Myers-Smith et al. 115 2020), especially the role of phenology as ecological dynamics change under warming (Keenan and Richardson 2015; 116 Myers-Smith et al. 2019; Bjorkman et al. 2020; Samplonius et al. 2020).

117

Arctic vegetation change plays a key role in regional and global feedback loops (Liston et al. 2002; Sturm et al. 2005; Pearson et al. 2013; Grosse et al. 2016) and carbon budgets (Piao et al. 2007; McGuire et al. 2009; Parker et al. 2021). As phenology changes, we expect compositional shifts and increased growth during longer growing seasons (Myneni et al. 1997; Ernakovich et al. 2014; Panchen and Gorelick 2017). Biome-wide shifts in growth rates and community composition could have profound implications for global carbon budgets through biomass accumulation (Piao et al. 2007) and decomposition (DeMarco et al. 2014). Warming drives earlier leaf emergence (Ernakovich et al. 2014; Park et al. 2016), which has been linked with increased plant productivity using remote sensing observations (Myneni et al. 125 1997). And thus, studies of satellite-derived spectral greening trends have linked changes in phenology to changes in
126 plant productivity (Myneni et al. 1997; Park et al. 2016; Kim et al. 2021). Furthermore, accurate Earth-system models
127 depend on our understanding of plant growth-climate relationships and ecosystem-climate feedbacks (Sturm et al. 2005;
128 Loranty and Goetz 2012; Richardson et al. 2013; Pearson et al. 2013; Fisher et al. 2018; Bonan and Doney 2018).
129 Despite underpinning global models, uncertainty remains in the expected association between phenology and growth of
130 Arctic plants and whether warmer temperatures or longer growing seasons are the primary drivers of increasing tundra
131 plant productivity.

132

133 In this study, we use dendroecology to test the correspondence between *in situ* phenology observations, environmental 134 factors, and radial growth of Salix arctica Pall. (Salicaceae) on Qikiqtaruk - Herschel Island in the Western Canadian 135 Arctic. Salix arctica has a circum-Arctic distribution, woody tissues which enable dendrology, and exceptionally 136 closely monitored phenology at the site (Myers-Smith et al. 2019), offering a suitable species for this study. At this site, 137 S. arctica phenology has advanced in both spring and autumn, though autumn only marginally, overall lengthening the 138 growing season by two days per decade (Myers-Smith et al. 2019). We test three questions: (1) Do longer growing 139 seasons facilitate greater shrub radial growth? (2) Of phenological metrics, does leaf emergence date, senescence date, 140 or growing season length best explain shrub radial growth? (3) Do climatic factors (air temperature, precipitation, sea 141 ice concentration, or snowmelt), or maximum plant productivity (estimated through the spectral Normalised Difference 142 Vegetation Index, NDVI), explain shrub radial growth better than phenology? We hypothesise that: (1) A longer 143 realised period of growth will increase radial growth; (2) Growing season length will explain shrub radial growth better 144 than leaf emergence or senescence date as it encompasses the cumulative change in the growth period; (3) Growing 145 season length will best explain shrub radial growth relative to temperatures or other variables as short growing seasons 146 in Arctic systems are limiting plant growth (Myers-Smith et al. 2019).

147

#### 148 Methods

#### 149 Study site

We studied *S. arctica* phenology and radial growth on Qikiqtaruk – Herschel Island, YT, Canada (69.57°N, 138.90°W)
(Fig. 1). The island is approximately 100 km<sup>2</sup> in area, with soils formed of marine and glacial deposits atop ice-rich
permafrost (Burn and Zhang 2009). Qikiqtaruk sits at the northerly extent of tall shrubs, particularly *Salix richardsonii*Hook. (Salicaceae) (Myers-Smith et al. 2011b), which feature heavily in its flora alongside *Eriophorum vaginatum* L.

154 (Cyperaceae) tussock tundra and dwarf shrub heath rich in *S. arctica* (Myers-Smith et al. 2019).



- *Fig. 1* A map showing the location of Qikiqtaruk (69.57°N, 138.90°W) within Canada and Sentinel 2 false-colour *images showing the location of the transects on Qikiqtaruk. The purple dots represent the ends of the five transects, and the purple box shows the area within which all samples were taken. The orange square shows the location of the phenology plots*
- 1
- 159
- 160 Salix arctica *sampling*
- 161 As woody perennials, shrubs grow annual rings of wood which record radial growth over time (Myers-Smith et al.
- 162 2015b). Here, we focus on *S. arctica*, a prostrate willow with a circum-Arctic range which reaches as far north as the
- 163 north coast of Greenland (Argus 2007). We collected 38 shrub samples on a coastal floodplain on the east side of the

- 164 island in the Ice Creek watershed, a site of known vegetation change (Myers-Smith et al. 2019). The vegetation is
- 165 dominated by Salix spp., defined by patches of S. richardsonii, a canopy-forming willow, with S. arctica at ground-
- 166 level. Cross sections of 3–5 cm in length were taken from the base of the thickest stem of six to eight individuals each
- along five parallel transects on the Qikiqtaruk floodplain in July 2016. Individuals were at least 10 m apart and transects
  were 50–100 m apart to lower the risk of sampling clones, particularly given the sprawling prostrate growth form of *S*.
- 169 *arctica* (Argus 2007).
- 170

#### 171 Dendroecological methods

We sliced thin sections (~  $25 \mu$ m) of each sample with a sledge microtome, then photographed (Online Resource 1) and measured the rings along four radii per sample using the ObjectJ package (1.04a) for ImageJ (2.0.0-rc-59/1.51j). We visually crossdated samples to check for partly missing rings, then averaged the radii for each individual and crossdated again to check for entirely missing rings between samples.

176

177 We removed the first two years of radial growth data for each individual to account for age-related growth effects, and 178 the data from the year of sampling, as the growth for the season was not completed (Myers-Smith et al. 2015b). To 179 maintain a minimum of four years of radial growth data after removing the data from 2016 and the first two years, the 180 sample size was reduced to 28 individuals, running from 2002-2015 (Online Resource 1). Individuals with fewer than 181 seven years of radial growth were removed and we calculated the basal area increments from the ring width data. We 182 then detrended the basal area increment data, fitting a smoothing spline (dplR package for R, f = 0.5, nvrs = 0.67), 183 which removed variation in radial growth beyond interannual variation. We visually assessed different detrending 184 methods, testing spline, negative exponential and no detrending, and found that spline detrending best removed 185 individual growth trends to help focus our analyses on interannual variation in radial growth (Myers-Smith et al. 186 2015b). We therefore used the detrended basal area increment data for the main text analysis. We also conducted the 187 same analysis using ring widths and the negative exponential detrending method and found similar results to our main 188 text analysis, these are presented in the supplementary materials (Online Resource 3). For the statistical analysis, all 189 variables were normalised between 1 and -1, so that effect sizes could be compared.

190

191 Phenology data

192 Phenological observations have been collected since 2001 by Yukon Parks Rangers on Qikiqtaruk every 2–3 days from

193 April (snowmelt) until September (leaf senescence) along established transects (~250 m from our study site). The

- rangers record phenology throughout the growing season (Myers-Smith et al. 2019) in line with ITEX protocol (Prevéy
- et al. 2021). The phenological dates used in this study are the date of first leaf bud burst (leaf emergence) and the date

196 of first vellowing of leaves (leaf senescence) averaged across the 20 S. arctica individuals in the observation transect. 197 We calculated growing season length (GSL) as the number of days between leaf emergence and senescence. Please note 198 that the individuals from the phenological modelling are not the same individuals (destructively) sampled for the 199 dendroecological growth time-series. However, previous findings indicate that S. arctica phenology is generally 200 consistent across individuals at the site (Myers-Smith et al. 2019, Assmann et al. 2020). Considering further the close 201 proximity of the phenological monitoring sites and dendroecological transects we are therefore confident that the 202 phenometrics from the phenological monitoring are representative for the individuals sampled for the dendroecological 203 analysis.

204

## 205 Other environmental data

206 We also compared radial growth to seasonal air temperatures and precipitation, snowmelt, sea ice concentration, and 207 productivity. The temperature data came from Environment Canada Qikiqtaruk-Herschel Island weather station (ID 208 1560) and precipitation data came from the ERA5 gridded dataset produced by Copernicus Climate Change Service and 209 the European Centre for Medium-Range Weather Forecasts (Hersbach et al. 2020). We collated data into seasons 210 (spring: April-May, summer: June-July, autumn: August-September, winter: October-March) including the lagged 211 data for the preceding summer and autumn, as monthly resolution was higher than useful for this study. Snowmelt data 212 are from the Qikiqtaruk phenology dataset, where the date at which transects are free of snow is recorded (Myers-Smith 213 et al. 2019). We used the onset of sea ice melt data (Assmann et al. 2019), determined using the NOAA/NSIDC Climate 214 Data Record (CDR) v3 Passive Microwave Sea Ice Concentrations (Meier et al. 2017). Productivity data are the annual 215 maxima (estimated by smoothing trends in the data with a generalised additive model) of the MODIS MOD13A1v6 216 NDVI satellite dataset (Myers-Smith et al. 2020).

217

### 218 *Statistical analysis*

We selected our models *a priori*, using single predictor models to compare individual variables and assess their predictive power on shrub radial growth. We used a Bayesian framework for our analyses including weakly informative priors of a half Student-t prior with three degrees of freedom. The effect sizes of models were assessed relative to their credible intervals (95%). If the credible intervals for the estimated slope did not cross zero, we considered an effect to be significant. If the credible interval of the model slope sits at zero or fluctuates between overlapping zero and not overlapping zero, we consider the effect to be marginally significant.

225

226 To test the relationships between variables we used hierarchical linear mixed-effect models, with year, transect, and 227 individual as random effects, and individuals nested within transects. We used mixed models due the hierarchical

228	structure of our data, caused by non-independence of individuals within transects and within a given year. Shrubs also
229	shared conditions in each year, making them non-independent. The variability in absolute growth among individuals
230	was high, as seen during crossdating, but due to sampling a single species within a relatively small area, similar relative
231	growth responses were expected across groups. Thus, we chose not to use random slopes, only random intercepts. As a
232	secondary analysis presented in the supplementary materials (Online Resource 2, Online Resource 3), we used a
233	frequentist framework and compared models using AIC to see whether the models fit the data better than a null model
234	using the conventional threshold ( $\Delta AIC_{null} \ge 2$ ) (Akaike 1974). We also calculated conditional and marginal pseudo-R <sup>2</sup>
235	to test the absolute model fit including and excluding random effects respectively (Nakagawa and Schielzeth 2013). We
236	used models with maximum likelihood estimation for AIC comparisons, and models with restricted maximum
237	likelihood estimation for pseudo-R <sup>2</sup> and effect size values. Residuals of models were visually assessed for normality
238	with fitted-residual plots, and temporal autocorrelation (first- or second-order) with correlograms (Online Resource 2).
239	We did not detect a signal of temporal autocorrelation, and residuals were similar across all models (Online Resource
240	2). We tested the correlation among all environmental and phenological variables and correlation coefficients varied
241	between -0.76 and 0.75 (Online Resource 1).

All statistical tests were carried out in R (3.6.3), via RStudio (1.2.1335), including the brms package for Bayesian
analysis: Code and data are available at the following GitHub repository:

245 <u>https://github.com/ShrubHub/ShrubRingPhenoHub</u>

246

# 247 Results

248 Contrary to our first hypothesis that a longer realised period of growth will increase radial growth, we found that 249 growing season length had no effect on S. arctica basal area increments, nor was there a lagged effect from the previous 250 growing season (Table 1, Fig. 2). Contrary to our second hypothesis that growing season length will explain shrub 251 radial growth better than leaf emergence or senescence date, it was leaf senescence date that was the best predictor of 252 radial growth amongst these variables. Radial growth was significantly greater in years with an earlier leaf senescence 253 date (negative effect), and we detected no relationship between radial growth and leaf emergence date, growing season 254 length, or previous growing season length (Table 1, Fig. 2). Contrary to our third hypothesis that growing season length 255 will best explain shrub radial growth relative to temperatures and other variables, we found that rather than growing 256 season length, summer temperature was the best predictor amongst all variables. Summer temperature explained 2.8% 257 of the variation in radial growth (Fig. 3, marginal pseudo- $R^2$ , Online Resource 2), with higher temperature coinciding 258 with higher annual radial growth. However, the overall explanatory power of the models was low with no model 259 explaining more than 2.8% of variation without including random effects (marginal pseudo-R<sup>2</sup>, Online Resource 2).



Fig. 2 Radial growth corresponded weakly with phenological variables and more strongly with summer temperature.
The relationships of radial growth with both leaf senescence and summer temperature are statistically significant.
Scatter plots show the four phenological and two temperature variables we hypothesised to have relationships with
radial growth (basal area increment, indexed) in a given year over the period 2002-2015. Trendlines are predictions
from the hierarchical Bayesian models, dashed trendlines indicate a non-significant effect, the shaded areas represent
95%, 80%, and 50% credible intervals of the model estimates

268 All other variables aside from leaf senescence date and summer temperature were non-significant predictors of radial 269 growth (Fig. 3). Leaf emergence date; current and previous years' growing season length; temperatures from the winter, 270 spring, autumn, and the previous year; all precipitation models; and snowmelt date did not explain variation in radial 271 growth (Fig. 3). We used year as a random effect in our models (individual and transect level growth variation are 272 accounted for during detrending and scaling of the radial growth index) and its effect was significant, indicating 273 variation in radial growth among years beyond the effect of the bioclimatic variable of interest in that year (Online 274 Resource 2, Online Resource 3). The model results agreed between the different detrending approaches, and though 275 exact effect sizes differed slightly, the results were similar when models used detrended ring width data or used 276 detrended basal area increment (Online Resource 3).

277

**278Table 1:** Statistical results for the hierarchical Bayesian models relating radial growth (basal area increment) to**279**phenology, temperature, precipitation, NDVI, sea ice concentration, and snow-free date (results for ring widths are**280**included in Online Resource 3). All models span the period 2002–2015. Asterisk (\*) indicates a model where the 95%**281**credible intervals do not overlap zero and which we therefore consider to be significant. Sample depth per year as**282**follows: 2002, n=5; 2003, n=6; 2004, n=8; 2005, n=10; 2006, n=14; 2007, n=16; 2008, n=20; 2009, n=23; 2010,**283**n=25; 2011, n=27; 2012-5, n=28 (Online Resource 1). Only the fixed effects are shown here, the results for the**284**random effects are available in Online Resource 2.

Predictor variable	Estimate	Standard Error	Lower 95% CI	Lower 95% CI	R-hat	Bulk ESS	Tail ESS	Effect
Leaf emergence	-0.04	0.06	-0.16	0.08	1.00	2494	2063	Fixed
Leaf senescence	-0.10	0.04	-0.18	-0.01	1.00	2613	1760	Fixed
Growing season length	-0.05	0.06	-0.18	0.07	1.00	2268	1700	Fixed
Previous growing season length	-0.02	0.06	-0.13	0.09	1.00	3014	2128	Fixed
Previous summer temperature	0.00	0.06	-0.12	0.13	1.00	2905	1892	Fixed
Previous autumn temperature	0.09	0.05	-0.01	0.20	1.00	2743	2005	Fixed
Winter temperature	-0.02	0.05	-0.12	0.07	1.00	1887	1917	Fixed
Spring temperature	0.02	0.05	-0.09	0.12	1.00	2618	2475	Fixed
Summer temperature	0.12	0.05	0.01	0.22	1.00	3546	2036	Fixed
Autumn temperature*	0.07	0.06	-0.04	0.18	1.00	2968	2152	Fixed
Previous summer precipitation	-0.03	0.06	-0.14	0.08	1.00	2698	2094	Fixed
Previous autumn precipitation	-0.01	0.05	-0.11	0.09	1.00	2389	2016	Fixed
Winter precipitation	-0.01	0.06	-0.13	0.12	1.00	2999	2697	Fixed
Spring precipitation	-0.01	0.05	-0.10	0.08	1.00	2559	1856	Fixed
Summer precipitation	0.02	0.06	-0.10	0.13	1.00	2681	2159	Fixed
Autumn precipitation	-0.01	0.06	-0.12	0.10	1.00	2781	2215	Fixed
MODIS NDVI	0.01	0.06	-0.10	0.12	1.00	2164	1501	Fixed
Minimum sea ice extent	-0.14	0.12	-0.38	0.11	1.00	3588	2755	Fixed
Sea ice melt onset date	-0.10	0.06	-0.22	0.03	1.00	2373	2286	Fixed
Date snow free	-0.04	0.09	-0.21	0.13	1.00	1493	724	Fixed



Fig. 3 Only models of leaf senescence and summer temperature significantly explained variation in radial growth, most
variables showed no relationship to radial growth (Table 1). This plot shows standardised effect sizes (slopes) of
hierarchical Bayesian models of phenological events (purple), seasonal temperature (red), seasonal precipitation
(green), NDVI (yellow), minimum sea ice extent, sea ice concentration, and snowmelt date (all blue) on radial growth.
The centre line is the effect and error bars are 95% credible intervals. For ease of comparison between effect sizes,
explanatory variables in this analysis are variance-scaled from -1 to 1

# 294 Discussion

295 Through a unique study of long-term *in situ* phenology monitoring and dendroecology, we compared interannual

variation in phenology, environmental conditions, and NDVI to interannual variation in radial growth of S. arctica on

297 Qikiqtaruk. We found that summer temperatures and leaf senescence - but not leaf emergence or growing season length

298 – explained variation in radial growth for the widespread Arctic shrub S. arctica (Table 1, Fig. 3). Precipitation, sea ice,

snowmelt, and NDVI did not correspond with variability in interannual radial growth in our study. Thus, we did not

300 find support for the hypothesised relationship between phenology and radial growth. Our results suggest that factors

other than the timing of the growing period from leaf emergence through senescence, such as temperature, can exert a
 larger influence on shrub growth in this tundra ecosystem. These findings ultimately have implications for how tundra

303 shrub growth is modelled and thus the projection of Arctic carbon budgets.

- 304
- 305 Phenology

306 We found no evidence that earlier leaf emergence and longer growing seasons corresponded with increased radial 307 growth in S. arctica, including growth in the following year (Table 1, Fig. 3). Results for preliminary analyses including 308 other willow species from this and other sites have reached similar conclusions (Angers-Blondin 2019). Earlier leaf 309 emergence did not result in greater S. arctica radial growth; however, we did find evidence of greater radial growth in 310 years with earlier leaf senescence. Although we cannot identify a particular biological mechanism linking earlier leaf 311 senescence with enhanced radial growth, early leaf senescence was correlated with warmer summer temperatures 312 (Pearson's product-moment correlation, df = 263, p < 0.001,  $\rho = -0.60$ , Online Resource 1). Individuals may reach a 313 threshold after intense early growth (Rumpf et al. 2014), allowing for early cessation of growth aboveground. 314 Alternatively, growth and leaf age could be deterministic (Oberbauer et al. 2013; Keenan and Richardson 2015; 315 Semenchuk et al. 2016; Parker et al. 2017), with growth ending at a fixed time after growth begins each year. Or earlier 316 leaf senescence could occur in years with warmer temperatures without a mechanistic link between the two variables. 317 Our findings are in line with previous evidence that the timing of tundra plant senescence is driven at least in part by 318 non-climatic factors (Arft et al. 1999; Oberbauer et al. 2013). Taken together, our results suggest that shifts to earlier 319 shrub leaf emergence and longer growing seasons are not necessarily driving changes in tundra shrub growth, contrary 320 to interpretations of satellite remote sensing data (Myneni et al. 1997; Zeng et al. 2011, 2013; Arndt et al. 2019) and 321 reviews (Ernakovich et al. 2014).

- 322
- 323 Temperature

We found that higher summer temperatures increased the radial growth of *S. arctica* at our site. The summer is the peak
season for growth and individuals are sensitive to warming in this period (Andreu-Hayles et al. 2020), as observed
across the biome (Myers-Smith et al. 2015a; Myers-Smith and Hik 2018) from dendrochronology (Forbes et al. 2010;
Blok et al. 2011; Myers-Smith et al. 2011a; Li et al. 2016; Weijers et al. 2018; Le Moullec et al. 2019; Prendin et al.
2022), repeat photography (Sturm et al. 2001; Tape et al. 2006), and experiments (Elmendorf et al. 2012a, 2015;
Khorsand Rosa et al. 2015; Frei and Henry 2021). Temperature-growth relationships are heterogeneous across the

tundra biome, with relatively low climate sensitivity observed on Qikiqtaruk compared with other mid-latitude tundra

331 locations (Myers-Smith et al. 2015a). Growth response to temperature has decreased over time in another dwarf willow 332 species, Salix polaris Wahlenb. (Salicaceae), at Bjørnøya, Svalbard (Owczarek et al. 2021), suggesting that growth 333 responses may not be fixed over long timescales. The growth response to early leaf senescence suggests the importance 334 of resource accumulation for growth in the following season, shrubs may senesce above-ground but remain active 335 below-ground for longer periods. There is increasing evidence that above-ground phenology may be asynchronous with 336 below-ground root growth (Blume-Werry et al. 2016, 2017; Ögren 2017; Liu et al. 2022), though root phenology itself 337 may not respond to autumn warming (Schwieger et al. 2018). Snow cover insulates shrubs from winter and spring 338 temperatures (Kelsey et al. 2020; Rixen et al. 2022), and Krab et al. (2018) found diverging shrub radial growth 339 responses to winter temperature, spring warming, and snowmelt among species. Vaccinium vitis-idaea L. (Ericaceae) 340 grew more with delayed snowmelt with a contrasting reduction in growth in *Empetrum nigrum* L. (Ericaceae). We, 341 however, found no association between temperatures in the previous year and radial growth, and no relationship for 342 winter, spring, and autumn temperatures and radial growth (Table 1, Fig. 3).

343

#### 344 *Hydrology*

345 We did not find a strong influence of summer precipitation, sea ice or snowmelt on interannual variation in radial 346 growth S. arctica in this study. Growth of Arctic shrubs can be moisture-limited (Keuper et al. 2012; Ackerman et al. 347 2017; Buchwal et al. 2020; Weijers 2022). Moisture sensitivity of growth can depend on temperature (Li et al. 2016), 348 and can vary within (Thompson and Koenig 2018) and between sites (Myers-Smith et al. 2015a). Soils on Oikigtaruk 349 are frequently saturated, likely reducing the impacts of drought locally (Myers-Smith et al. 2019). We did not detect any 350 influence of precipitation from summer rain, snowmelt, or cloud cover (Table 1, Fig. 3). Decreasing snow cover reduces 351 soil insulation in winter and limits productivity increase under warming and earlier phenology in Alaska (Kelsey et al. 352 2020). The lack of a precipitation signal detected in our study could be influenced by our use of gridded climate datasets 353 due to a lack of a complete local record for precipitation at this site. Gridded climate datasets poorly capture spatially-354 variable precipitation, due to the paucity of Arctic meteorological stations and the high spatial variability of 355 precipitation (Macias-Fauria et al. 2014; Myers-Smith and Myers 2018). For sea ice, we found that lower annual 356 minima and earlier melt are weakly associated with increased radial growth of S. arctica, although phenology for this 357 species was not found to vary with sea ice extent (Assmann et al. 2019). Sea ice could influence plant growth and 358 phenology through interactions with local climate (Post et al. 2009; Bhatt et al. 2010; Kerby and Post 2013; Macias-359 Fauria et al. 2017; Assmann et al. 2019) and drought-stress (Forchhammer 2017; Buchwal et al. 2020). We found no 360 relationship between snowmelt date and radial growth, which is consistent with the primary mechanism of snowmelt 361 controlling phenology and so influencing plant growth (Assmann et al. 2019; Myers-Smith et al. 2019). Taken together, 362 our results suggest that temperature rather than growing season length, precipitation or sea ice dynamics was the

primary factor controlling S. arctica radial growth on Qikiqtaruk.

364

365 NDVI

366 We observed no correlation between NDVI and interannual variation in S. arctica radial growth, consistent with results 367 for other shrub species at the site (Myers-Smith et al. 2019). While satellite datasets do not entirely correspond with 368 each other (Guay et al. 2014), and shrub biomass cannot be directly estimated from NDVI alone (Cunliffe et al. 2020), 369 NDVI is easily-scaled, well-studied, and part of a broader picture of complex Arctic tundra vegetation change (Myers-370 Smith et al. 2020). Arctic shrubification has been linked with satellite-derived Arctic greening trends (Macias-Fauria et 371 al. 2012), and comparing ground observations to spectral greening observed by satellites improves broad-scale 372 interpretation of these trends (Myers-Smith et al. 2020). Correlation of NDVI and shrub growth has been found in some 373 studies (Forbes et al. 2010; Macias-Fauria et al. 2012), but is not universal and varies with site and the time of year 374 (Blok et al. 2011; Brehaut and Danby 2018; Andreu-Hayles et al. 2020). Taken together, these results suggest that 375 satellite spectral greening indices are not capturing all of the variation in plant productivity, including the length of the 376 snow-free season, indicated by analyses with shrub radial growth (Angers-Blondin 2019; Berner et al. 2020).

377

#### 378 Study limitations

379 While our findings bring together phenology and dendrochronology, two important fields of study of Arctic change, 380 there are limitations. Sampling stem elongation (primary growth) and root collars rather than stems alone would 381 improve the capture of interannual variation in shrub productivity. Primary and secondary (radial) growth can be driven 382 by different controls (Bret-Harte et al. 2002; Campioli et al. 2012a, 2012b), so study of annual stem increments or other 383 measures of shrub growth in addition to radial growth would more robustly address questions of shrub growth (Myers-384 Smith et al. 2015b). Root collars show greater climate sensitivity (Ropars et al. 2017) and less response to individual 385 conditions than stems (Sonesson and Callaghan 1991; Sadras and Denison 2009; Myers-Smith et al. 2015b), yet root 386 collars are challenging to find and excavate in clonal species such as S. arctica, and are more destructive to sample. We 387 were not easily able to locate root collars consistently at this site (Angers-Blondin 2019). The destructive nature of 388 dendrological sampling also prevented us from sampling the individuals in the long-term phenology transect directly, so 389 we sampled nearby individuals as phenology is consistent across the site (Myers-Smith et al. 2019). Future research 390 across different sites and species using localised climate and microenvironmental variables may shed more light on the 391 relationships between plant phenology and growth.

392

**393** *Future study* 

394 Modern techniques facilitate below-ground monitoring of tundra plant phenology and root growth (Iversen et al. 2015;

395 Sloan et al. 2016; Blume-Werry et al. 2016, 2017), exposing an overlooked dimension of tundra dynamics. An 396 increasing number of studies indicate phenological asynchrony above and below ground at sub-Arctic sites (Ögren 397 2017; Blume-Werry et al. 2017), with below-ground root growth extending into the late summer and autumn in now-398 thawed soils. Fungal symbiotes such as mycorrhizae can influence plant growth and carbon exchange in tundra shrubs 399 which could be altering growth-climate interactions (Clemmensen et al. 2006; Compant et al. 2010; Deslippe et al. 400 2011). Iler et al. (2013) suggest that phenology responses to warming are reaching physiological limits in some Arctic 401 and alpine species, potentially reducing the magnitude of future change. Collins et al. (2021) found that reproductive 402 and vegetative phenologies are affected differently by experimental warming, which could alter ecosystem dynamics 403 via trophic mismatches and resource-allocation (Post and Forchhammer 2008; Clausen and Clausen 2013; Kerby 2015; 404 Wheeler et al. 2015). There has been relatively little investigation of plant senescence and the drivers of the end of the 405 growing season, creating uncertainty in our understanding of plant responses to warming across the growing season.

406

407 Innovative techniques such as drone-derived biomass estimates could also help with scaling up to landscape-wide 408 analyses (Cunliffe et al. 2020). Newer approaches to studying tundra plant phenology such as time lapse cameras (a.k.a. 409 phenocams) are overcoming inherent challenges of data collection in the Arctic (Westergaard-Nielsen et al. 2017; 410 Richardson et al. 2018; Parmentier et al. 2021). Local observations of plant phenology and growth can be scaled up 411 using drone and satellite data to bridge scale gaps and form a landscape perspective on tundra productivity change 412 (Riihimäki et al. 2019; Assmann et al. 2019, 2020; Cunliffe et al. 2020). Challenges of scaling and data collection are 413 being met by technological solutions, allowing us to see Arctic change from new angles and more clearly than ever 414 before. Though further research is required, particularly for Arctic systems (Diepstraten et al. 2018), the increasing 415 scope of monitoring of above- and below-ground plant responses encompassing phenology and growth allows for the 416 investigation of key knowledge gaps about tundra ecosystem responses to global change.

417

## 418 Conclusions

Our findings demonstrate that plant phenology does not necessarily predict growth in an Arctic shrub, but that warmer temperatures in the summer are associated with increased annual radial growth. Interannual variation in precipitation, sea ice, snow cover, and MODIS NDVI for the landscape were not strongly related to radial growth. Our results indicate that future Arctic warming will likely enhance shrub growth and encroachment (Tape et al. 2006; Myers-Smith et al. 2011a; García Criado et al. 2020). Where this growth is not limited by water or nutrients (Mack et al. 2004; Myers-Smith et al. 2015a; Ackerman et al. 2017), there may be significant consequences for water, energy, and carbon fluxes (Loranty and Goetz 2012; Pearson et al. 2013; Parker et al. 2021). Taller shrub canopies could influence soil

426 temperatures, litter decomposition rates, nutrient cycling and ultimately the tundra carbon cycle (Sturm et al. 2005; Blok

- 427 et al. 2010; Myers-Smith et al. 2011a; DeMarco et al. 2014; Way and Lapalme 2021). While questions remain in these
- 428 complex systems, studying shrub phenology and growth data for other sites and species and incorporating a below-
- 429 ground perspective on plant phenology (Iversen et al. 2015) and growth will paint a clearer panarctic picture of plant
- 430 responses to rapid Arctic warming (Myers-Smith et al. 2020). Investigating the magnitude and direction of change on-
- 431 the-ground in tundra ecosystems is necessary to validate assumptions that underpin remote sensing studies (Myers-
- 432 Smith et al. 2019; Piao et al. 2019; Cunliffe et al. 2020), strengthening our understanding of tundra plant responses to
- 433 warming. Teasing apart the complex mechanisms between climate change and plant growth in tundra ecosystems is
- 434 vital to improve projections of how Arctic vegetation change influences global climate.

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- 737 Online Resource 1 Boyle, J.S.\*, Angers-Blondin, S., Assmann, J.J., Myers-Smith, I.H. Summer temperature but not
- 738 growing season length influences radial growth of Salix arctica in coastal Arctic tundra Polar Biology
- \* School of Geosciences, University of Edinburgh, Edinburgh, Scotland, United Kingdom; email:
- 740 joe.scott.boyle@gmail.com
- 741
- 742 Here, we present a thin section image (Fig. A), sample depth figures (Fig. B), and correlation matrix for environmental
- 743 and phenological variables (Fig. C). All code and data used for this study and generated during our analyses are
- 744 available in a GitHub repository (<u>https://github.com/ShrubHub/ShrubRingPhenoHub</u>).
- 745



- 746 Fig. A An example of one of the four thin section radii of S. arctica as used in this study. This stem came from transect
- 747 5, and was 18 years old when it was sampled. The rings can be made out clearly, with the outer ones being the
- 748 youngest and those with wider rings and a larger surface area being deposited in years of higher growth



Fig. B Sample depth plots showing the age distribution of the samples from different transects and the number of samples representing different years across the different sample transects. The oldest samples (up to 28 years) were taken from transect 5 and the youngest (from 3 years) from transect 3 (indicated by colour). The number of total samples is lower in the second plot as it shows the data after processing where some have been removed for having too few years' data. The dotted line indicates the first year of phenology data — the cutoff point for the data used in this study

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nerecet unnermerer	toute up us in the second propriet in the training the second sec	)* )
Lest est rear Sprear in the in man	Provide in the strate of the s	
Snow melt 0.75 0.36 -0.17-0.53 0.16 -0.24-0.65 -0.6 -0.14-0.08-0	0.12 0.34 0.38 -0.18 0.05 -0.5 0.35 0.1 0.2	
Leaf emergence 0.43 0.02 -0.75 0.04 -0.13-0.46-0.76-0.39-0.08-0	0.03 0.35 0.16 0.08 0.17 <mark>-0.51-0.15</mark> 0.44 0.28 - 0.8	
Leaf senescence -0.04 0.28 0.32 -0.1 -0.21 -0.25 -0.6 -0.31 -0.	<b>0.25-0.19</b> 0.19 0.02 0.2 <b>-0.24</b> 0.26 0.28 0.37	
GSL 0.2 0.06 0.33 0.63 -0.03 -0.15 -0.03	-0.6	
Prev. summer temp. 0.58 -0.25-0.15-0.03 0 -0	0.19 0 0.68 0.17 0.2 -0.12 0.29 -0.21-0.16 - 0.4	
Prev. autumn temp0.08-0.1 0.21 0.3 -0	0.36-0.12 0.29 0.42 0.41 -0.19 0.07 -0.32-0.03	
Winter temp. 0.62 -0.15-0.42 0	.34 0 -0.35 0.1 0.05 0.5 -0.35 0.23 0.16 - 0.2	
Spring temp. 0.25 -0.31 0	.06 -0.47-0.38 0.03 -0.04 0.63 0.1 0.02 -0.21	
Summer temp. 0.51-0	0.03 0.04 0.05 -0.06-0.24 0.1 0.41 -0.39-0.52 - 0	
Prev. summer precip	$0.2 \ 0.32 \ 0.19 - 0.26 - 0.31 - 0.21 \ 0.12 - 0.54 - 0.38$	)
Prev. autumn pre	eCiD. 0.51 -0.45-0.09-0.24 0.01 -0.19-0.18	
Winter	precip. <u>-0.25</u> 0.36 -0.09 0.43 -0.3 -0.23	ŀ
Sp	ring precip. 0.42 <mark>-0.13-0.34</mark> 0.23 0.32	
S	ummer precip. 0.1 -0.02 0.29 0.3	;
	Autumn precip0.19 0.36 _0.07	ł
	MODIS max. NDVI -0.51-0.26	,
	Sea ice min. extent 0.461	

755 Fig. C Correlation matrix for environmental and phenological variables used in the overall analyses. Correlation

*coefficients varied between -0.76 and 0.75* 

757 Online Resource 2 - Boyle, J.S.\*, Angers-Blondin, S., Assmann, J.J., Myers-Smith, I.H. Summer temperature – but not

758 growing season length – influences radial growth of Salix arctica in coastal Arctic tundra Polar Biology

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760 joe.scott.boyle@gmail.com

761

762 Here, we present summary tables of a frequentist mixed-model analysis using basal area increment (Table A), as well as

the full results table for the Bayesian analysis using basal area increment, including fixed and random effects (Table B).

764 We also present temporal autocorrelation plots (Fig. A). All code and data used for this study and generated during our

765 analyses are available in a GitHub repository (<u>https://github.com/ShrubHub/ShrubRingPhenoHub</u>).

- 766
- 767 Table A: Statistical results for the linear mixed effect models relating radial growth (basal area increment) to

768 phenology, temperature, precipitation, Normalised Differential Vegetation Index (NDVI), sea ice concentration, and

results for ring widths are included in Table S3). All models span the period 2002–2015. Asterisk (\*)

**770** *indicates a model for which*  $\Delta AIC_{null} \leq -2$ . *All*  $\Delta AIC_{null} > -2$  recorded as 0, as differences below 2 are statistically

771 *undetectable.* Sample depth for each year as follows: 2002, n=5; 2003, n=6; 2004, n=8; 2005, n=10; 2006, n=14;

772 2007, n=16; 2008, n=20; 2009, n=23; 2010, n=25; 2011, n=27; 2012-5, n= 28 (Online Resource 1)

Duadiatan yaniahla		Pseu	ıdo-R <sup>2</sup>	Effect Size		
rredictor variable		Marginal	Conditional	Effect	SE	
Leaf emergence	0	0.3%	3.0%	-0.042	0.057	
Leaf senescence	-4.6*	2.8%	2.9%	-0.097	0.035	
Growing season length	0	0.4%	3.1%	-0.052	0.057	
Previous growing season length	0	0.1%	2.9%	-0.026	0.053	
Previous summer temperature	0	<0.1%	3.2%	-0.002	0.058	
Previous autumn temperature	-2.3*	2.0%	3.1%	0.095	0.046	
Winter temperature	0	0.2%	3.1%	-0.023	0.042	
Spring temperature	0	0.1%	3.2%	0.018	0.049	
Summer temperature	-4.5*	2.8%	2.8%	0.121	0.044	
Autumn temperature	0	1.0%	3.0%	0.071	0.051	
Previous summer precipitation	0	0.2%	3.2%	-0.031	0.051	
Previous autumn precipitation	0	<0.1%	3.0%	-0.011	0.049	
Winter precipitation	0	<0.1%	3.1%	-0.006	0.058	
Spring precipitation	0	<0.1%	3.2%	-0.007	0.041	
Summer precipitation	0	<0.1%	3.2%	0.016	0.053	
Autumn precipitation	0	<0.1%	3.2%	-0.006	0.050	
MODIS NDVI	0	<0.1%	3.0%	0.121	0.090	
Minimum sea ice extent	0	0.6%	2.5%	-0.136	0.119	
Sea ice melt onset date	0	1.3%	2.8%	-0.096	0.058	
Date snow free	0	0.2%	3.0%	-0.042	0.058	

774 *Table B*: Statistical results for the hierarchical Bayesian models relating radial growth (basal area increment) to

phenology, temperature, precipitation, Normalised Differential Vegetation Index (NDVI), sea ice concentration, and

snow-free date. All models span the period 2002–2015. Asterisk (\*) indicates a model where the 95% credible intervals

777 do not overlap zero and which we therefore consider to be significant (significant fixed effects are in bold to distinguish

- them from random effects). Sample depth for each year as follows: 2002, n=5; 2003, n=6; 2004, n=8; 2005, n=10;
- **779** 2006, *n*=14; 2007, *n*=16; 2008, *n*=20; 2009, *n*=23; 2010, *n*=25; 2011, *n*=27; 2012-5, *n*= 28 (Online Resource 1)

Predictor variable	Estimate	Standard	Lower	Lower	R-hat	Bulk	Tail	Effect
	Estimate	Error	95% CI	95% CI	IX-nat	ESS	ESS	Enect
Leaf emergence	-0.04	0.06	-0.16	0.08	1.00	2494	2063	Fixed
Leaf senescence	-0.10	0.04	-0.18	-0.01	1.00	2613	1760	Fixed
Growing season length	-0.05	0.06	-0.18	0.07	1.00	2268	1700	Fixed
Previous growing season length	-0.02	0.06	-0.13	0.09	1.00	3014	2128	Fixed
Previous summer temperature	0.00	0.06	-0.12	0.13	1.00	2905	1892	Fixed
Previous autumn temperature	0.09	0.05	-0.01	0.20	1.00	2743	2005	Fixed
Winter temperature	-0.02	0.05	-0.12	0.07	1.00	1887	1917	Fixed
Spring temperature	0.02	0.05	-0.09	0.12	1.00	2618	2475	Fixed
Summer temperature	0.12	0.05	0.01	0.22	1.00	3546	2036	Fixed
Autumn temperature*	0.07	0.06	-0.04	0.18	1.00	2968	2152	Fixed
Previous summer precipitation	-0.03	0.06	-0.14	0.08	1.00	2698	2094	Fixed
Previous autumn precipitation	-0.01	0.05	-0.11	0.09	1.00	2389	2016	Fixed
Winter precipitation	-0.01	0.06	-0.13	0.12	1.00	2999	2697	Fixed
Spring precipitation	-0.01	0.05	-0.10	0.08	1.00	2559	1856	Fixed
Summer precipitation	0.02	0.06	-0.10	0.13	1.00	2681	2159	Fixed
Autumn precipitation	-0.01	0.06	-0.12	0.10	1.00	2781	2215	Fixed
MODIS NDVI	0.01	0.06	-0.10	0.12	1.00	2164	1501	Fixed
Minimum sea ice extent	-0.14	0.12	-0.38	0.11	1.00	3588	2755	Fixed
Sea ice melt onset date	-0.10	0.06	-0.22	0.03	1.00	2373	2286	Fixed
Date snow free	-0.04	0.09	-0.21	0.13	1.00	1493	724	Fixed
Leaf emergence*	0.06	0.03	0.00	0.13	1.00	1141	1461	Random
Leaf senescence*	0.04	0.03	0.00	0.10	1.00	1280	1715	Random
Growing season length*	0.06	0.03	0.00	0.13	1.00	1157	1517	Random
Previous growing season length*	0.06	0.03	0.00	0.13	1.00	1114	1702	Random
Previous summer temperature*	0.06	0.03	0.01	0.14	1.00	1248	1620	Random
Previous autumn temperature*	0.04	0.03	0.00	0.11	1.00	1465	1940	Random
Winter temperature*	0.06	0.03	0.00	0.13	1.00	1115	1667	Random
Spring temperature*	0.06	0.03	0.00	0.13	1.00	1161	982	Random
Summer temperature*	0.04	0.03	0.00	0.10	1.00	1431	1796	Random
Autumn temperature*	0.05	0.03	0.00	0.12	1.00	1153	1521	Random
Previous summer precipitation*	0.06	0.03	0.00	0.13	1.00	1031	1133	Random
Previous autumn precipitation*	0.06	0.03	0.01	0.13	1.00	858	1160	Random
Winter precipitation*	0.06	0.03	0.00	0.13	1.01	1076	1470	Random
Spring precipitation*	0.06	0.03	0.00	0.13	1.00	1147	1154	Random
Summer precipitation*	0.00	0.03	0.01	0.14	1.00	065	1271	Random
Autumn procinitation*	0.00	0.03	0.00	0.13	1.00	1217	1271	Random
MODIS NDVI*	0.00	0.03	0.01	0.14	1.00	1217	1307	Random
Minimum and inc autout*	0.00	0.03	0.00	0.13	1.00	1001	1421	Random Doudom
Nimmum sea ice extent*	0.05	0.03	0.00	0.12	1.00	1031	1000	Random
Sea ice meit onset date*	0.05	0.03	0.00	0.12	1.00	886	1080	Kandom Danal
Date snow free*	0.06	0.03	0.00	0.13	1.00	896	1502	Kandom



781 Fig. A Autocorrelation plots for each environmental and phenological variable used in the overall analyses (using

<sup>782</sup> basal area increment). We found no strong temporal autocorrelation for any of the variables tested

783 Online Resource 3 - Boyle, J.S.\*, Angers-Blondin, S., Assmann, J.J., Myers-Smith, I.H. Summer temperature – but not

784 growing season length – influences radial growth of Salix arctica in coastal Arctic tundra Polar Biology

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- 787
- 788 Here, we present summary tables of Bayesian (Table A) and frequentist (Table B) analyses using ring width detrended
- 789 with a negative exponential fit. We also present figures of the Bayesian analysis with ring width data (Figs. A-C). All
- code and data used for this study and generated during our analyses are available in a GitHub repository
- 791 (<u>https://github.com/ShrubHub/ShrubRingPhenoHub</u>).
- 792
- 793 Table A: Statistical results for the hierarchical Bayesian models relating ring width to phenology, temperature,
- 794 precipitation, NDVI, sea ice concentration, and snow-free date (Table S3). All models span the period 2002–2015.
- 795 Asterisk (\*) indicates a model where the 95% credible intervals do not overlap zero and which we therefore consider to
- 796 *be significant. Sample depth for each year as follows: 2002, n=5; 2003, n=6; 2004, n=8; 2005, n=10; 2006, n=14;*
- **797** 2007, n=16; 2008, n=20; 2009, n=23; 2010, n=25; 2011, n=27; 2012-5, n= 28 (Online Resource 1).

Predictor variable	Estimate	Standard Error	Lower 95% CI	Lower 95% CI	R-hat	Bulk ESS	Tail ESS	Effect
Leaf emergence	0.07	0.09	-0.11	0.25	1.00	2845	1993	Fixed
Leaf senescence	-0.04	0.07	-0.18	0.11	1.00	3499	2431	Fixed
Growing season length	-0.13	0.08	-0.29	0.04	1.00	2705	2259	Fixed
Previous growing season length	-0.03	0.08	-0.20	0.13	1.00	3361	2036	Fixed
Previous summer temperature	-0.05	0.10	-0.24	0.15	1.00	1082	710	Fixed
Previous autumn temperature	-0.09	0.07	-0.24	0.06	1.00	2198	1685	Fixed
Winter temperature	-0.10	0.06	-0.22	0.03	1.00	1888	1530	Fixed
Spring temperature	-0.08	0.07	-0.23	0.07	1.00	2921	1942	Fixed
Summer temperature	0.13	0.08	-0.03	0.29	1.00	2214	1809	Fixed
Autumn temperature*	0.15	0.08	0.00	0.31	1.00	3778	2774	Fixed
Previous summer precipitation	-0.12	0.08	-0.27	0.03	1.00	2284	1877	Fixed
Previous autumn precipitation	0.02	0.08	-0.14	0.19	1.00	2589	2102	Fixed
Winter precipitation	-0.02	0.09	-0.19	0.16	1.00	3253	2962	Fixed
Spring precipitation	-0.04	0.06	-0.17	0.09	1.00	3083	2402	Fixed
Summer precipitation	-0.03	0.09	-0.22	0.15	1.00	1384	635	Fixed
Autumn precipitation	0.08	0.08	-0.09	0.24	1.00	3214	2193	Fixed
MODIS NDVI	0.07	0.08	-0.10	0.22	1.00	2257	1500	Fixed
Minimum sea ice extent	-0.17	0.20	-0.55	0.22	1.00	3939	2493	Fixed
Sea ice melt onset date	-0.13	0.10	-0.32	0.06	1.00	3636	2929	Fixed
Date snow free	0.04	0.11	-0.18	0.25	1.00	2699	2482	Fixed
Leaf emergence*	0.06	0.03	0.00	0.14	1.00	1181	1608	Random
Leaf senescence*	0.04	0.03	0.00	0.10	1.00	1300	1488	Random
Growing season length*	0.06	0.03	0.00	0.12	1.00	1034	1005	Random
Previous growing season length*	0.06	0.03	0.00	0.13	1.00	1087	1326	Random
Previous summer temperature*	0.06	0.03	0.00	0.14	1.01	780	872	Random
Previous autumn temperature*	0.06	0.03	0.00	0.13	1.00	1157	1638	Random
Winter temperature*	0.06	0.03	0.01	0.13	1.00	1126	1813	Random
Spring temperature*	0.06	0.03	0.00	0.13	1.00	1067	1164	Random
Summer temperature*	0.04	0.03	0.00	0.10	1.00	1331	1830	Random

Autumn temperature*	0.05	0.03	0.00	0.12	1.00	1076	1401	Random
Previous summer precipitation*	0.06	0.03	0.01	0.13	1.00	1090	1851	Random
Previous autumn precipitation*	0.06	0.04	0.00	0.14	1.00	803	795	Random
Winter precipitation*	0.06	0.03	0.00	0.14	1.00	1028	1299	Random
Spring precipitation*	0.06	0.03	0.00	0.13	1.00	992	992	Random
Summer precipitation*	0.06	0.03	0.00	0.14	1.00	1029	1384	Random
Autumn precipitation*	0.04	0.03	0.00	0.11	1.00	1325	1900	Random
MODIS NDVI*	0.06	0.03	0.01	0.13	1.00	1077	1540	Random
Minimum sea ice extent*	0.05	0.03	0.00	0.12	1.00	1088	1393	Random
Sea ice melt onset date*	0.05	0.03	0.00	0.11	1.00	1265	1489	Random
Date snow free*	0.06	0.03	0.00	0.13	1.00	1217	1434	Random

*Table B:* Statistical results for the linear mixed effect models relating ring width to phenology, temperature,

800 precipitation, NDVI, sea ice concentration, and snow-free date. All models span the period 2002–2015. Asterisk (\*)

801 indicates a model for which  $\Delta AIC_{null} \leq -2$ . All  $\Delta AIC_{null} > -2$  recorded as 0, as differences below 2 are statistically

*undetectable. Sample depth for each year as follows: 2002, n=5; 2003, n=6; 2004, n=8; 2005, n=10; 2006, n=14;* 

2007, *n*=16; 2008, *n*=20; 2009, *n*=23; 2010, *n*=25; 2011, *n*=27; 2012-5, *n*= 28 (Online Resource 1).

D		Pseu	ıdo-R <sup>2</sup>	Effect Size		
Predictor variable		Marginal	Conditional	Effect	SE	
Leaf emergence	0	0.4%	2.7%	0.072	0.084	
Leaf senescence	0	0.2%	2.9%	-0.037	0.067	
Growing season length	0	1.2%	2.6%	-0.131	0.080	
Previous growing season length	0	0.1%	2.6%	-0.033	0.078	
Previous summer temperature	0	0.2%	2.6%	-0.051	0.084	
Previous autumn temperature	0	0.6%	2.7%	0.077	0.075	
Winter temperature	0	1.5%	2.6%	-0.098	0.054	
Spring temperature	0	0.8%	2.7%	-0.081	0.068	
Summer temperature	0	1.3%	2.8%	0.125	0.076	
Autumn temperature	-3.1*	2.2%	2.4%	0.155	0.065	
Previous summer precipitation	0	1.5%	2.5%	-0.118	0.065	
Previous autumn precipitation	0	0.1%	2.8%	0.023	0.071	
Winter precipitation	0	0.0%	2.7%	-0.016	0.085	
Spring precipitation	0	0.3%	2.7%	-0.038	0.059	
Summer precipitation	0	0.0%	2.8%	-0.022	0.078	
Autumn precipitation	0	0.8%	2.7%	-0.087	0.069	
MODIS NDVI	0	0.4%	2.4%	0.068	0.073	
Minimum sea ice extent	0	0.4%	2.3%	-0.178	0.180	
Sea ice melt onset date	0	1.0%	2.8%	-0.126	0.089	
Date snow free	0	0.1%	2.7%	0.037	0.101	



805 Fig. A Autocorrelation plots for each variable used in the overall analyses (using ring width). We found no strong

<sup>806</sup> temporal autocorrelation for any of the variables tested



Fig. B Annual ring width growth data corresponded weakly with phenological variables and more strongly with
summer and autumn temperatures. The relationship with autumn temperature is statistically significant. Scatter plots
show four phenological and two temperature variables' relationships to radial growth (ring width, indexed) in a given
year over the period 2002-2015. Trendlines are predictions from the hierarchical Bayesian models, dashed trendlines
indicate a non-significant effect, the shaded areas represent 95%, 80%, and 50% credible intervals of the model
estimates



Fig. C In the analysis of annual ring widths, no variables tested had a strong, direct relationship to growth, autumnal
temperature was marginally significant, with a credible interval at or just overlapping zero (Table A). This plot shows
standardised effect sizes (slopes) of hierarchical Bayesian models of phenological events (purple), seasonal
temperature (red), seasonal precipitation (green), NDVI (yellow), minimum sea ice extent, sea ice concentration, and
snowmelt date (all blue), on growth. The centre line is the effect and error bars are 95% credible intervals. For ease of
comparison between effect sizes, explanatory variables in this analysis are variance-scaled from -1 to 1